

Composition of the Earth and its reservoirs: Geochemical observables

Cin-Ty A. Lee

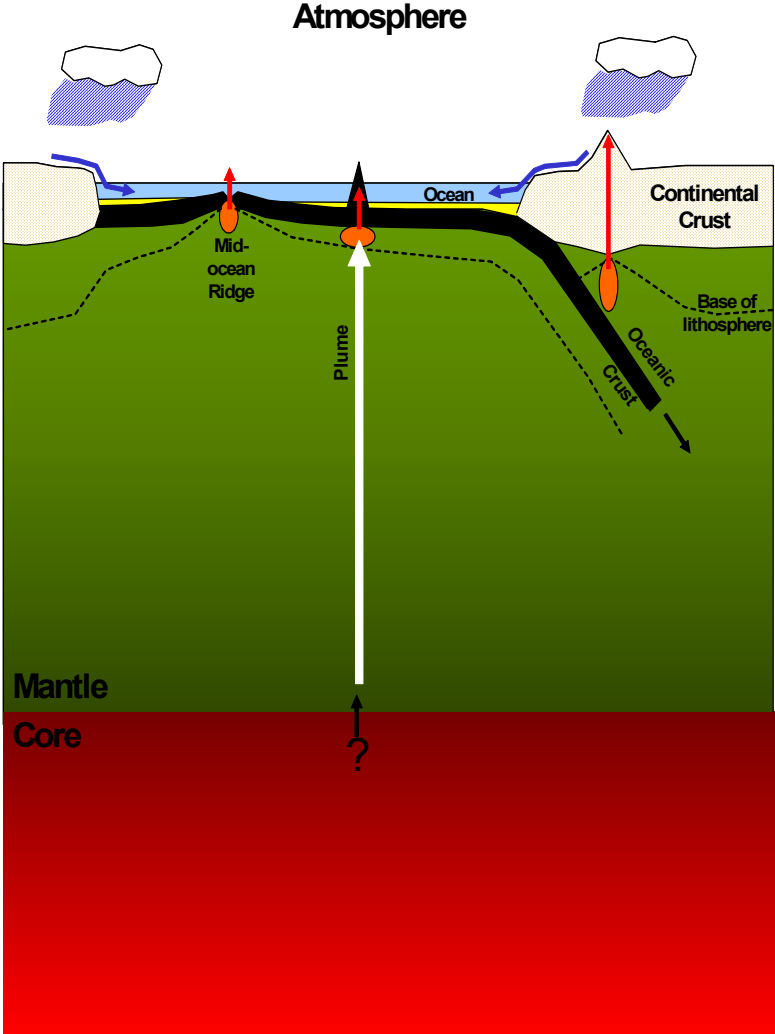
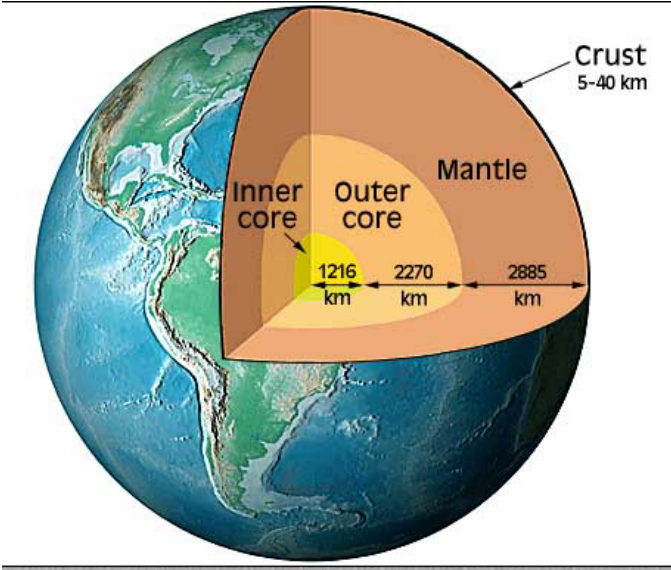
Rice University



RICE

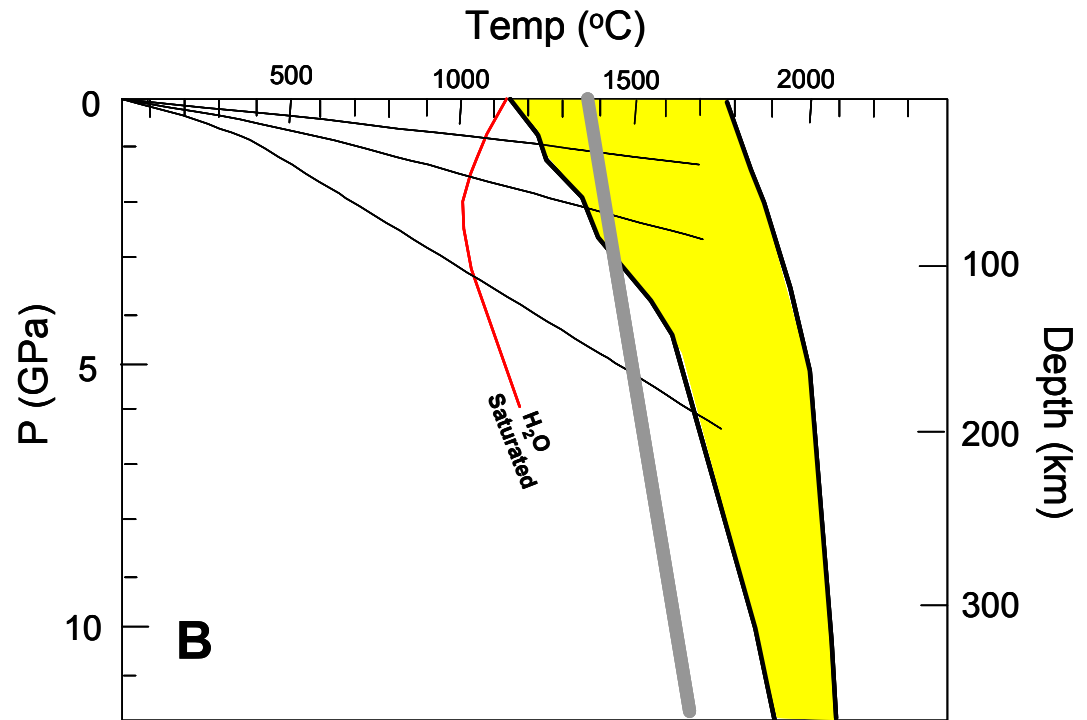
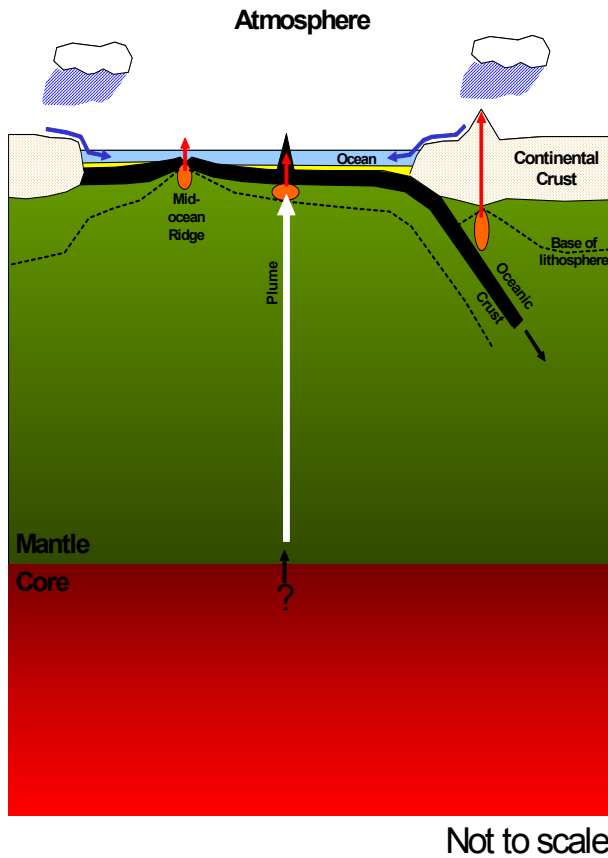
MYRES-I 2004

The Earth is dynamic and heterogeneous

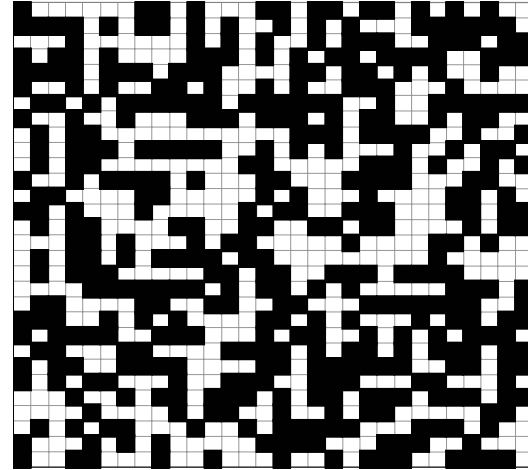
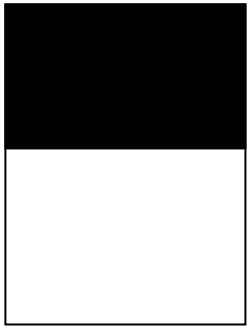


Not to scale

The Earth is differentiating: But are we mixing or unmixing reservoirs?



