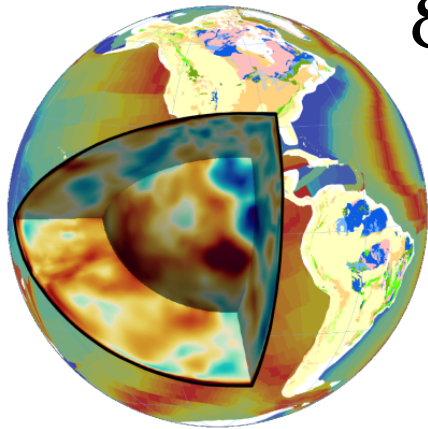


Geodynamics I: Subduction zone dynamics and global mantle flow



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The University of Texas at Austin

CIDER Summer School
Subduction Zone Dynamics



Berkeley CA
July 2017



INSTITUTE FOR GEOPHYSICS



TEXAS Geosciences

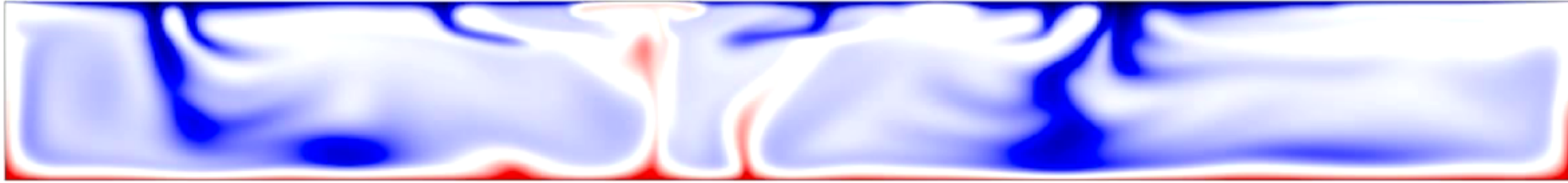
The University of Texas at Austin
Jackson School of Geosciences

Contents

- Plate driving forces
- Some fluid dynamics
- Global models
- Regional models
- Conveyor belts

Plate tectonics is the top boundary layer of mantle convection

Subduction: cold plume dynamics



fluid heated from below – Rayleigh-Benard convection

computation by A. McNamara

- Rayleigh number, Ra , controls the vigor of convection
- Symmetry of upwellings and downwellings broken by
 - *temperature dependence of viscosity*
 - *depth-dependent viscosity*
 - internal vs. bottom heating
 - fractionation (e.g. continents and thermo-chemical piles)

Top thermal boundary described by half space cooling

$$\mathbf{q} = -k\nabla T$$

Fick's law for heatflow

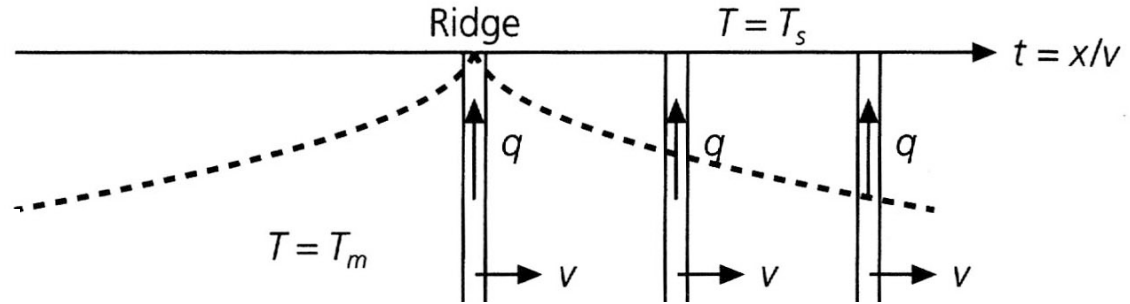
$$\frac{\partial T}{\partial t} = \kappa \frac{\partial^2 T}{\partial z^2}$$

thermal diffusion in 1-D
(no heat sources)

$$T = T_m \quad \text{at } x = 0$$

$$T = 0 \quad \text{at } z = 0$$

$$T = T_m \quad \text{at } z \rightarrow \infty$$

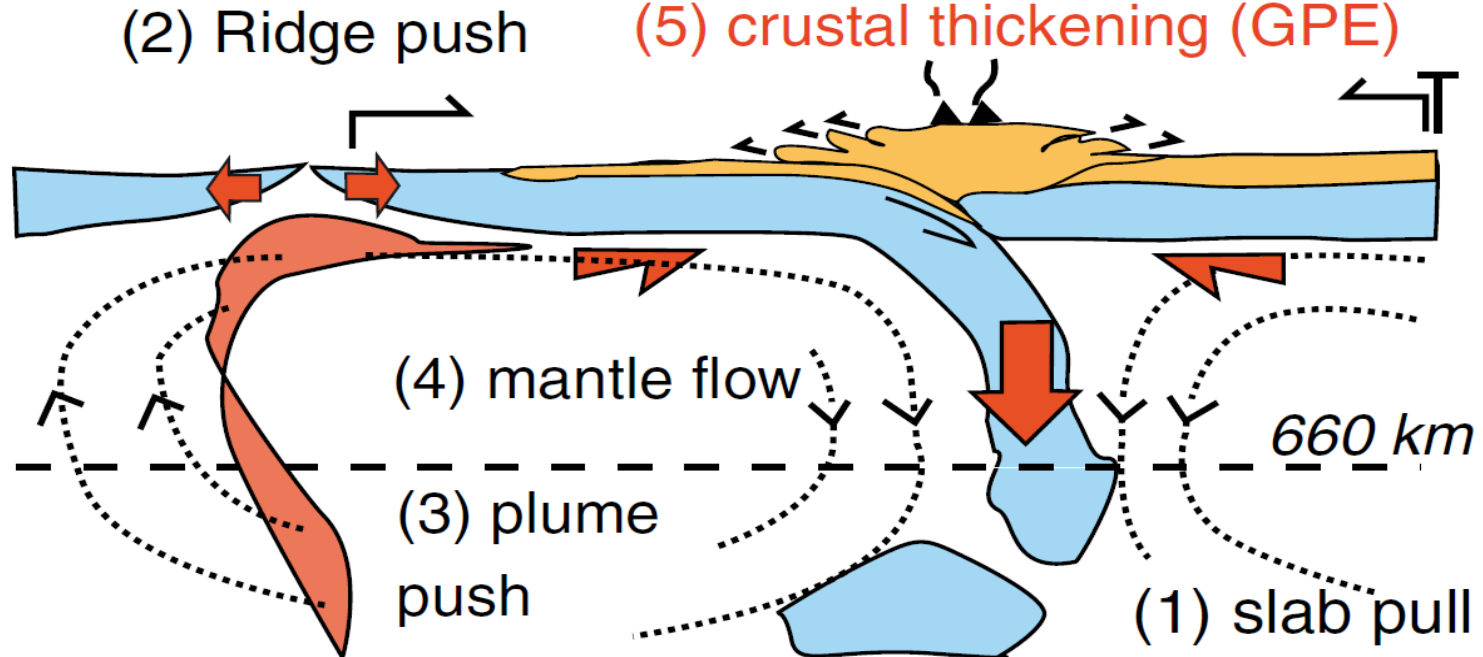


κ = thermal diffusivity
length $\sim \sqrt{(\kappa \times \text{time})}$

$$T(x, z) = T_m \operatorname{erf} \left(\frac{z}{2\sqrt{\kappa t(x)}} \right)$$

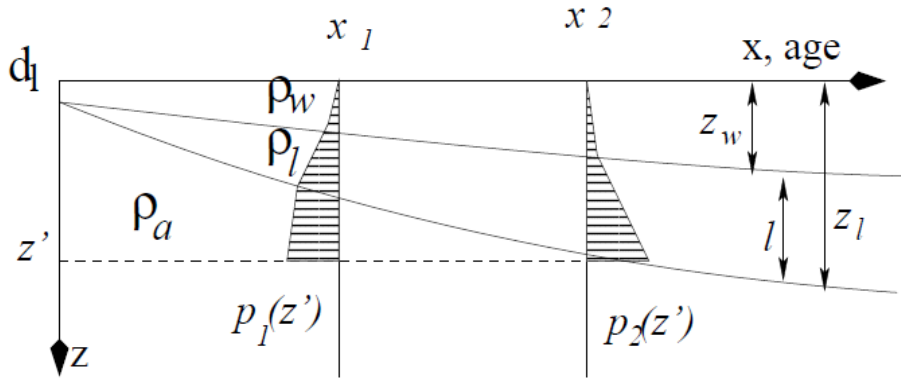
Integrals over components of thermo-chemical convection

(5) crustal thickening (GPE)



Lithospheric thickening

(AKA *ridge push*, oceanic GPE)



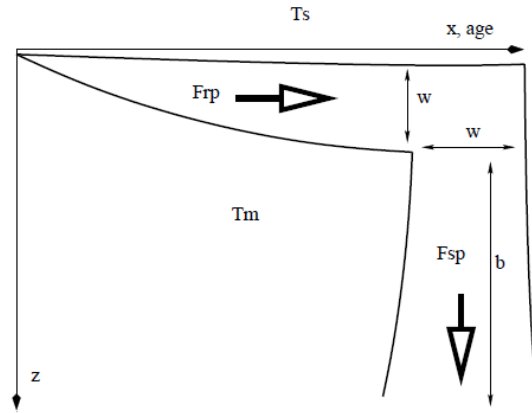
$$l = a(z_w - d_l) = a\Delta z_w$$

$$a = \frac{\rho_w - \rho_a}{\rho_a - \rho_l} \quad (\text{from isostasy})$$

$$z_w(x) = c \sqrt{\frac{x}{u_0}}$$

$$F = \frac{ac^2}{2u_0} g \Delta \rho \Delta x$$

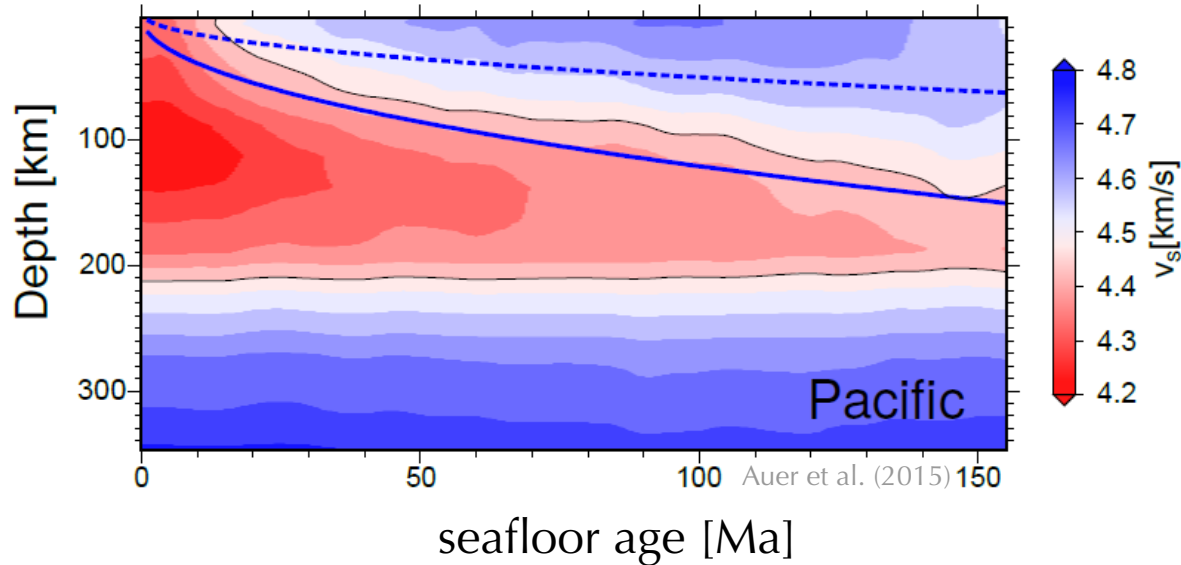
Subducting thermal boundary layer: *Slab pull*



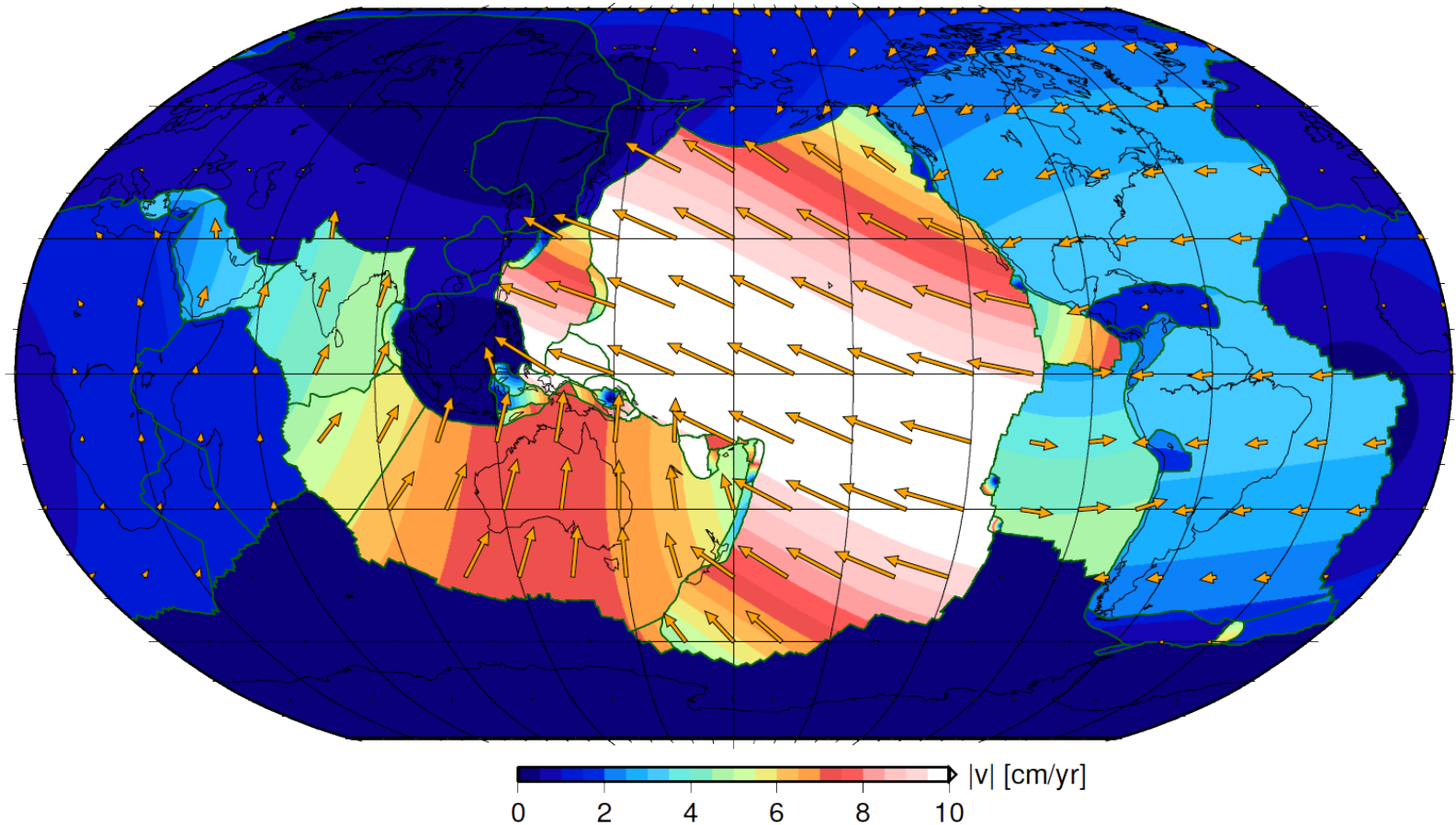
$$F_b = -2\rho_0 g \alpha b (T_m - T_s) \sqrt{\frac{\kappa t_s}{\pi}}$$

Force estimates from half-space cooling

- ridge push (lithospheric thickening) $\sim 10^{12}$ N/m
- slab pull $\sim 10^{13}$ N/m
- worried about tectonics? worry about subduction

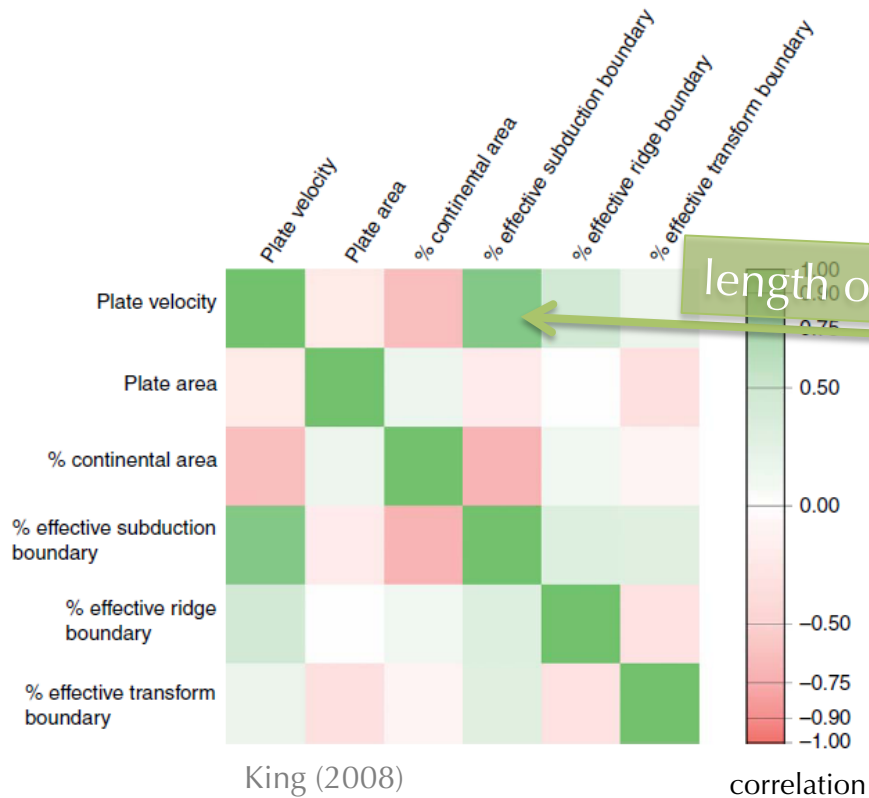


How to explain present-day plate tectonics and test our models?

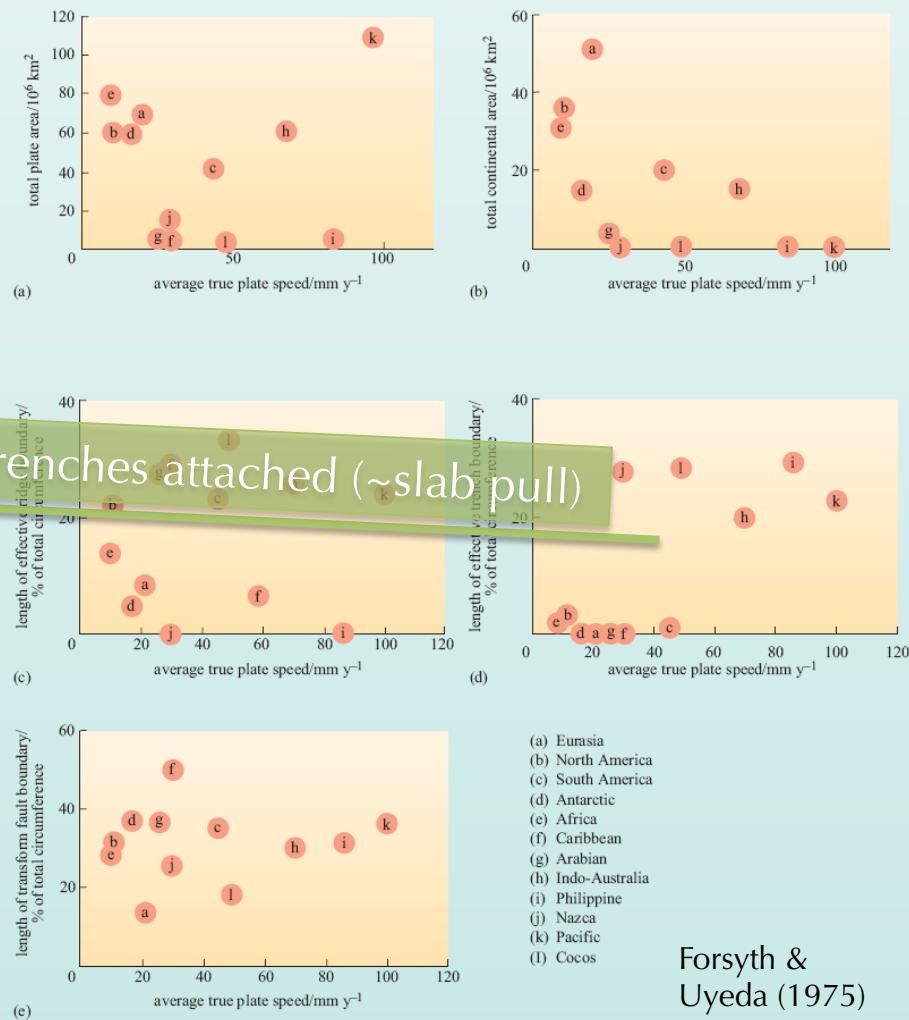


Bird (2002) plate boundaries and MORVEL56 velocities by Argus et al. (2011) in spreading-aligned reference frame (Becker et al., 2015)