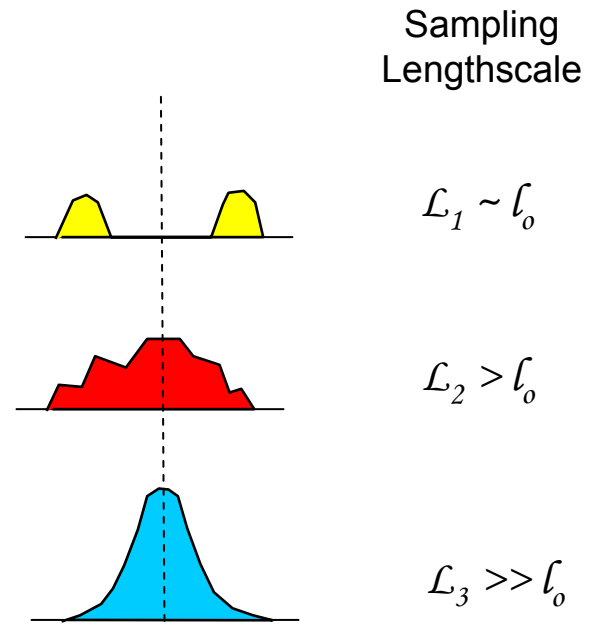
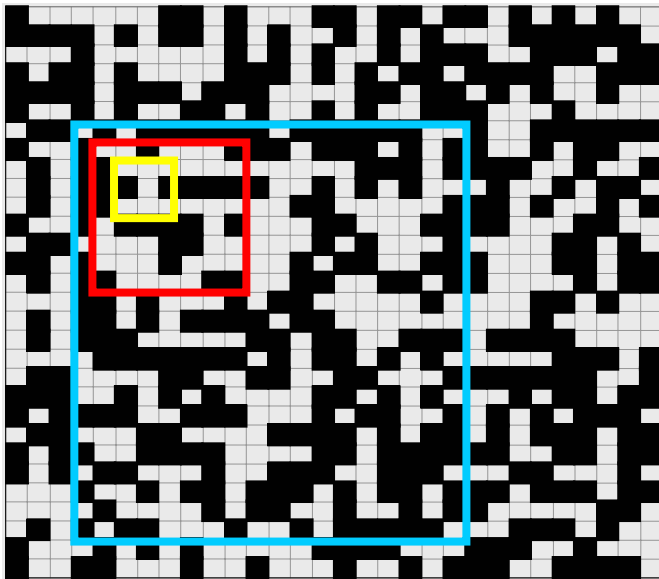
Definition of heterogeneity



Heterogeneity: a qualitative term describing how “well-mixed” a system is

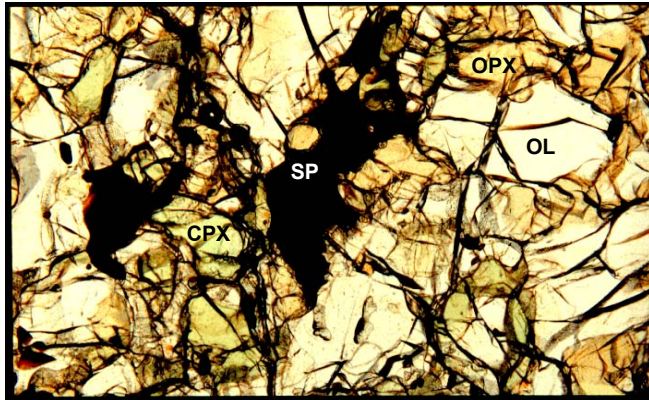
Reservoirs: chemically distinct regions in a system that have physically defined boundaries

At what point do we consider a heterogeneity a reservoir?

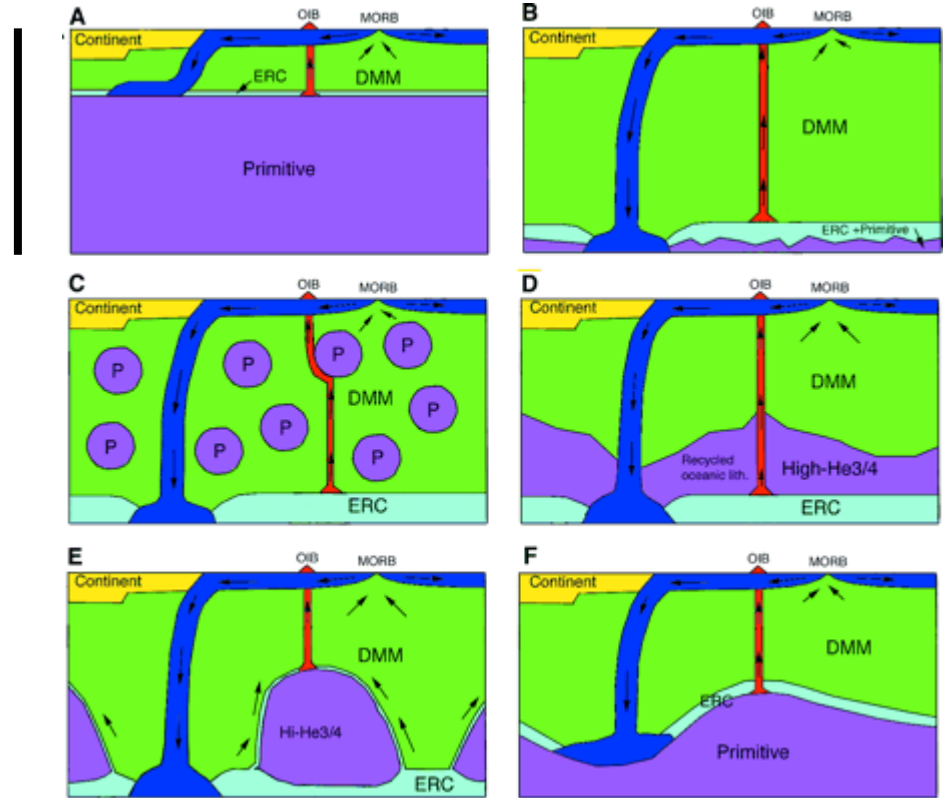


Tell me what lengthscale you're interested in...

3000 km



1 cm



Tackley, 2000

Forms of compositional heterogeneities

Major - wt. %

Minor - ~ 0.1 wt. %

Trace - < 100 ppm

Isotopic

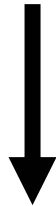
What types of compositional heterogeneities lead to variations in physical parameters?

Three Steps to Bliss

1. Bulk Earth Composition



2. Composition of reservoirs



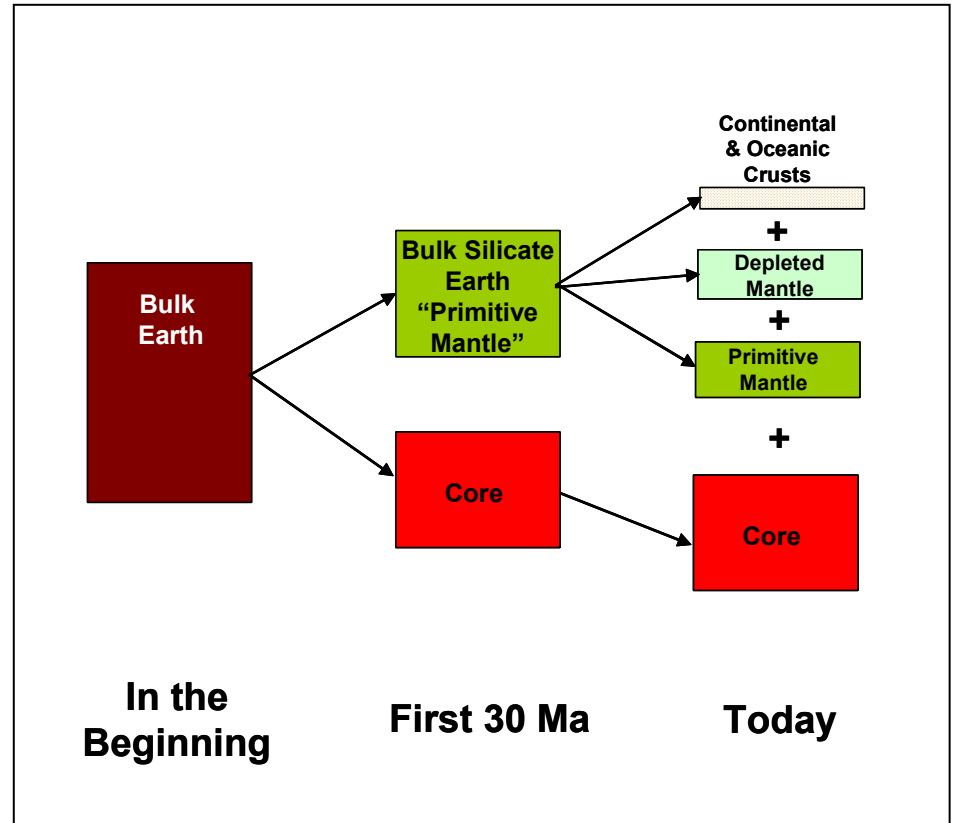
3. Quantifying the size and distribution of reservoirs

STEP 1 Towards a Bulk Earth Composition

Bulk Silicate Earth
(Primitive Mantle)

+

Core



What types of samples can we work with?