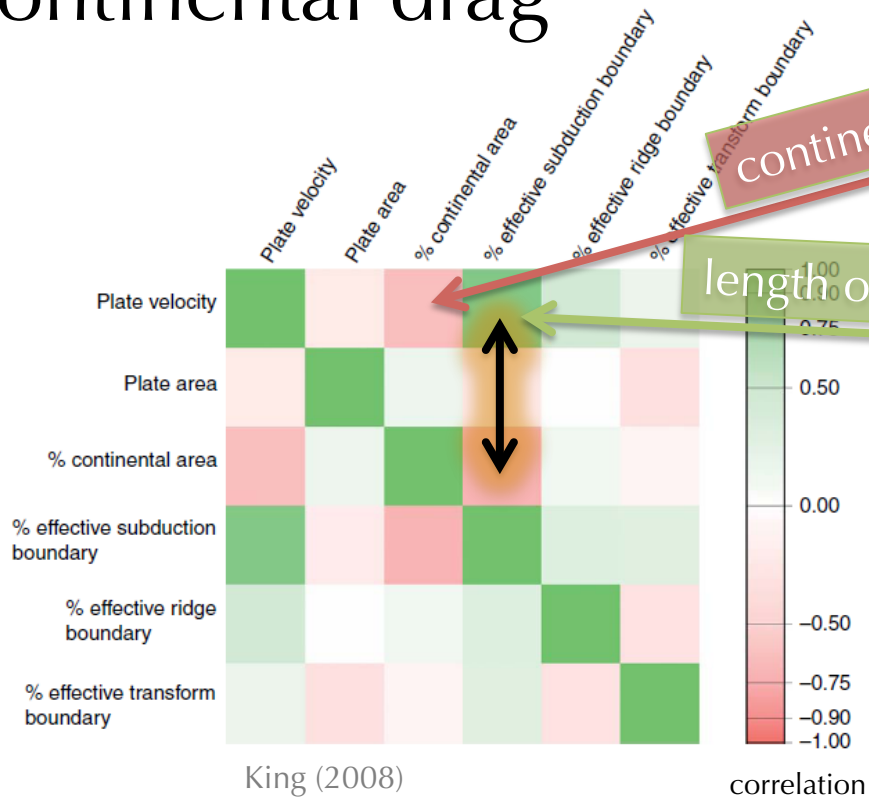
Different natural experiments in space: Plate speed correlation with various geometric/kinematic parameters



length of trenches attached (~slab pull)

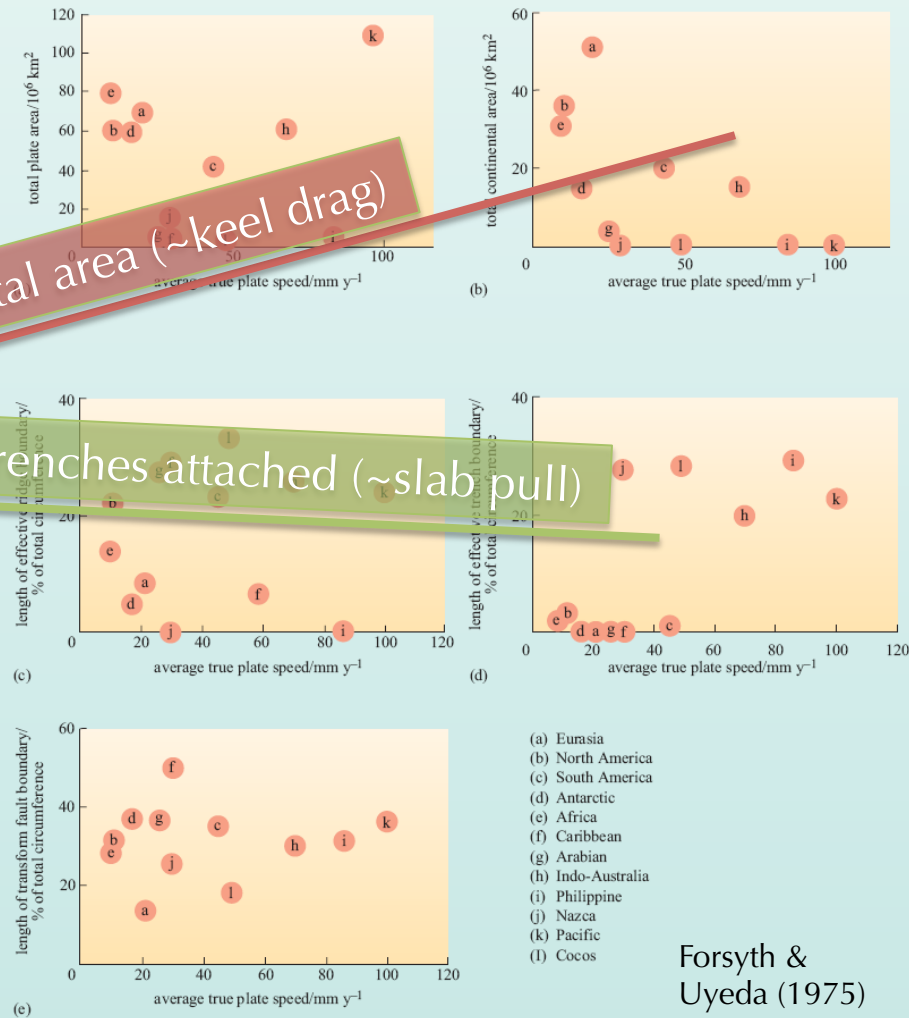


Slab pull vs. continental drag



continental area (~keel drag)

length of trenches attached (~slab pull)



Some fluid dynamics to build our own models

Static force balance (conservation of momentum) in any continuum

$$\frac{\partial \overset{\text{stress tensor}}{\sigma_{ij}}}{\partial x_j} = - \underset{\text{body forces}}{f_i} = - \underset{\text{e.g. buoyancy}}{\Delta \rho g_i}$$

Constitutive law for an incompressible, Newtonian fluid

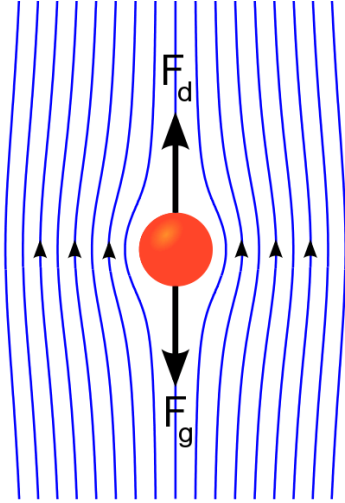
$$\underset{\text{conservation of mass}}{\frac{\partial v_i}{\partial x_i} = 0} \quad \sigma_{ij} = \underset{\text{dynamic pressure}}{-p \delta_{ij}} + \underset{\text{Newtonian viscosity}}{2\eta \dot{\epsilon}_{ij}} = \underset{\text{dynamic pressure}}{-p \delta_{ij}} + \underset{\text{2 x strain-rate}}{\eta \left(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i} \right)}$$

Stokes equation for constant viscosity (neglect inertia, OK at 10^{-25} level ($1/Pr$))

$$2\eta \frac{\partial \dot{\epsilon}_{ij}}{\partial x_j} - \frac{\partial p}{\partial x_i} + \Delta \rho g_i = 0$$

viscous dragdynamic pressure gradientbuoyancy force

Stokes sphere / Stokes sinker

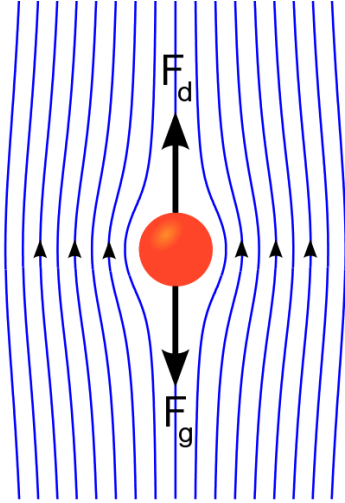


$$F_d = \text{area} \times \text{stress} = \text{area} \times \text{strain-rate} \times \text{viscosity } (\eta)$$
$$= \text{area} \times \text{velocity } (v) / \text{radius } (a) \times \text{viscosity}$$

$$F_g = \text{density contrast } (\Delta\rho) \times \text{gravitational acceleration } (g) \times \text{volume}$$

$$V_{\text{Stokes}} = C \frac{\Delta\rho g a^2}{\eta_m}$$

A Stokes solution: Stokes sphere



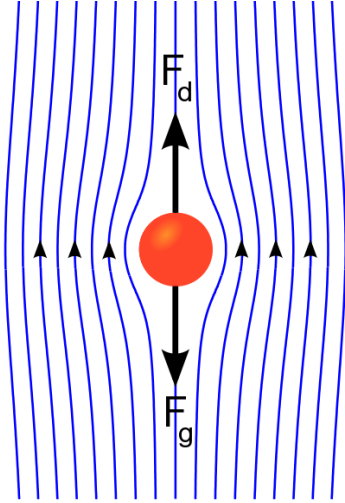
$$V_{\text{Stokes}} = C \frac{\Delta \rho g a^2}{\eta_m}$$

$$C = \frac{2 + 2\eta'}{6 + 9\eta'} \quad \eta' = \frac{\eta_s}{\eta_m}$$

Aside I:

→ A needle, as opposed to sphere, will sink with $0.5 \dots 2 v_{\text{Stokes}}$

Stokes sets the advective scale



$$V_{\text{Stokes}} = C \frac{\Delta \rho g a^2}{\eta_m}$$

Aside II:

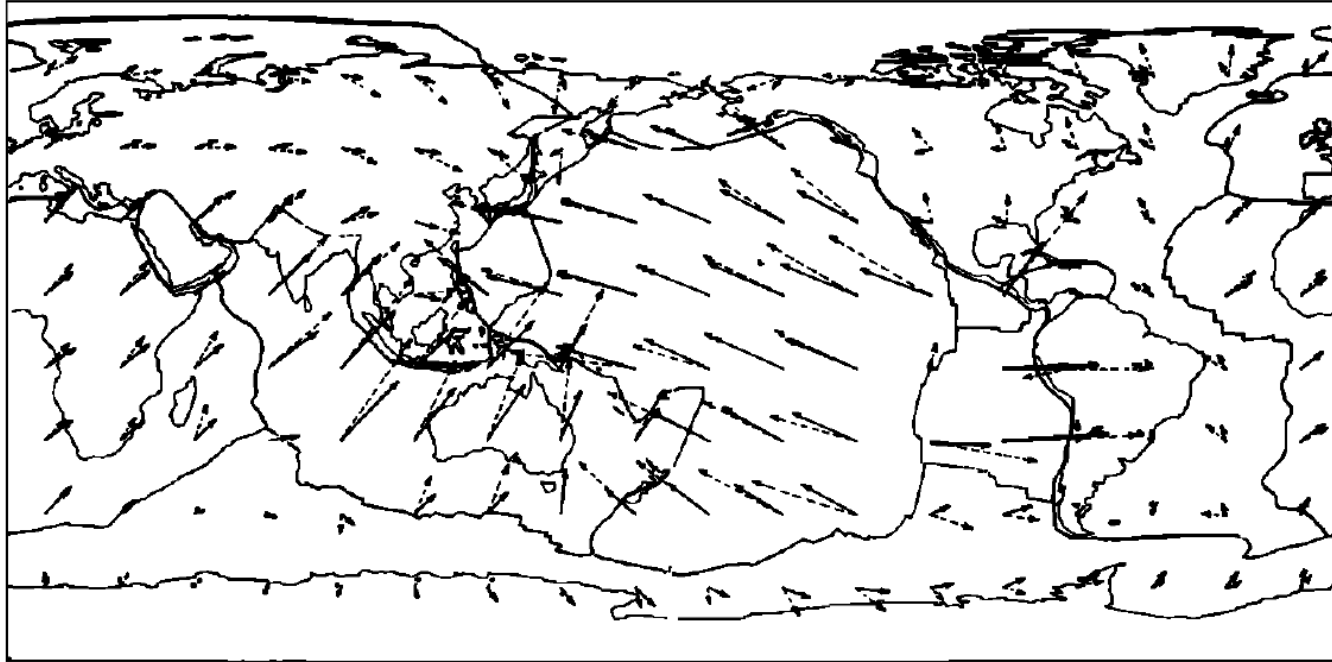
- Peclet# = ratio of diffusive to convective time scale, $Pe = t_{\text{diffusion}} / t_{\text{convection}}$
- $t_{\text{diff}} = a^2 / \kappa$, $t_{\text{conv}} = a / v_{\text{Stokes}}$, $\Delta \rho = \Delta T \alpha \rho_0$, then $Pe \rightarrow Ra$ (with a instead of L)

Density driven flow

$$v \propto \frac{\Delta\rho}{\eta}$$

$$\sigma \propto \eta \dot{\epsilon} \propto \eta \frac{v}{a} \propto \frac{\Delta\rho}{a}$$

Plate motions and global mantle flow



SHEAR STRESS AT BASE OF LITHOSPHERE

—→ Simple Drag

--→ Flow Model

Mantle circulation

- treat mantle and lithosphere as a fluid
- infinite Prandtl number (no inertia) approximation
 - Navier-Stokes turns into Stokes equation
- instantaneous solution for given density and boundary conditions
 - can solve in < 1 s for spherical Earth without lateral viscosity variations

force balance
(conservation of momentum)

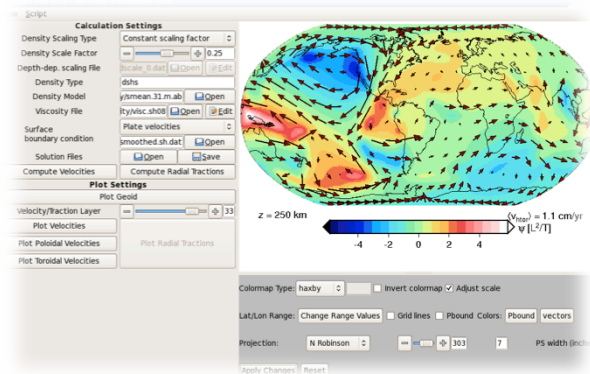
$$\frac{\partial \sigma_{ij}}{\partial x_j} = -Ra \tilde{T} \delta_{ir}$$

normalized temperature
thermal buoyancy

constitutive equation
(rheology)
and conservation of mass

$$\sigma_{ij} = -p\delta_{ij} + 2\eta\dot{\epsilon}_{ij}$$

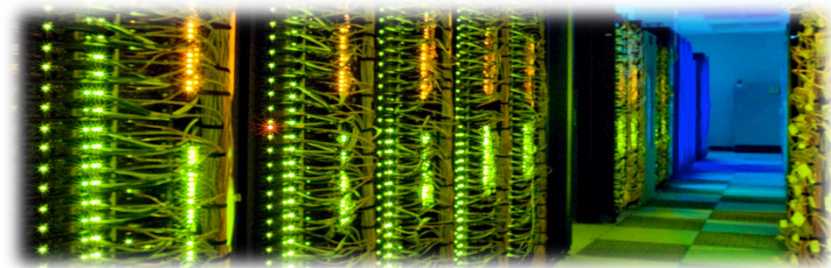
$$\frac{\partial v_i}{\partial x_i} = 0$$



HC of SEATREE GUI is part of UGESCE
or CIDER VirtualBox distribution

Mantle circulation

- Thermo-chemical heterogeneity and complex rheologies make things interesting
- Finite element methods best suited for lateral viscosity variations (can solve all of this, in ~hours, at < km resolution without approximations), with ~512 CPUs



force balance
(conservation of
momentum)

$$\frac{\partial \sigma_{ij}}{\partial x_j} = -Ra\tilde{T}\delta_{ir} + Ra_c\tilde{C}\delta_{ir}$$

chemical buoyancy

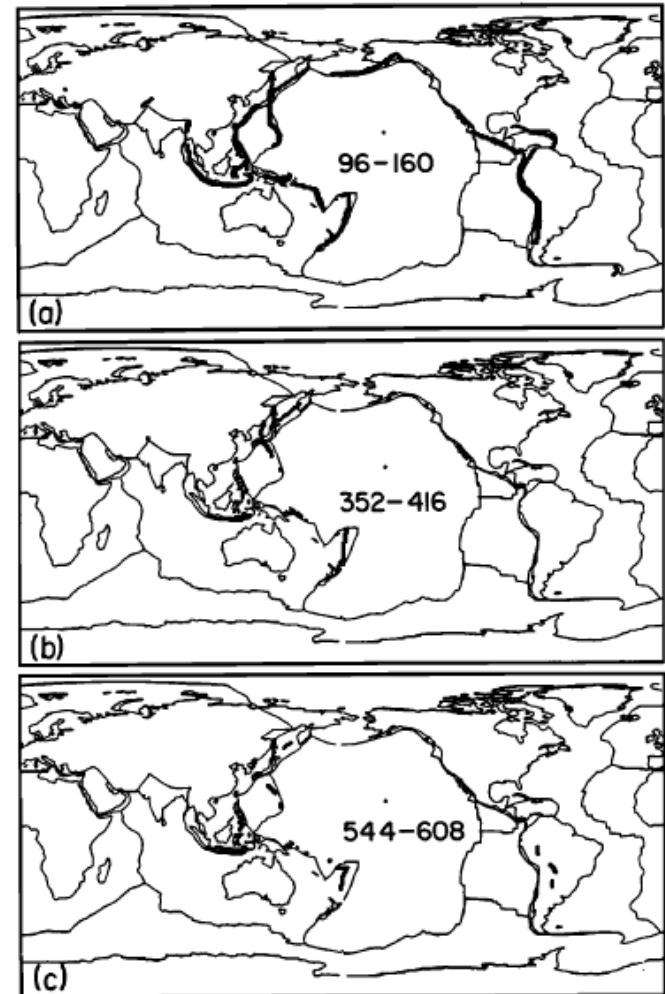
constitutive law
(rheology)

$$\sigma_{ij} = -p\delta_{ij} + 2\eta(\sigma, T, d, H_2O, \varepsilon)\dot{\varepsilon}_{ij}$$

non-linear rheology

How do we know the density variations in the mantle?

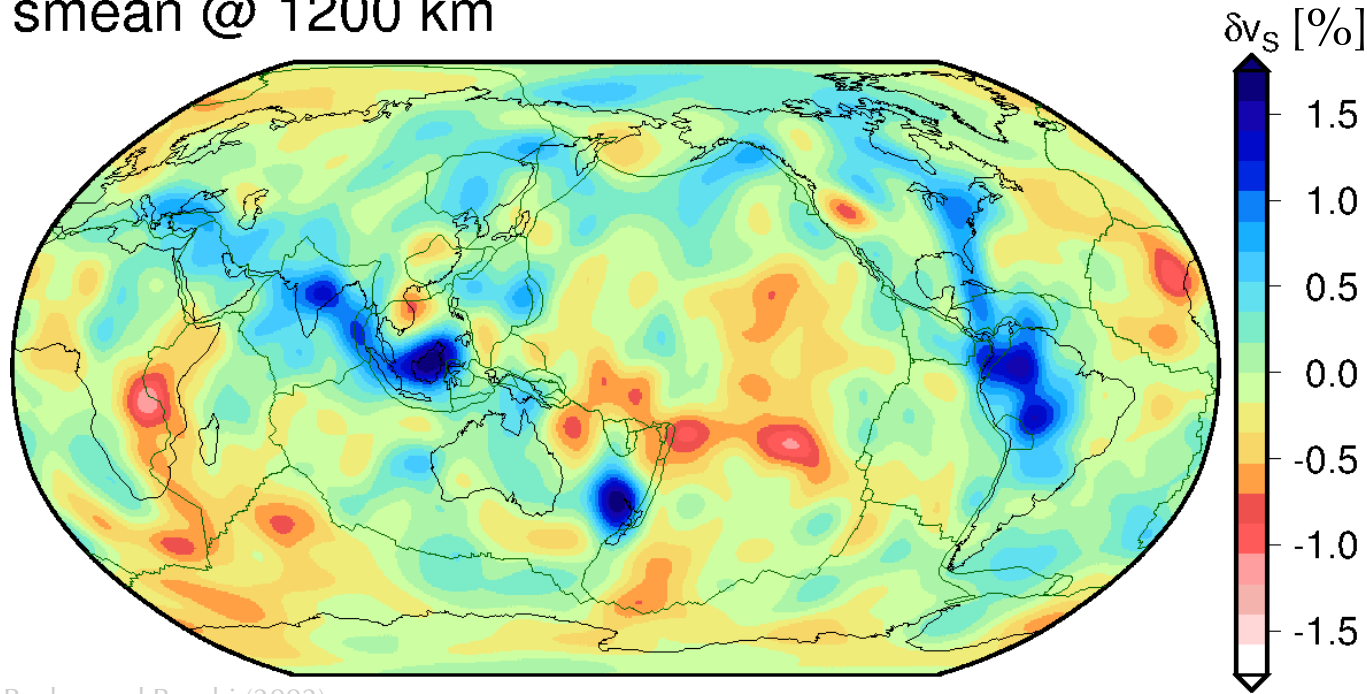
- upper mantle slabs from Wadati-Benioff zones



Hager (1984)

Seismic tomography also shows slabs

smean @ 1200 km

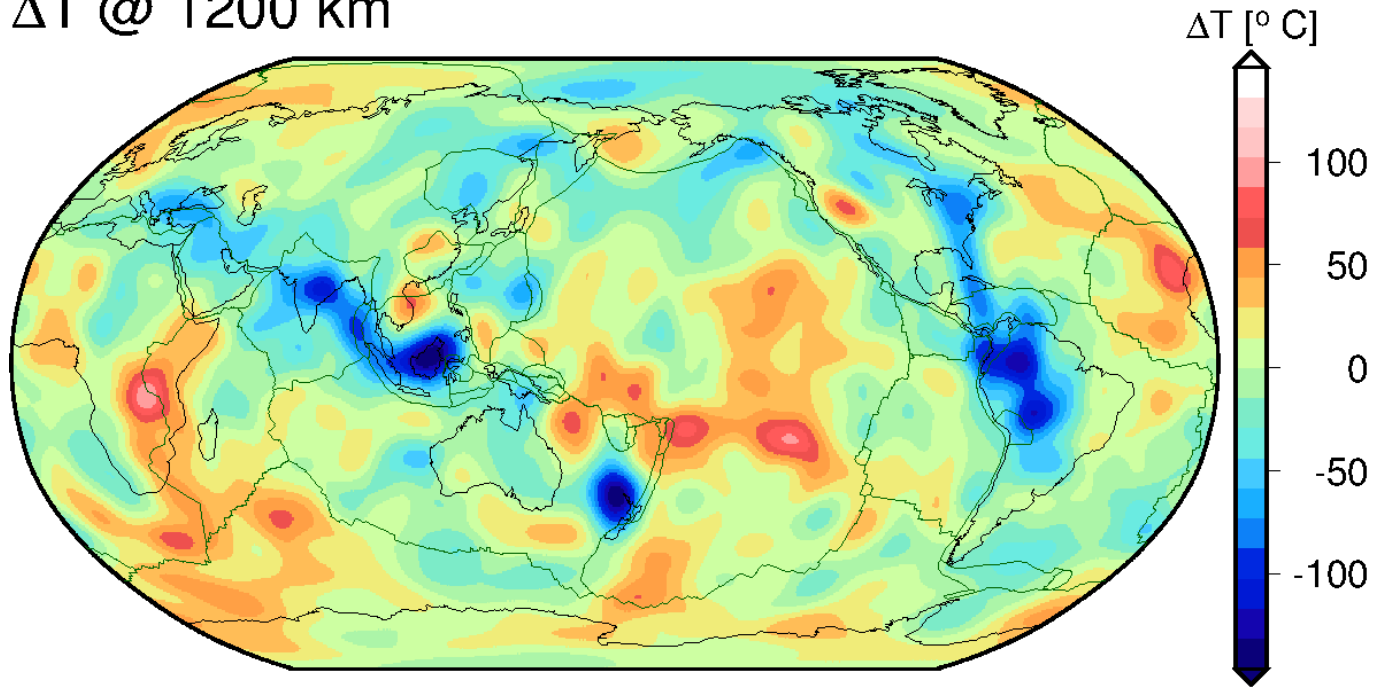


Becker and Boschi (2002)