Mantle rocks (xenoliths, massifs, ophiolites)

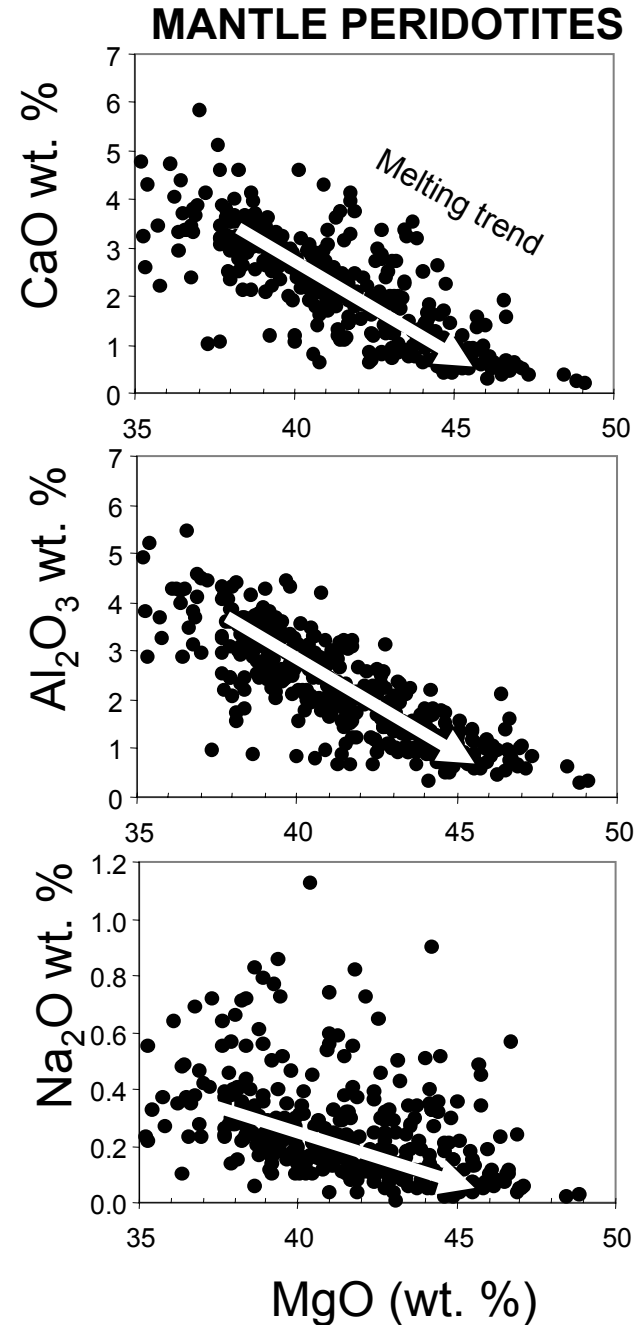
Lavas/Magmas

Sediments

Meteorites?

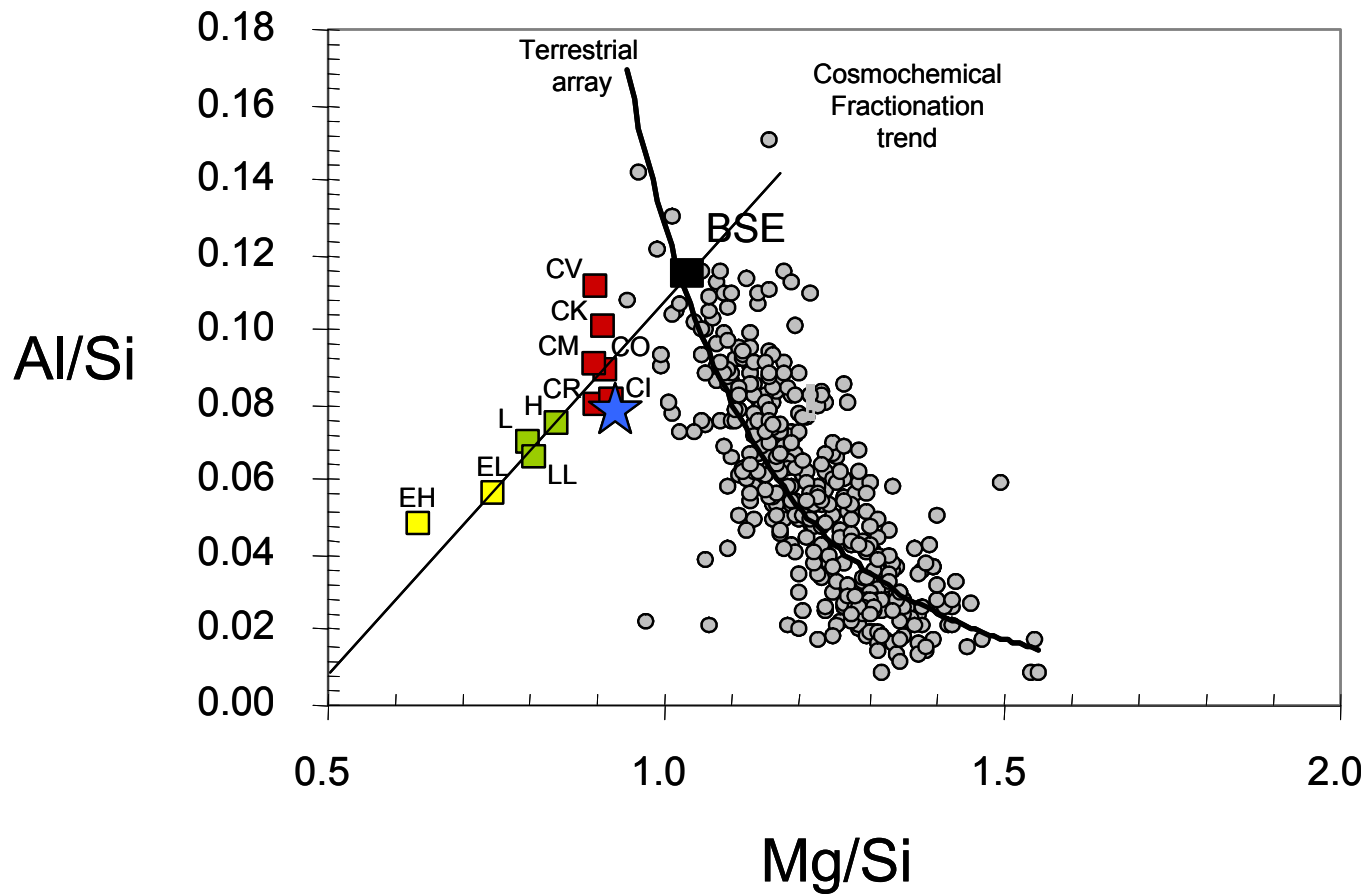
The search for the holy grail

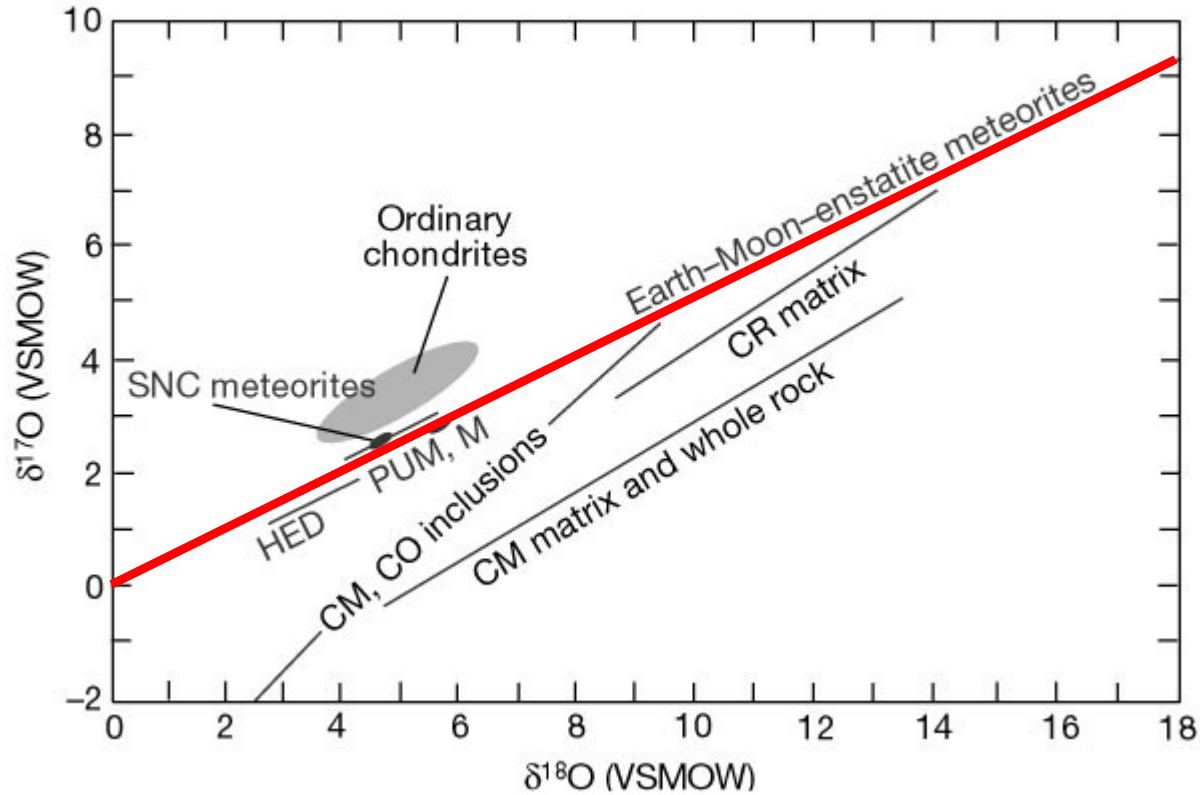
Do samples of “Primitive Mantle” exist?