- *S* wave models provide poor image of slabs in upper mantle
- *P* wave models poor outside subduction zones in upper mantle

Besides cratons and piles: $d \ln \rho / d \ln v_s = 0.2$

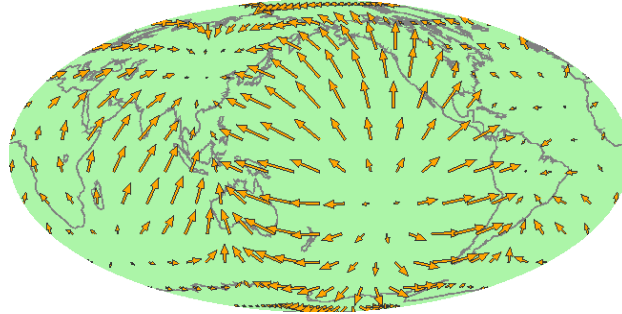
ΔT @ 1200 km



- Use mineral physics to convert velocity into temperature (density) anomalies

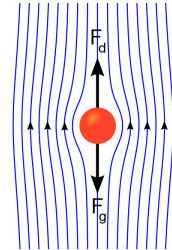
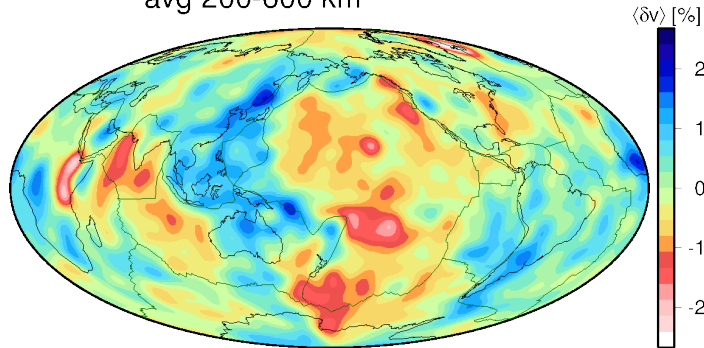
Tomography driven flow at surface

Flow model with free slip surface
(no lateral viscosity variations)



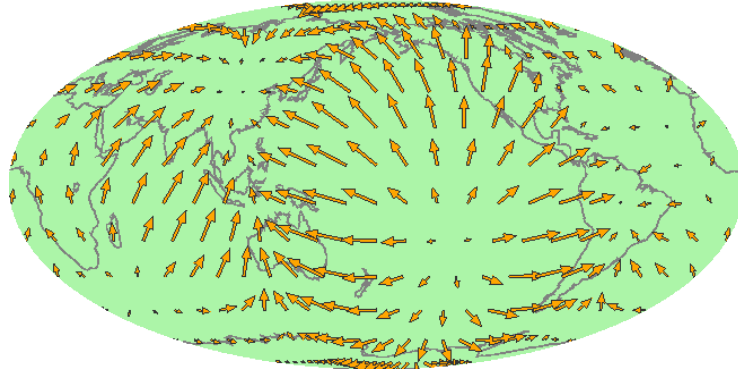
Upper mantle averaged tomography

avg 200-600 km



Surface motions compared to plate motions

Flow model with free slip surface and no LVVS



Observed plate velocities in hot spot reference frame

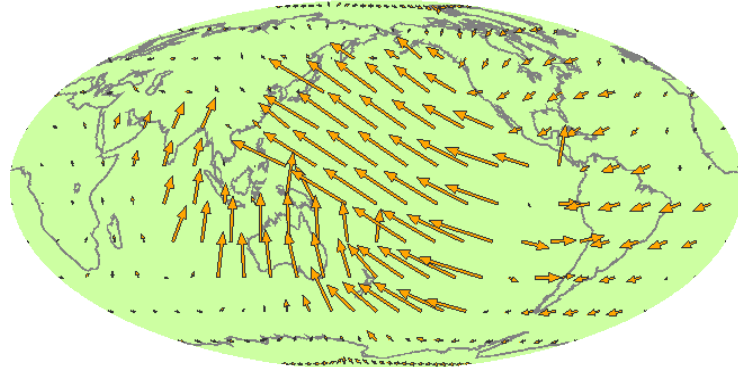
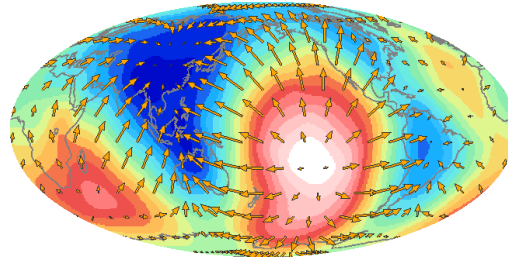
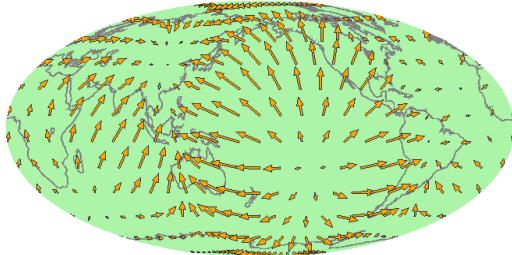


Plate motions require LVVs in lithosphere

Flow model with only radial viscosity variations

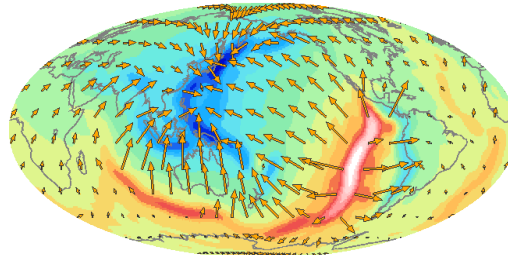
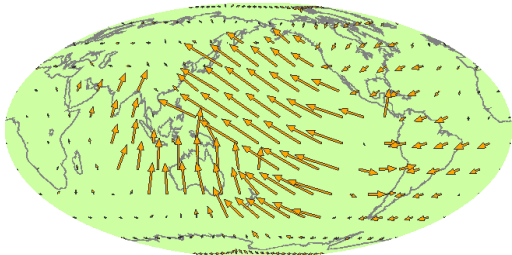
poloidal component



- No toroidal flow without lateral viscosity variations (no PT coupling)
- Strain-rates not very plate-like

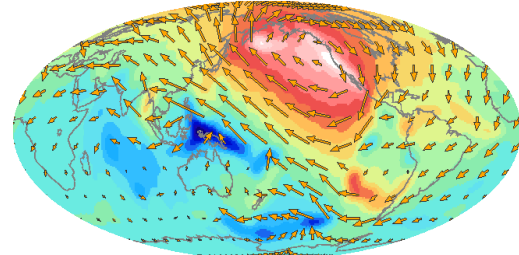
Observed plate velocities in hot spot reference frame

poloidal component



sources and sinks

toroidal component

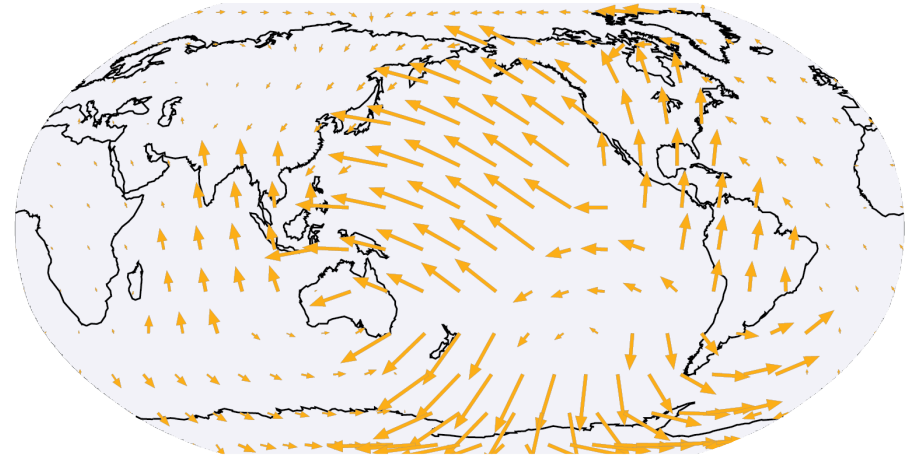


strike slip motion, spin

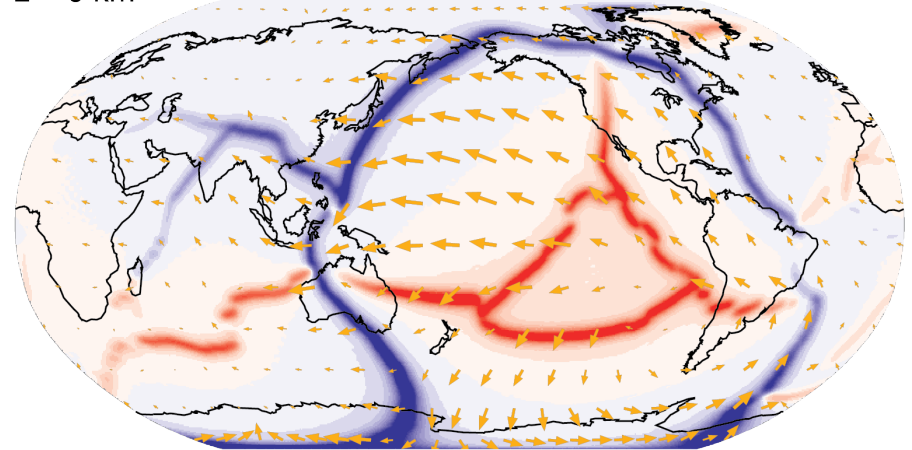
Plate-plate Interactions

- can compute interaction matrices for arbitrary plate geometries and viscosity distributions

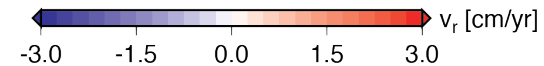
130 Ma



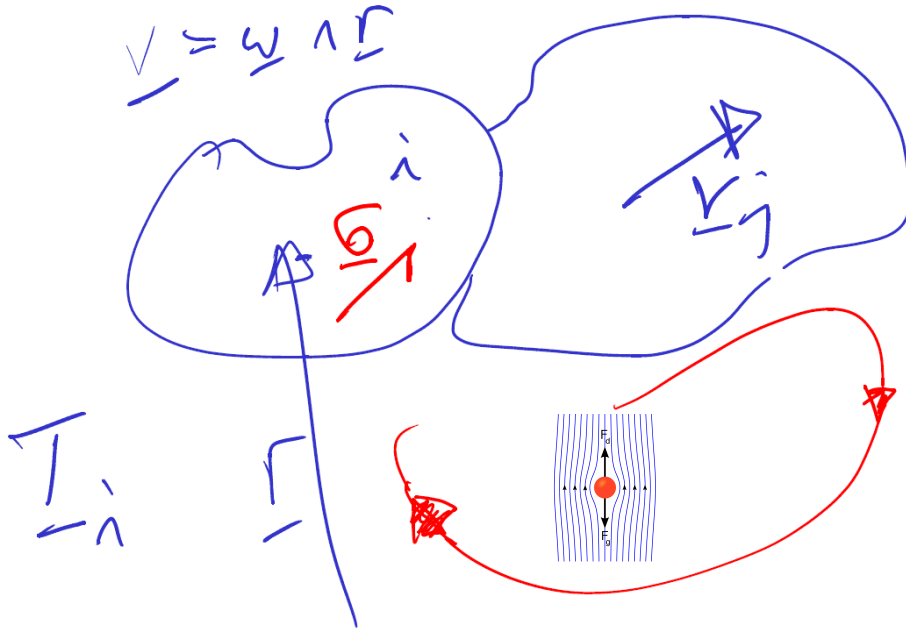
$z = 0$ km



$z = 248$ km



Solve for rigid plate motions given mantle tractions



$$\mathbf{T}_i = \int_{plate} dA \mathbf{r} \times \boldsymbol{\sigma}_{flow}$$

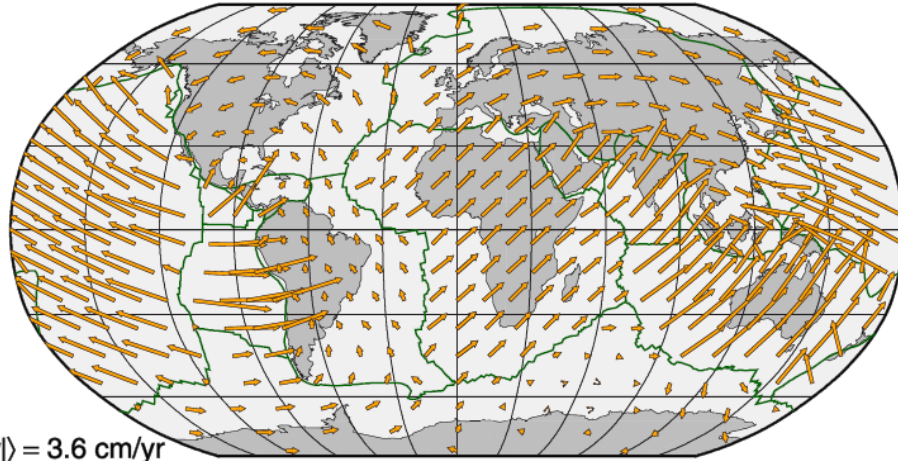
$$\mathbf{T}_{vd} = \mathbf{I}\boldsymbol{\omega} = \sum^M T_i$$

Note: Can construct plate motion interaction matrix \mathbf{I} , for any LVVs

Plate motion models ~work

plate motions

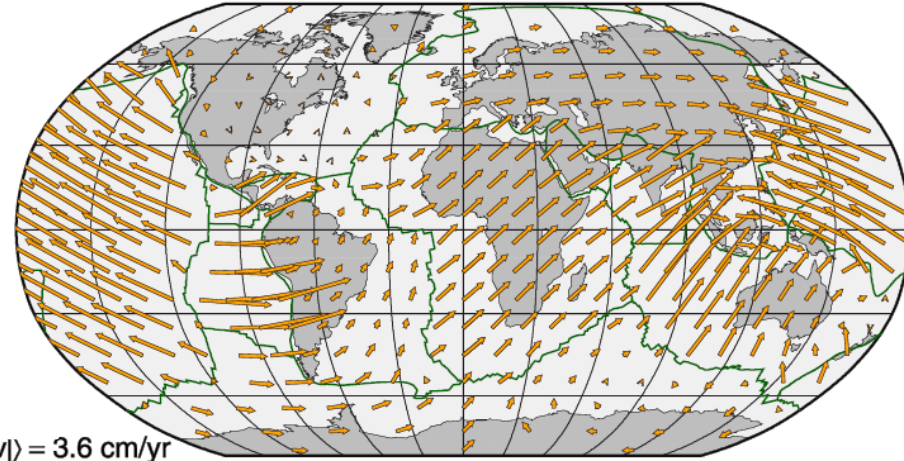
NUVEL1-A



$\langle |v| \rangle = 3.6 \text{ cm/yr}$

mantle flow model prediction

$r=0.94$

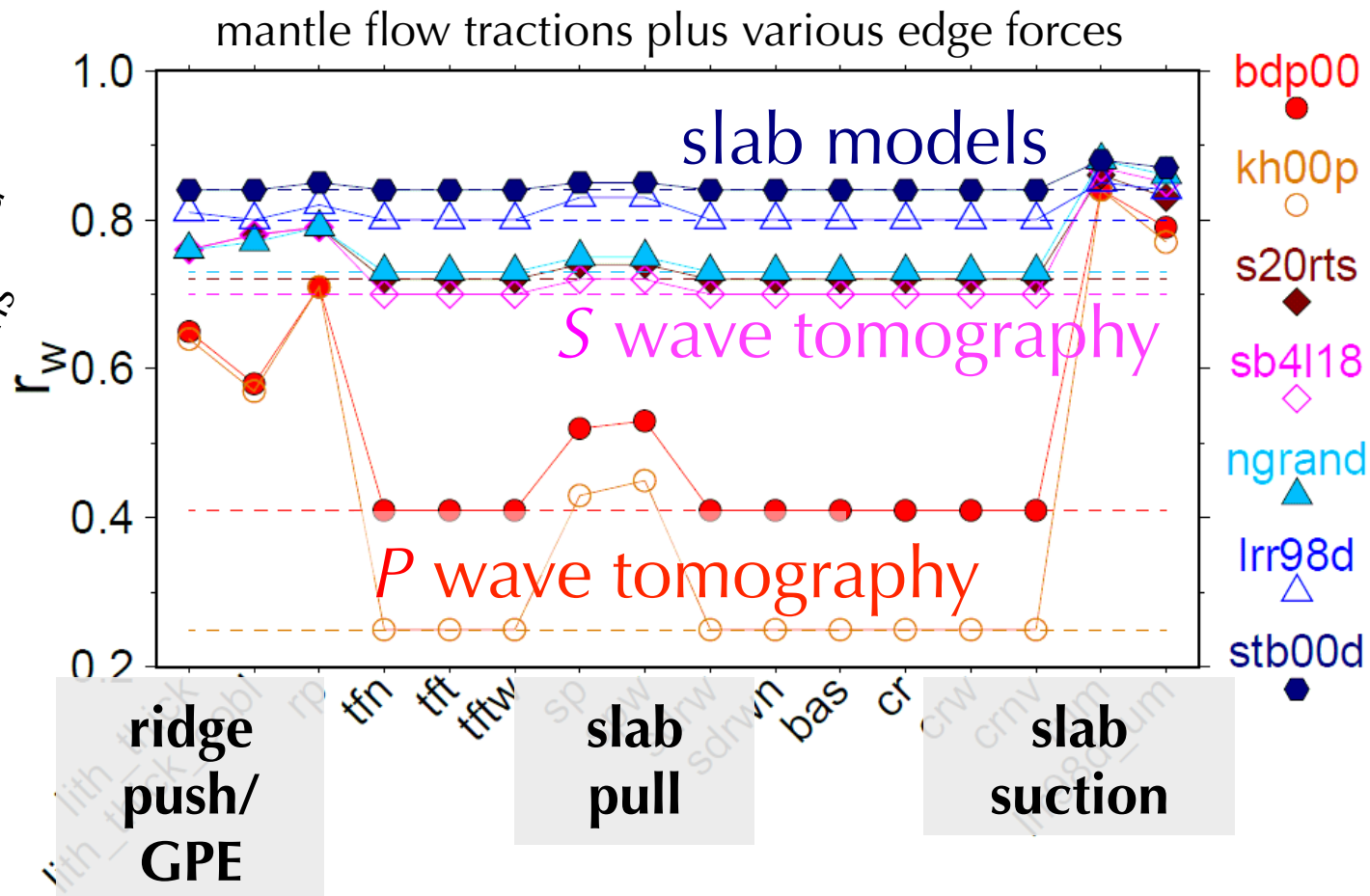


$\langle |v| \rangle = 3.6 \text{ cm/yr}$

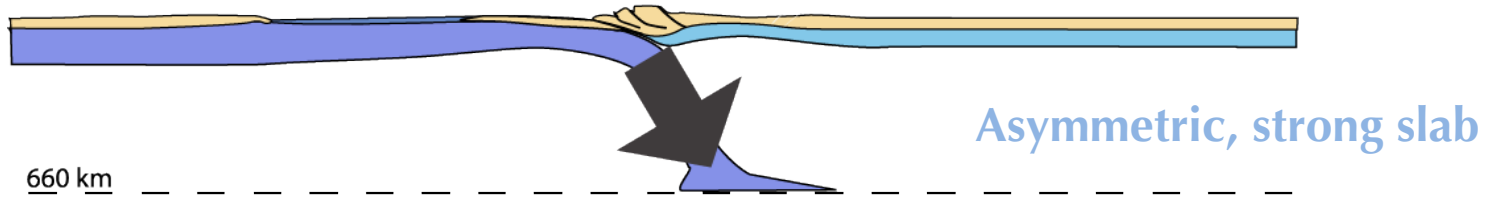
Faccenna et al. (2013)

**Slabs 70% vs.
rest 30%**

Correlation between model and
observed plate motions



Slab Pull



Slab Suction

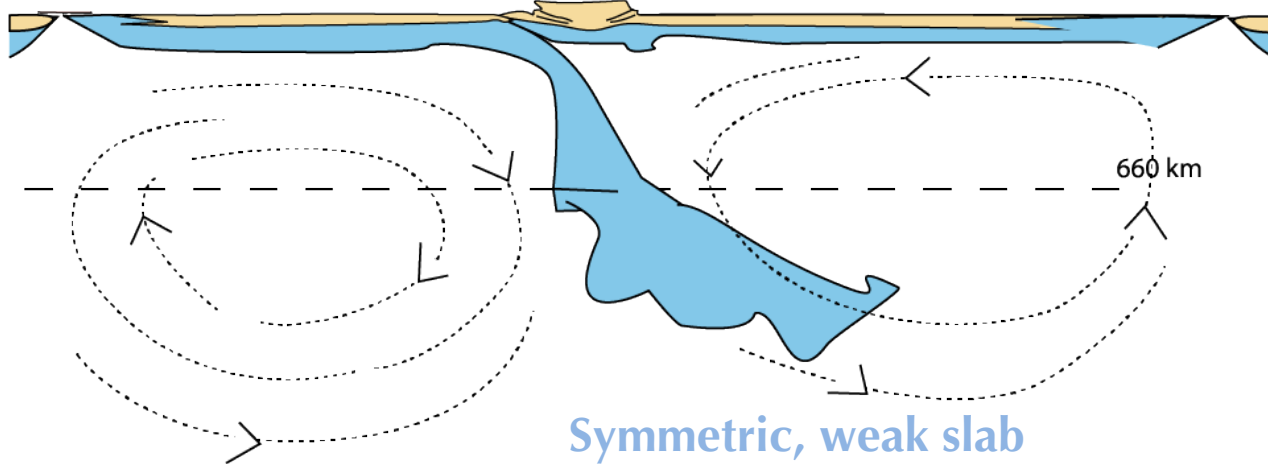
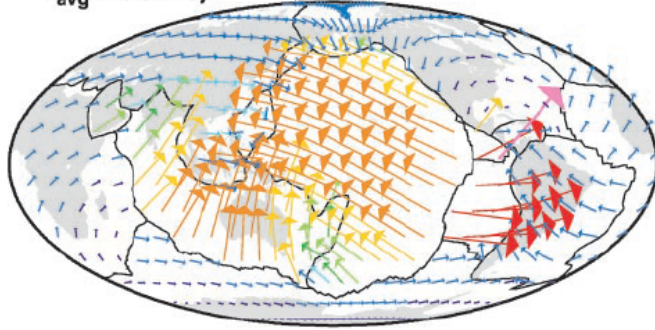