**Can we use “primitive” meteorites, e.g.
undifferentiated meteorites as a proxy for the
undifferentiated Earth?**

Yes and No



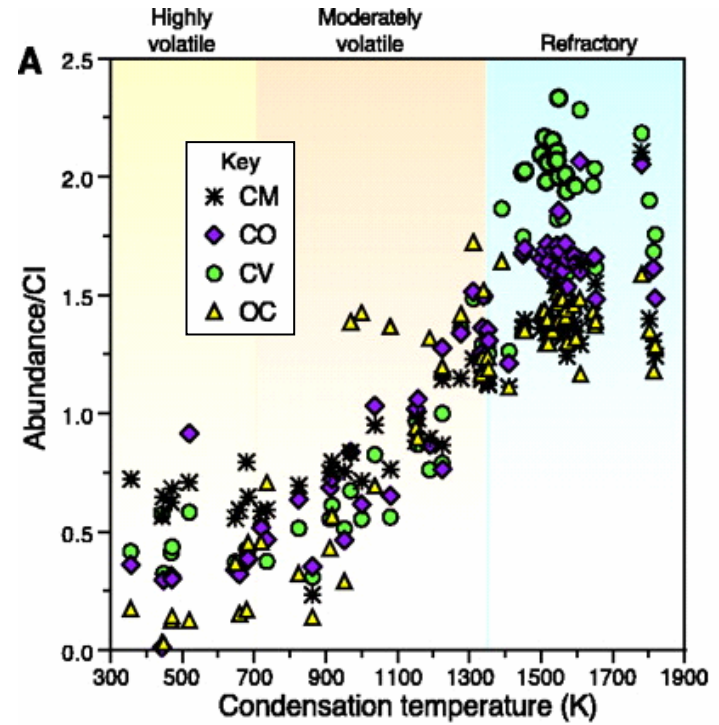
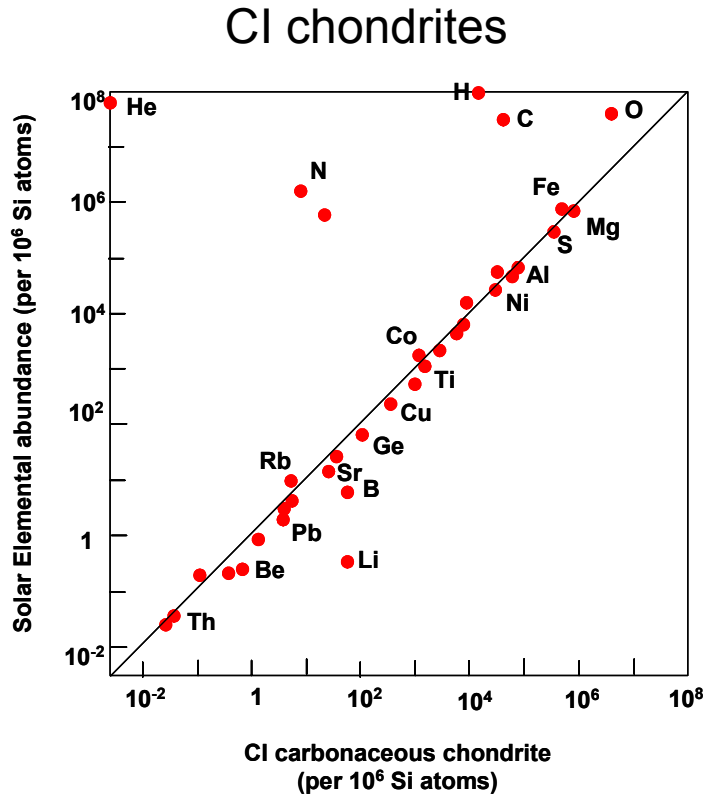


Drake and Righter, 2002; after:

- 4 Clayton, R. N. & Mayeda, T. K. Oxygen isotope studies in
0 carbonaceous chondrites. *Geochim. Cosmochim. Acta* **63**, 2089-
. 2104 (1999).
- 4 Clayton, R. N. Oxygen isotopes in meteorites. *Annu. Rev. Earth
1 Planet. Sci.* **21**, 115-149 (1993).
- .
- 4 Clayton, R. N., Mayeda, T. K., Goswami, J. N. & Olsen, E. J.
2 Oxygen isotope studies in ordinary chondrites. *Geochim.
. Cosmochim. Acta* **55**, 2317-2337 (1991).

**It is safe to say that
Earth is derived
from Earth-like
materials**

... but all is not lost



The Early Evolution of the Inner Solar System: A Meteoritic Perspective

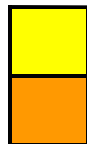
C. M. O'D. Alexander, A. P. Boss, R. W. Carlson

SCIENCE, Volume 293, Number 5527, pp. 64-68

50% condensation temperatures (pressure 1E-4 bars)

| | | | | | | | | | | | | | | | | | |
|------------|------------|------------|------------|----|------------|------------|------------|------------|------------|------------|-----------|------------|------------|------------|-----------|----------|----|
| H | | | | | | | | | | | | | | | | | He |
| Li 1225 | Be | | | | | | | | | | | B 964 | C | N | O | F 736 | Ne |
| Na 970 | Mg 1340 | | | | | | | | | | | Al 1680 | Si 1311 | P 1267 | S 648 | Cl | Ar |
| K 1000 | Ca 1520 | Sc 1644 | Ti 1590 | V | Cr 1300 | Mn 1190 | Fe 1336 | Co 1351 | Ni 1354 | Cu 1037 | Zn 660 | Ga 997 | Ge | As 1157 | Se 684 | Br | Kr |
| Rb 1080 | Sr | Y | Zr 1750 | Nb | Mo 1600 | Tc | Ru 1600 | Rh | Pd 1334 | Ag 952 | Cd | In 470 | Sn 720 | Sb 912 | Te 680 | I | Xe |
| Cs | Ba | La-Lu | Hf | Ta | W 1800 | Re 1800 | Os 1800 | Ir 1600 | Pt 1411 | Au 1225 | Hg | Tl 428 | Pb 496 | Bi 451 | Po | At | Rn |
| Fr | Ra | Ac-Lr | | | | | | | | | | | | | | | |

| | | | | | | | | | | | | | | |
|------------|------------|----|-----------|----|----|------------|----|----|----|----|----|----|------------|------------|
| La 1500 | Ce | Pr | Nd | Pm | Sm | Eu 1290 | Gd | Tb | Dy | Ho | Er | Tm | Yb 1420 | Lu 1590 |
| Ac | Th 1590 | Pa | U 1540 | Np | Pu | Am | Cm | Bk | Cf | Es | Fm | Md | No | Lr |



Refractory, >1400 K

Transitional ~1250-1350 K



Moderately volatile ~800-1250 K

Highly volatile <800 K

Nebular differentiation

Refractory

Moderately volatile

Volatile

Planetary differentiation

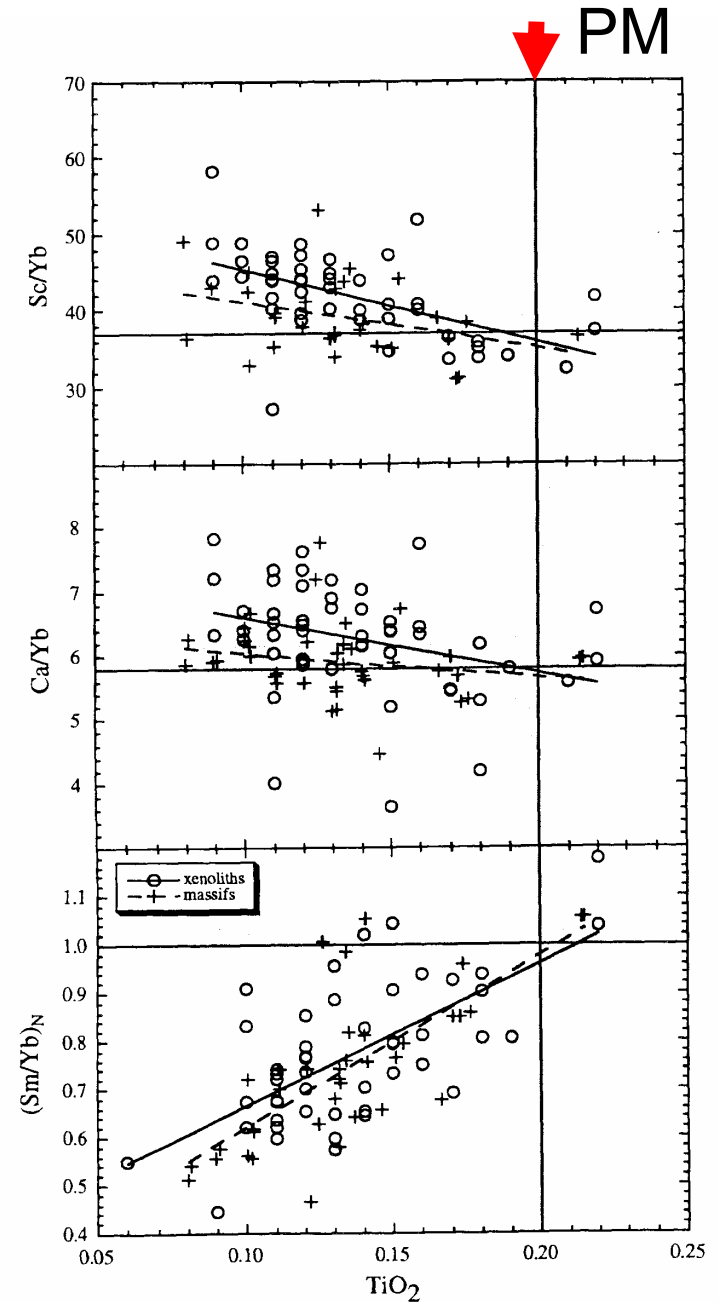
Lithophile – silicate loving

Siderophile – Fe loving

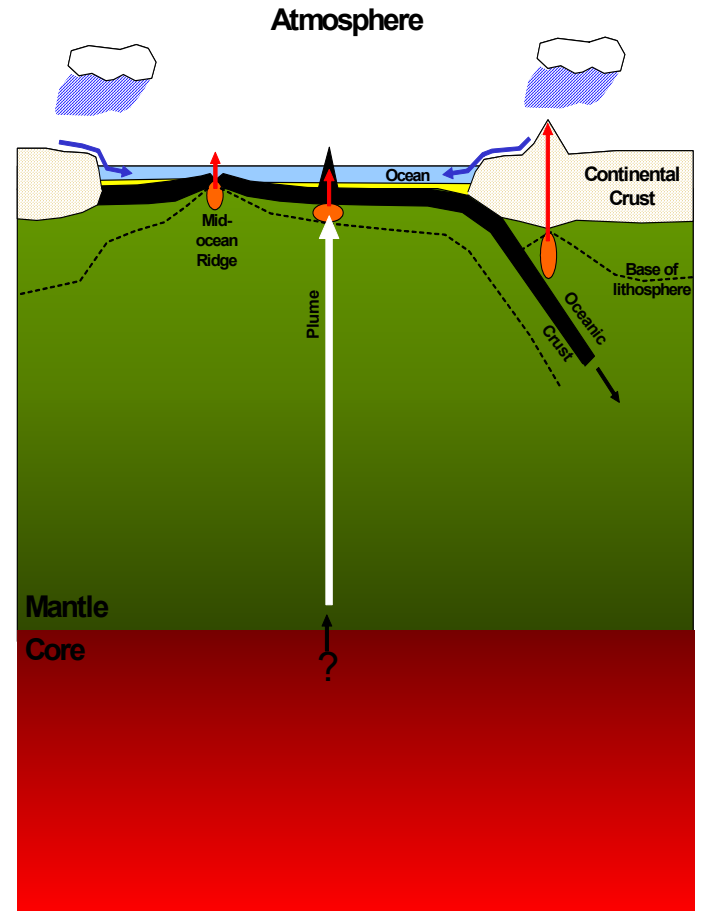
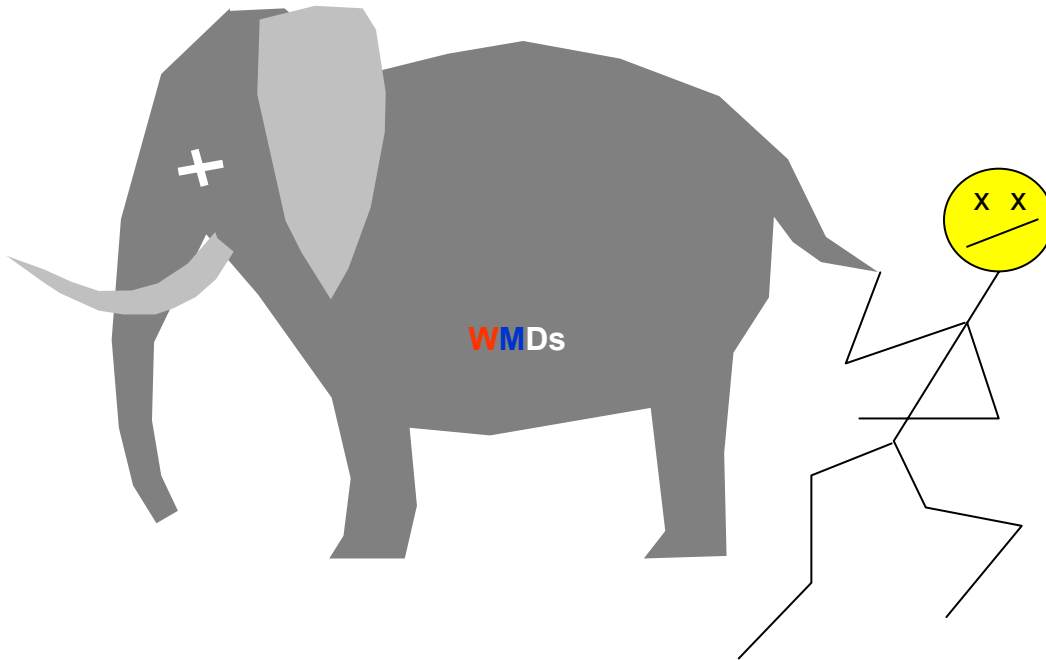
Atmophile - atmosphere

Establish concentrations in
“Primitive Mantle” using
refractory lithophile element
ratios for chondrites

Fig. 6 from
McDonough & Sun (1995)

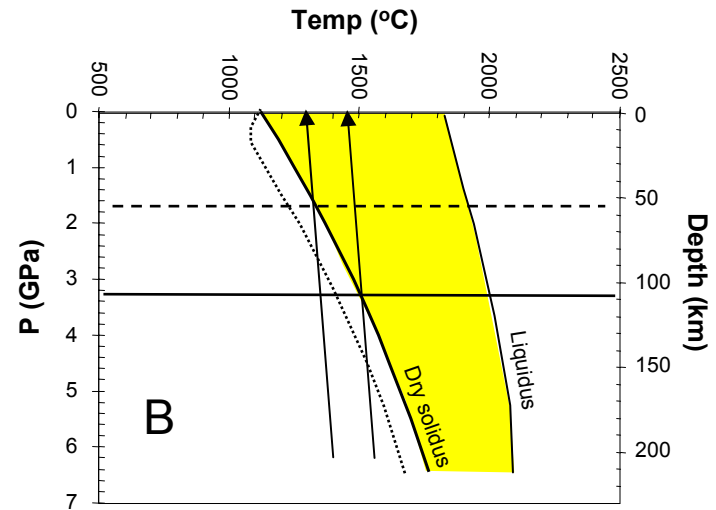
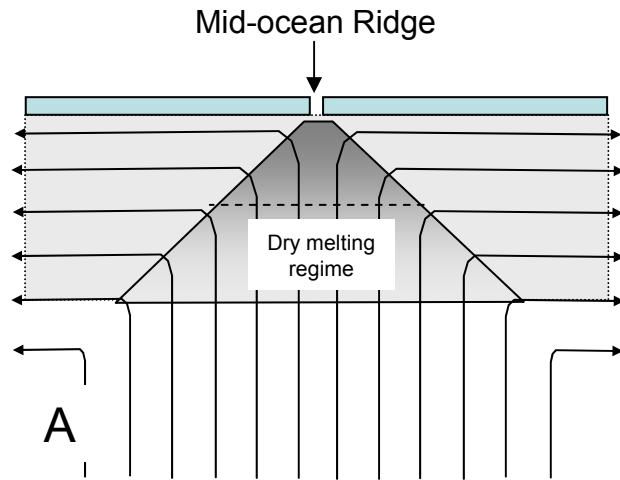