Plate speed vs. slab suction / slab pull

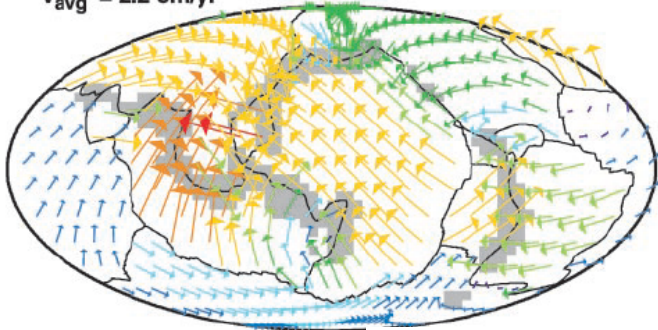
Observed Plate Velocities

$V_{avg} = 5.5 \text{ cm/yr}$



Slab Suction Only

$V_{avg} = 2.2 \text{ cm/yr}$



Combined Slab Pull and Slab Suction

$V_{avg} = 4.5 \text{ cm/yr}$

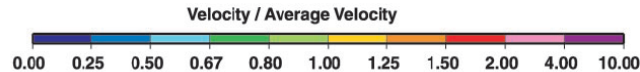
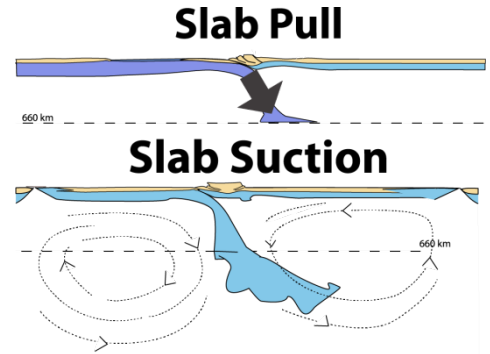
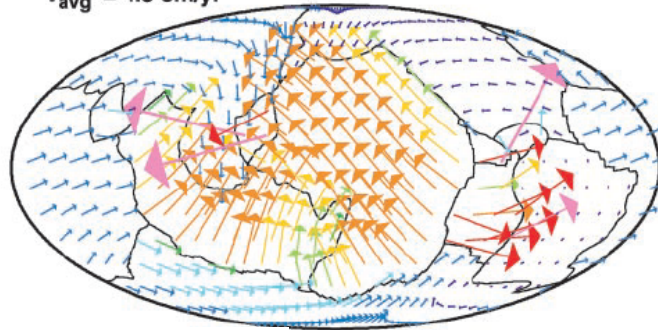
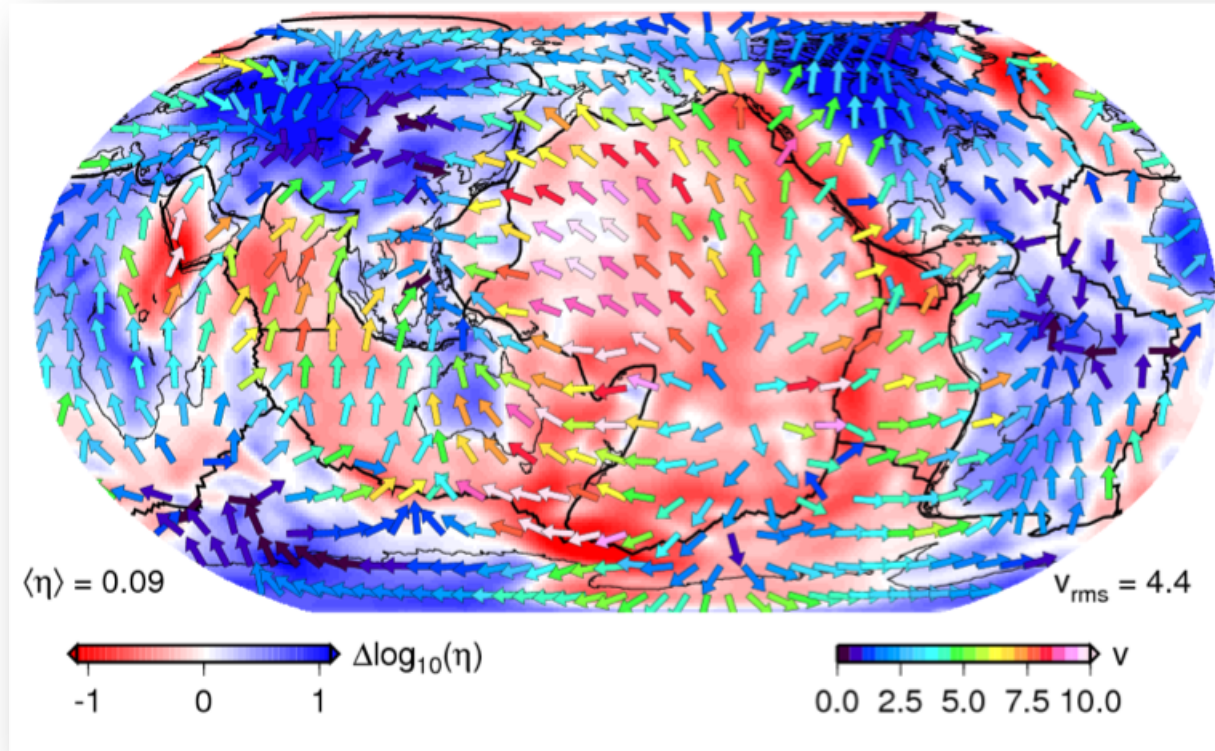
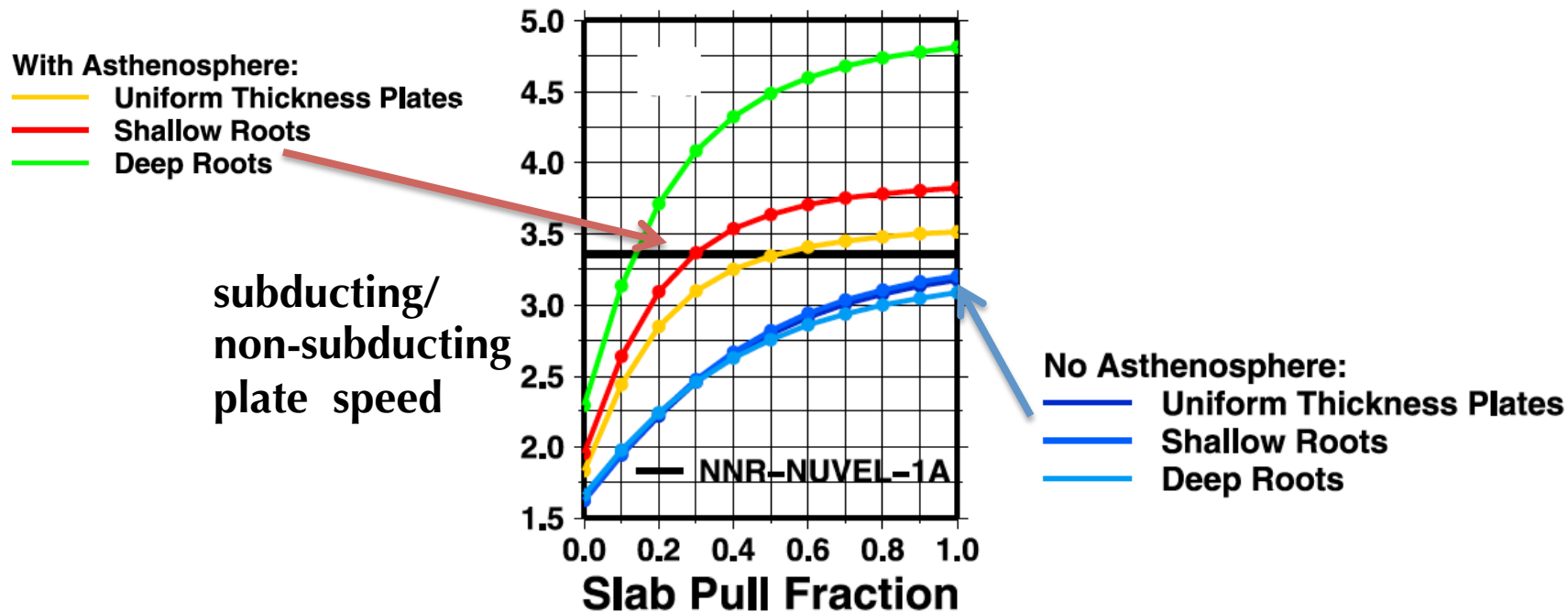


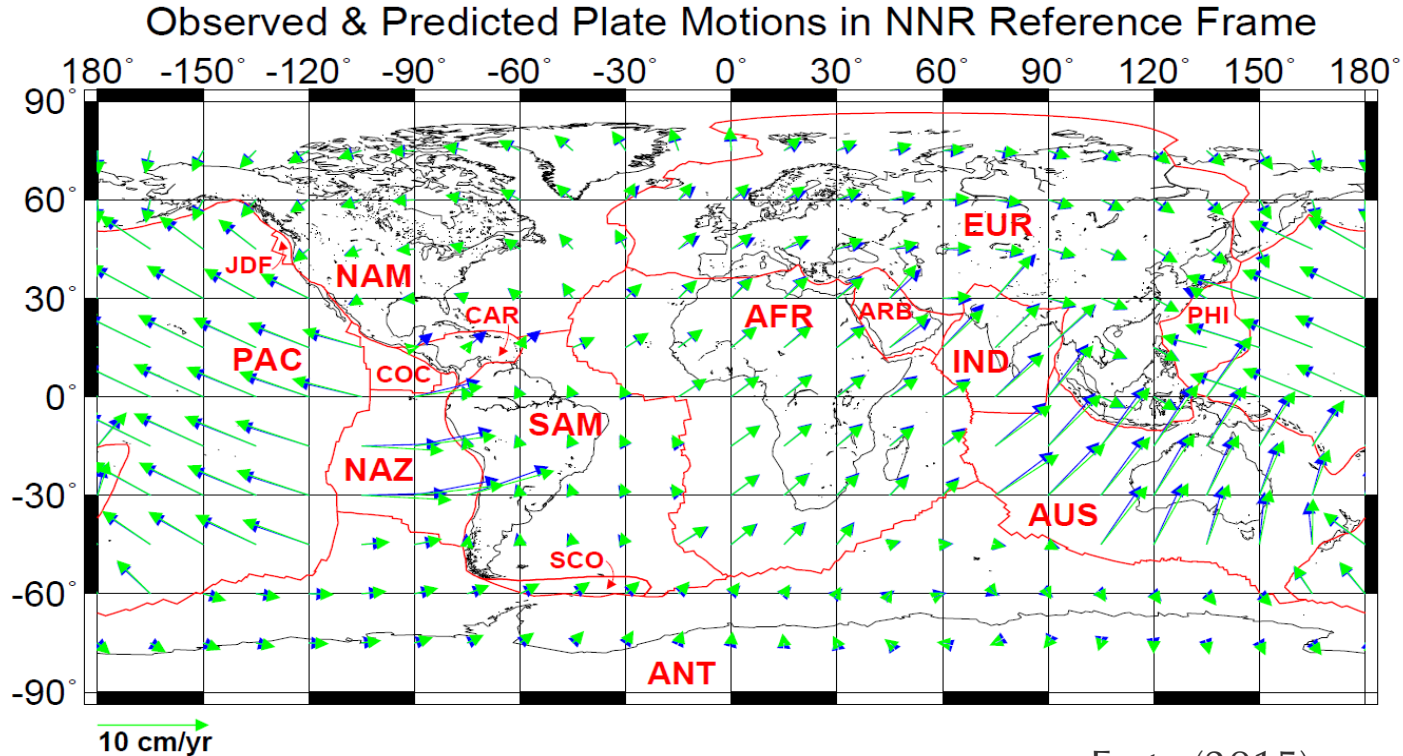
Plate speed vs. asthenospheric viscosity