STEP 2. RESERVOIR DOGS

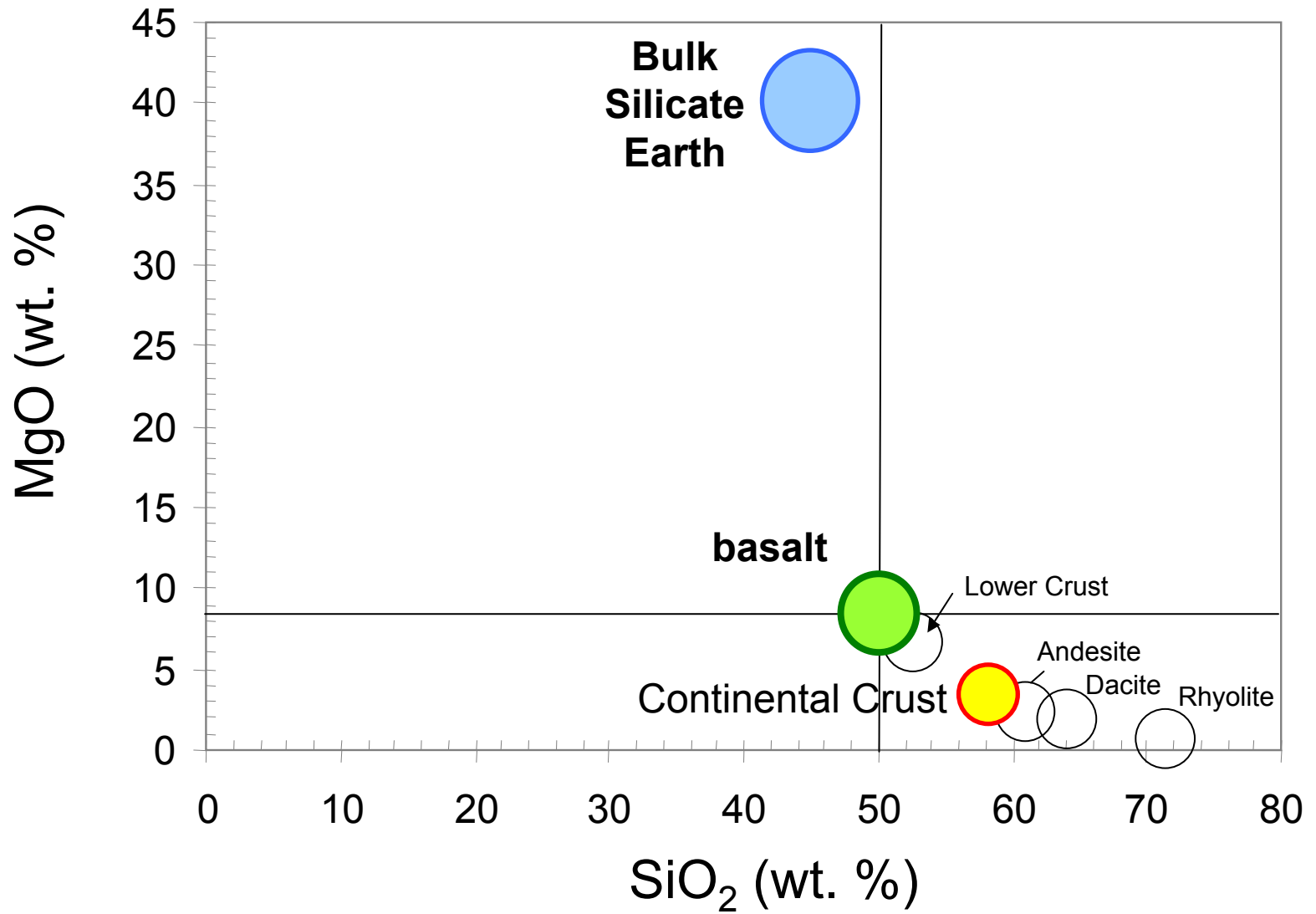


Not to scale

Partial melting as a major differentiation process



Major Element Effects



Partition coefficient

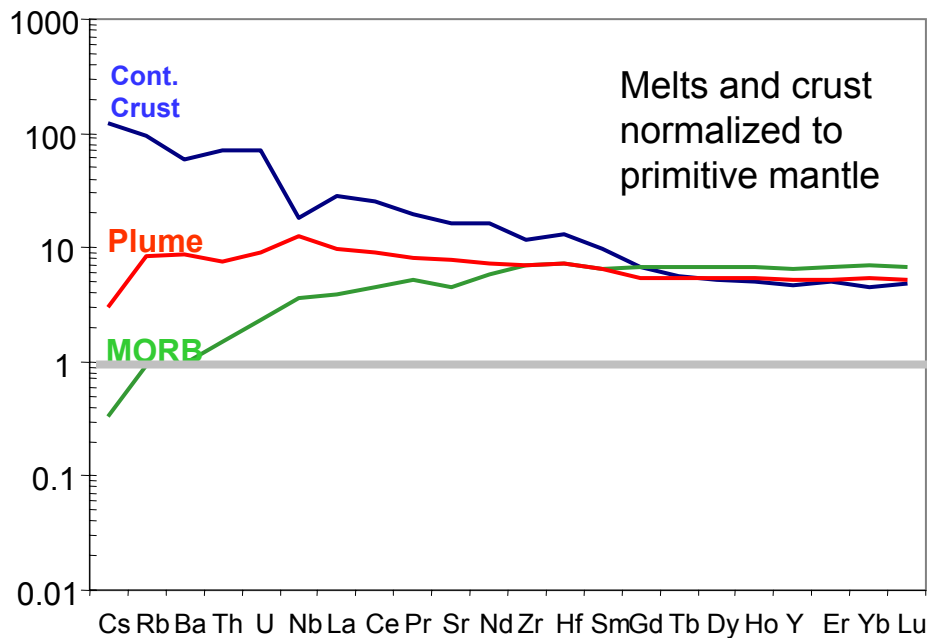


$$D = \frac{C_{\text{solid}}}{C_{\text{melt}}}$$

$D > 1$ compatible in solid

$D < 1$ incompatible in solid

Using magmas as “windows” to the mantle



Primitive
Mantle

MORB source is **depleted** in highly incompatible elements (DMM = depleted MORB mantle)

Continental Crust is enriched in highly incompatible elements

Plume is enriched or more primitive in character

incompatible



More compatible

Increasing D

Trace-elements may suffer from
fractionation during magmatic processes

Isotopes in general are NOT fractionated

Sm/Nd < 1

melt



Rb/Sr > 1

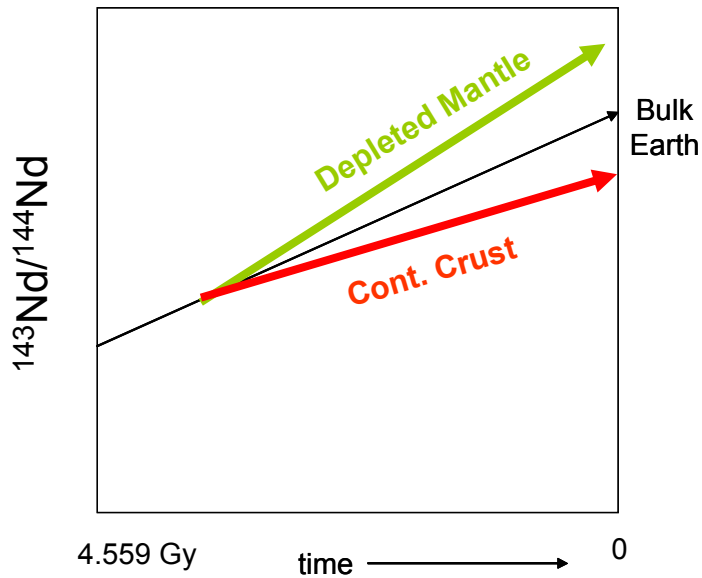
Sm/Nd > 1



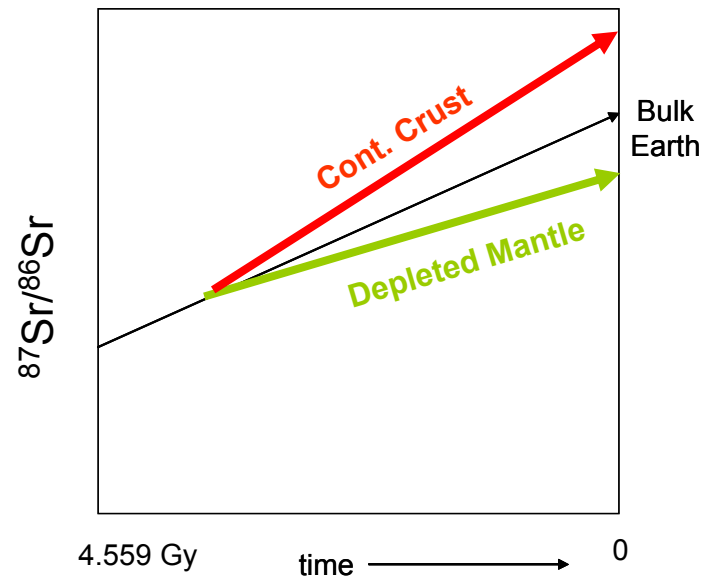
Rb/Sr < 1

Solid residue

$^{147}\text{Sm} - ^{143}\text{Nd}$

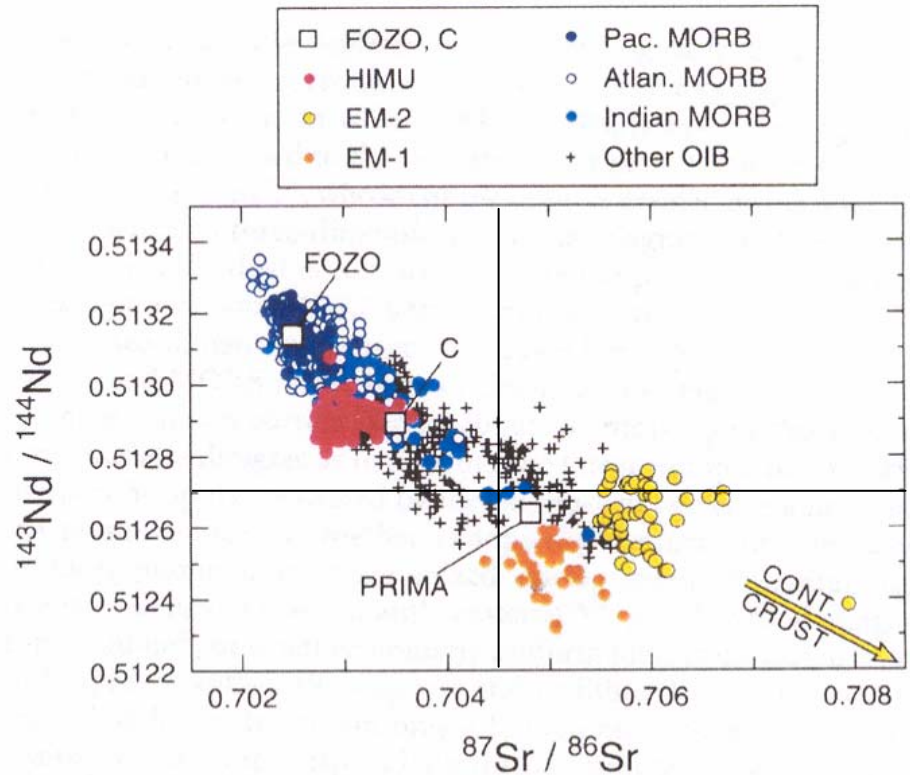


$^{87}\text{Rb} - ^{87}\text{Sr}$



Depleted Mantle
radiogenic Nd
unradiogenic Sr

Continental Crust
unradiogenic Nd
radiogenic Sr



Hofmann, 1997
Nature 385: 219-229

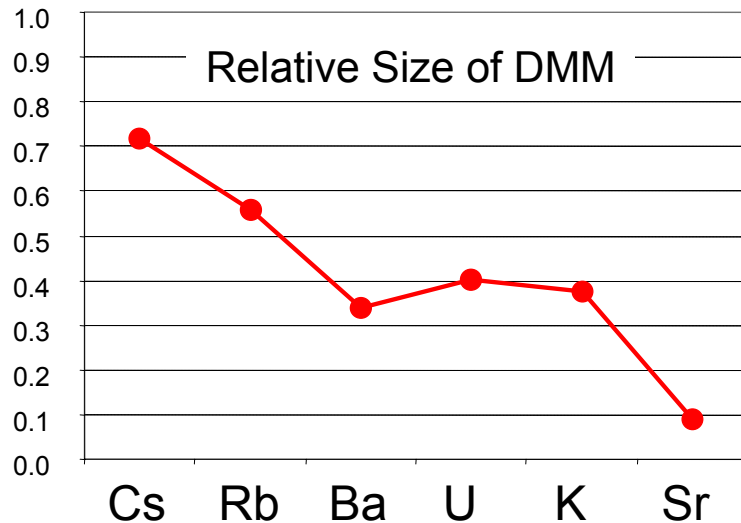
At least the upper part of the Earth's mantle is depleted in highly incompatible elements

This depleted portion appears to be complementary to the continental crust

How much of the mantle is depleted?

Mass Balance Magic

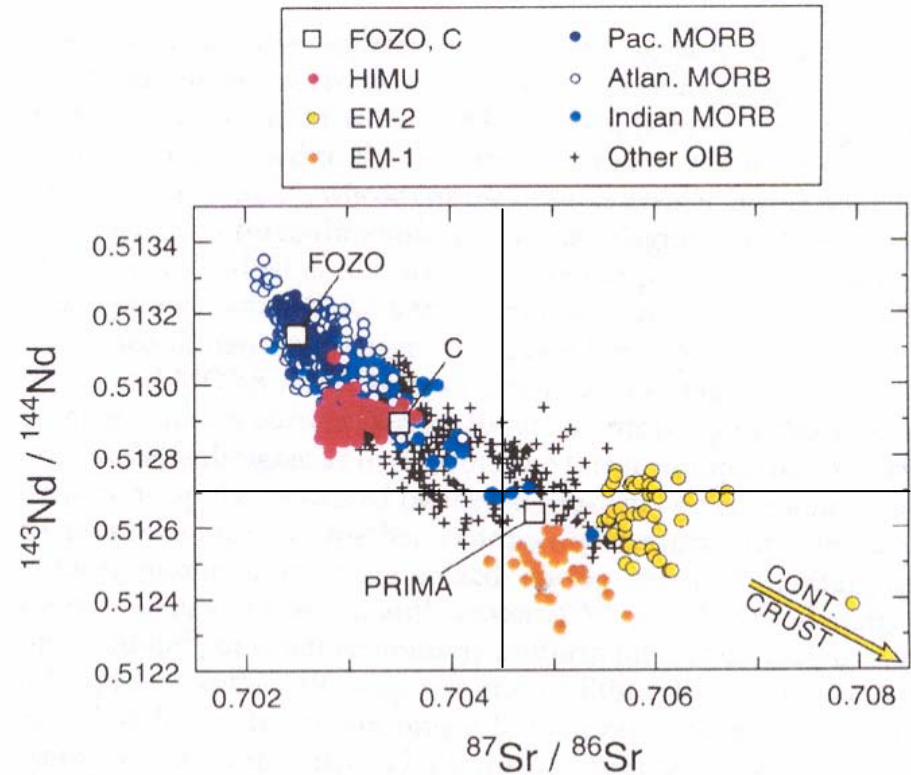
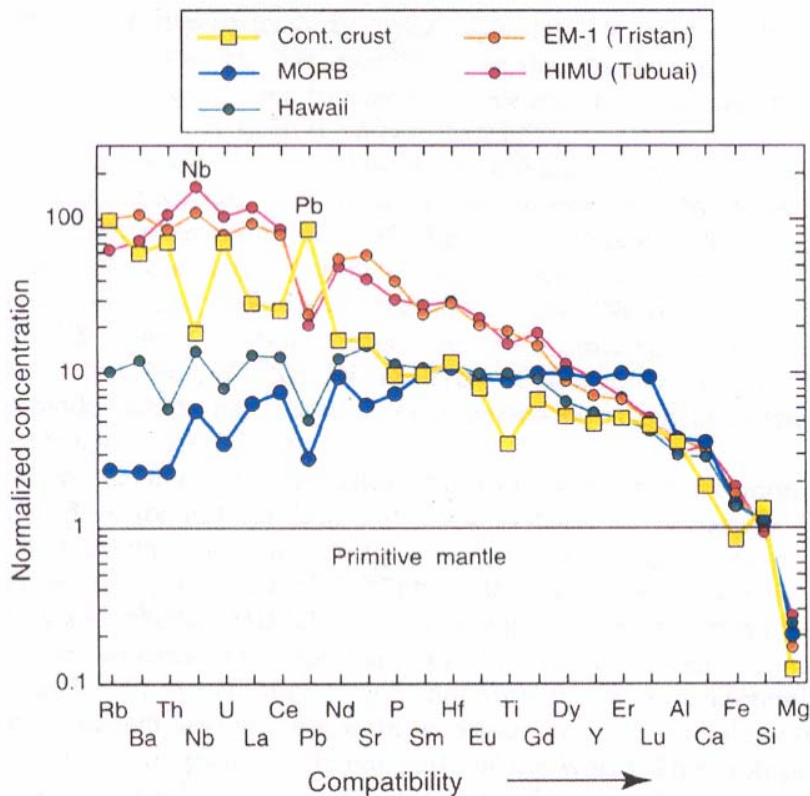
| | Concentration (ppm) | | | | DMM/PM | Relative Size of DMM |
|----|---------------------|-------|--------|--------|--------|----------------------|
| | Cont. Crust | MORB | PM | DMM | | |
| Cs | 2.6 | 0.007 | 0.021 | 0.0007 | 0.0 | 0.7 |
| Rb | 58 | 0.56 | 0.6 | 0.056 | 0.1 | 0.6 |
| Ba | 390 | 6.3 | 6.6 | 0.63 | 0.1 | 0.3 |
| U | 1.42 | 0.047 | 0.0203 | 0.0047 | 0.2 | 0.4 |
| K | 15606 | 600 | 240 | 60 | 0.3 | 0.4 |
| Sr | 325 | 90 | 19.9 | 9 | 0.5 | 0.1 |



Possibly 30 to 70% of mantle has been processed to make Cont. Crust

Mass Balance Says **NOTHING** about geometry

What do hotspots tell us? (Ocean Island Basalts – OIB)



Hofmann, 1997
Nature 385: 219-229

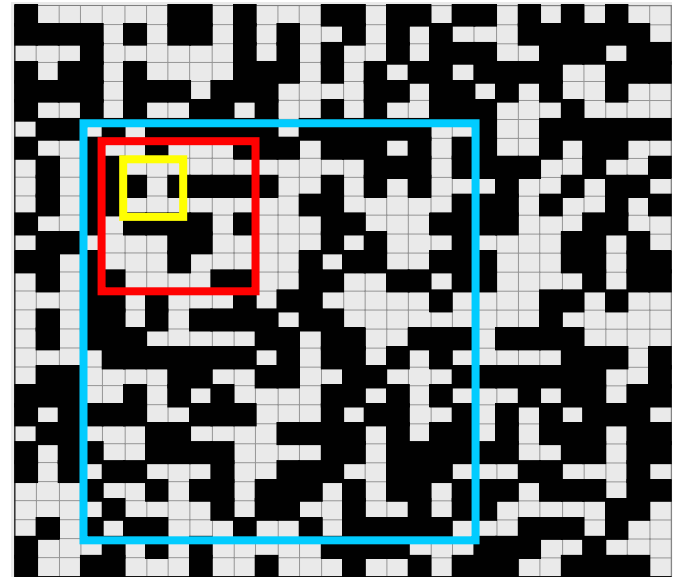
Hotspot magmas appear to be variably enriched in incompatible elements

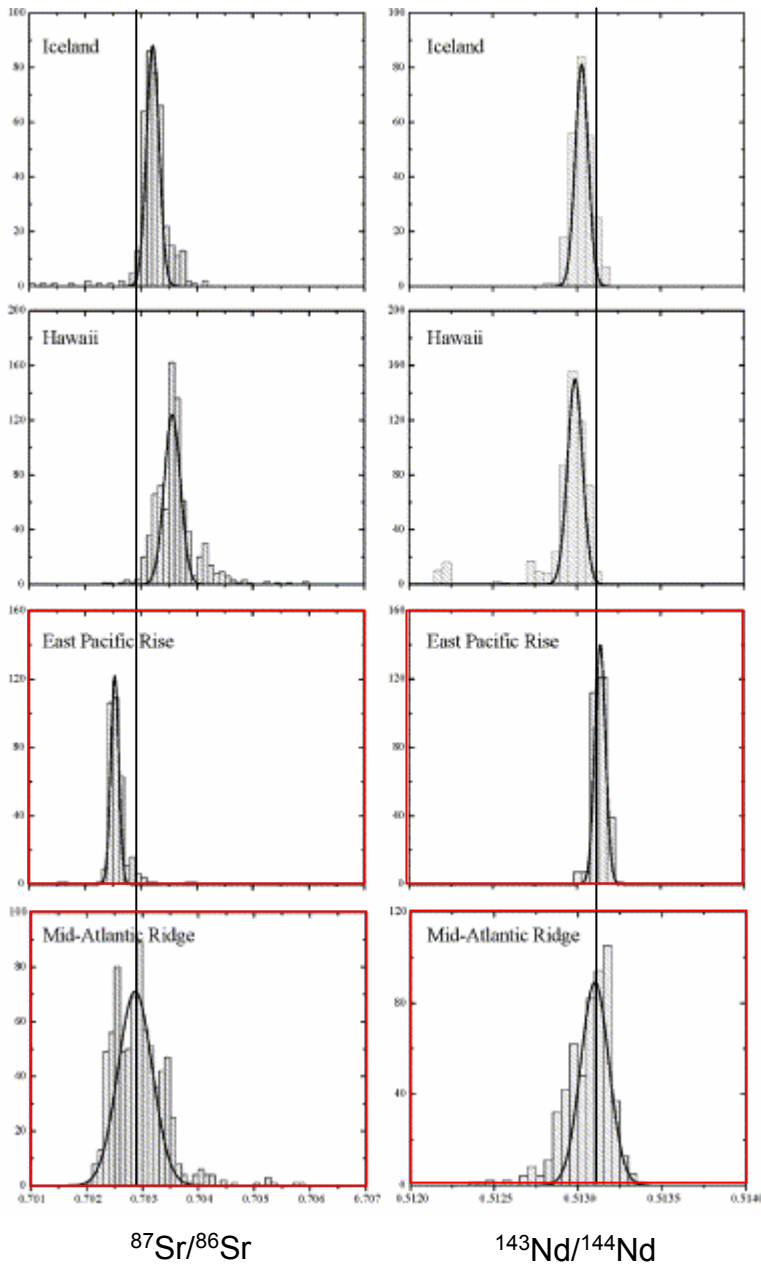
Source feature



Or

Sampling problem?





Meibom & Anderson 2003

Sampling problem is a possibility

BUT average isotopic composition of some OIBs differ from that of MORBs



OIB source must differ slightly in composition from that of MORB source

IF OIBs are associated with plumes +
IF plumes derive from the lower mantle

The lower mantle must be slightly enriched.

Do plumes exist?

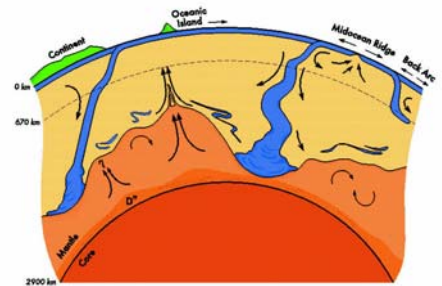
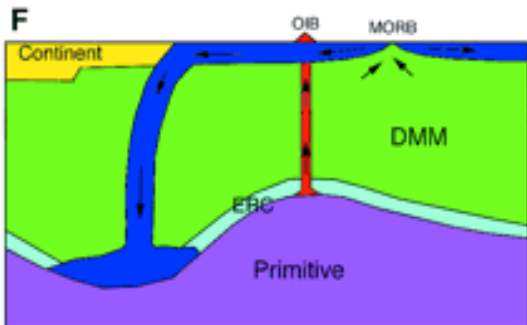
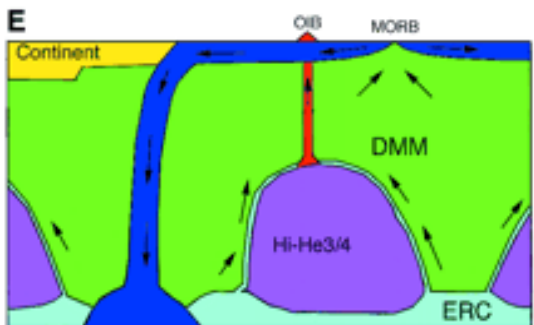
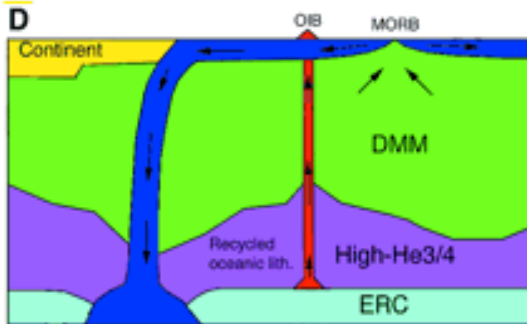
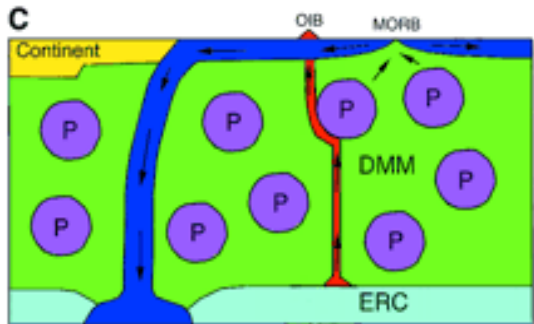
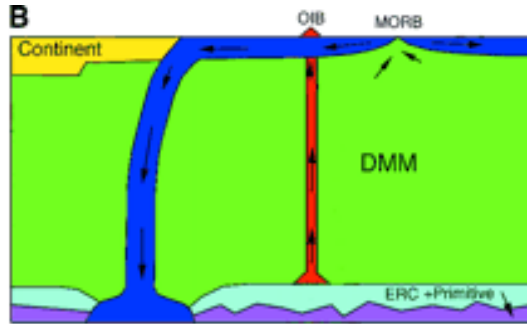
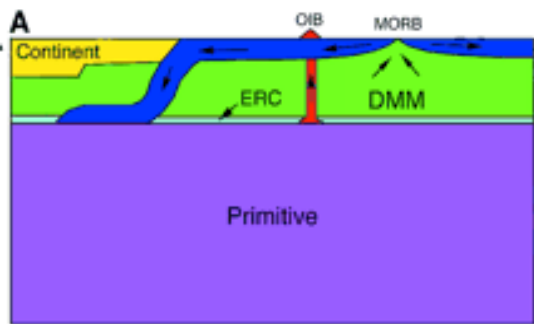
Other Reservoirs in the Silicate Earth

Subducted oceanic crust

SCLM – subcontinental lithospheric mantle

Subducted sediment

Primordial mantle???



Tackley, 2000

Where do we go from
here?

More of the same data?

New data?

NEW QUESTIONS?

YOU can be an armchair geochemist

Use compiled datasets

GEOROC

RidgePetDB

GERM

But don't abuse the data sets

New directions  **New data**

DIRECTION 1

Is a homogeneous Bulk Silicate Earth a valid assumption?

Was the primordial mantle stratified? If so, how does this affect geochemical mass balance conditions?

Is the lower part of the mantle Fe-rich?

Can we ignore mass fluxes across the core-mantle boundary?

What data may be able to answer these questions?

Short-lived radionuclides

^{182}Hf - ^{182}W , ^{146}Sm - ^{142}Nd , ^{107}Pd - ^{107}Ag

More precise data on first series transition elements

Fe, Mn, Ni, Co, Cr, V, Sc

Direction 2

Constrain the size and distribution of reservoirs

Geochemistry, Geophysics, Geodynamics, Petrology

G³P

Do geochemical heterogeneities reflect variations in physical properties that are detectable by geophysical methods?

Clearly, major-element variations affect physical properties of the mantle.

The physical effects of trace-element and isotopic heterogeneities are unclear. However, trace-element and isotopic heterogeneities may be correlated with variations in **oxygen fugacity** and **water content**, both of which may have profound effects on elasticity, conductivity, and viscosity.



Future?

- Effects of fO_2 and H_2O on physical properties of peridotite
- Characterizing the variation of fO_2 and H_2O in the mantle and what processes control these parameters

Geochemical research on

partitioning behavior of redox sensitive elements
valence state of redox-sensitive elements
quantification of H_2O content in the mantle

Conclusions

Geochemistry has provided a considerable knowledge base for our understanding of how the Earth works.

Further progress will require interdisciplinary collaboration and the generation of new questions and areas of focus.

More geochemical data will obviously be helpful, but existing databases should be mined to reveal areas that need more refinement.