Asthenosphere vs. slab viscosity tradeoff

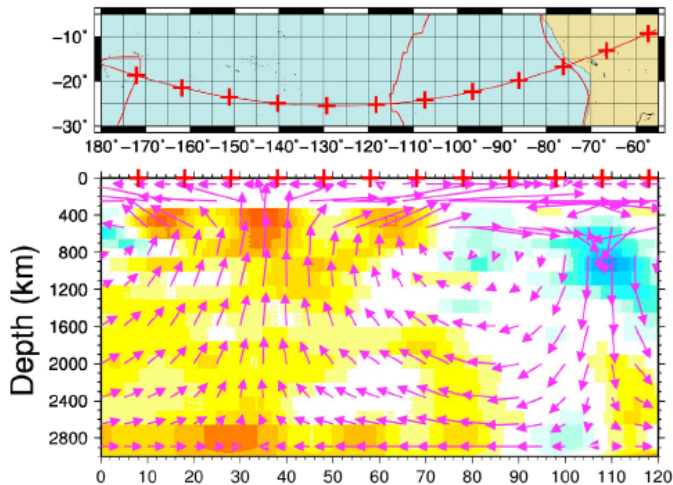


Some more uncertainties

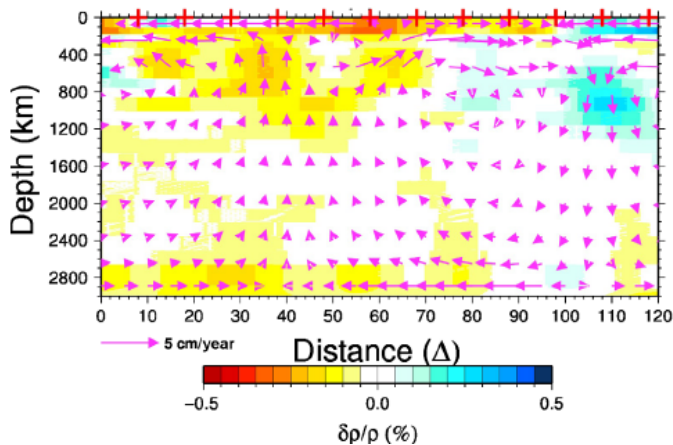


Forte (2015)

A S20RTS dens. & L4 visc.



C S20RTS dens. (from inversion) & V2 visc.

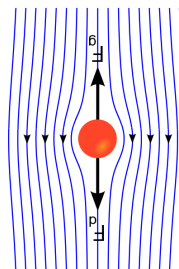


Importance of viscosity and density

(mineral physics,
compositional
anomalies)

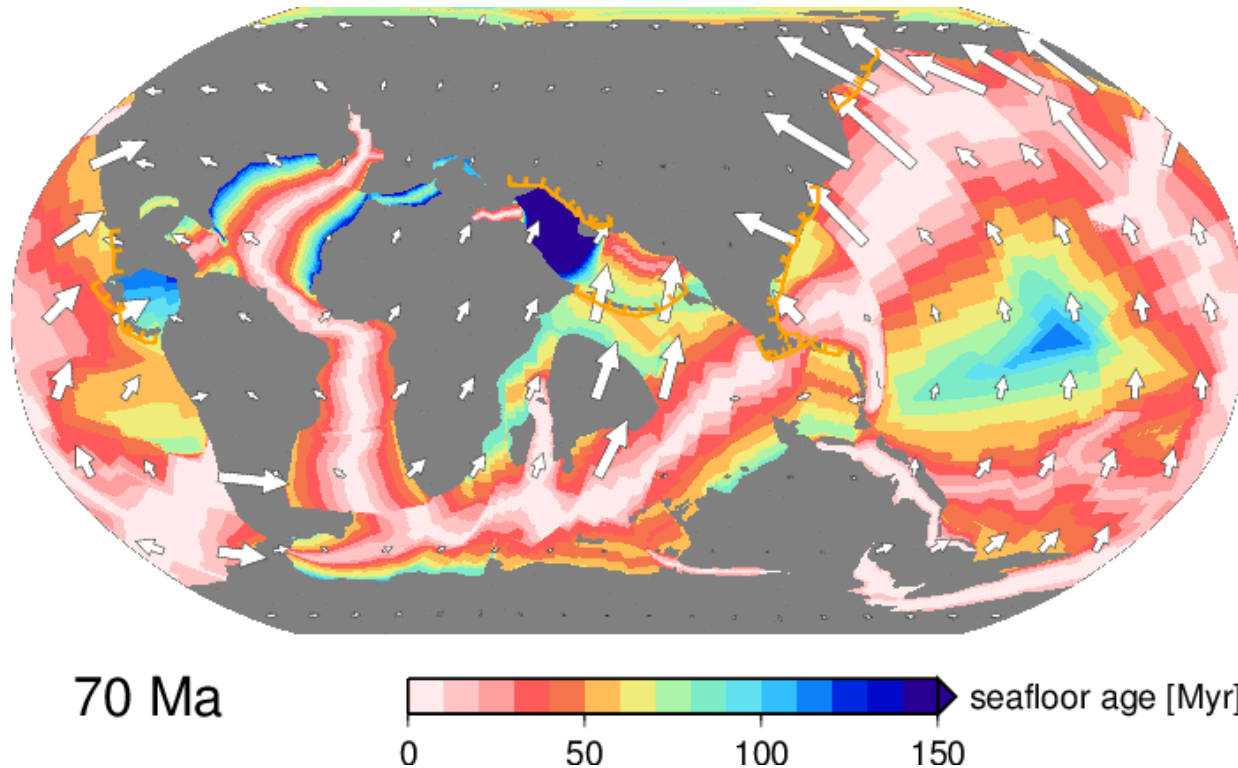
simple tomography to
density scaling, simple
viscosity structure

joint tomography /
dynamics inversion
for density, complicated
viscosity structure

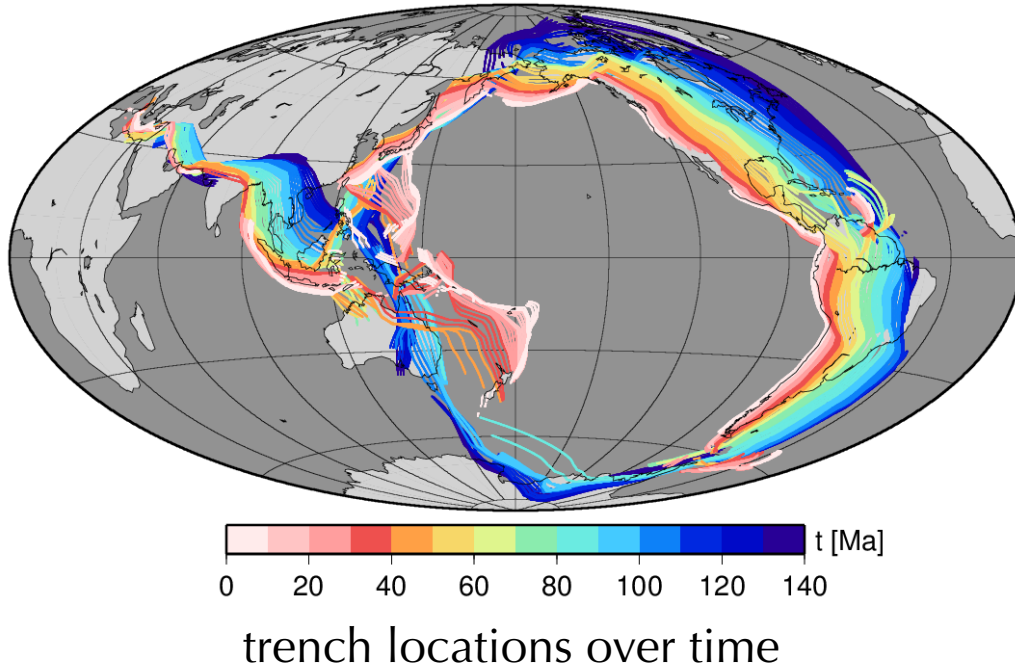


Rowley et al. (2016)
cf. Simmons et al. (2009)
Forte and Mitrovica (2004)
Forte (2015)

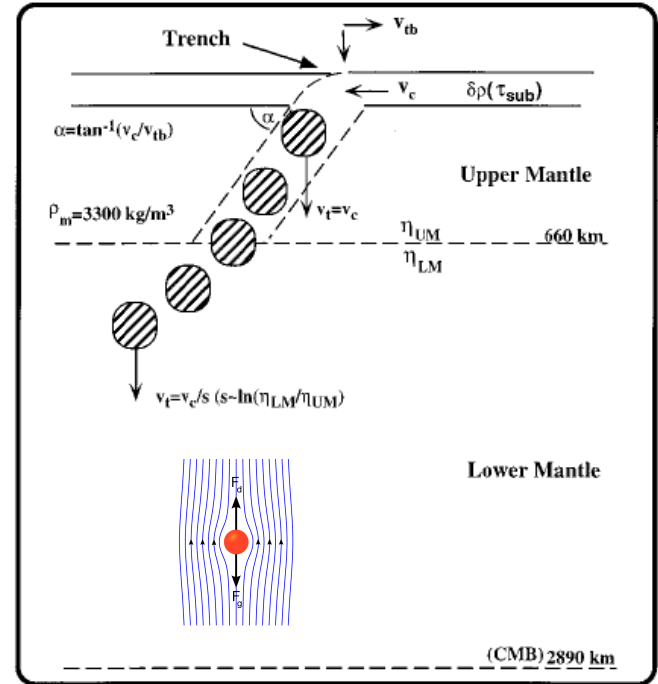
Global plate tectonics experiments in the past



Stokeslets link subduction history and mantle viscosity to present-day structure



Steinberger et al. (2010)

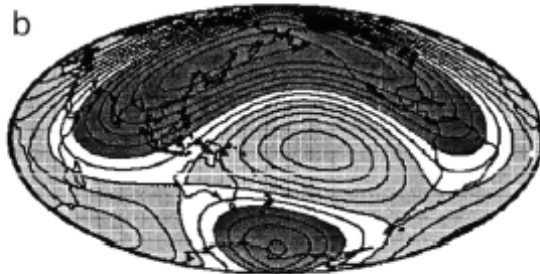


Ricard et al. (1993); Lithgow-Bertelloni et al. (1993, 1998)

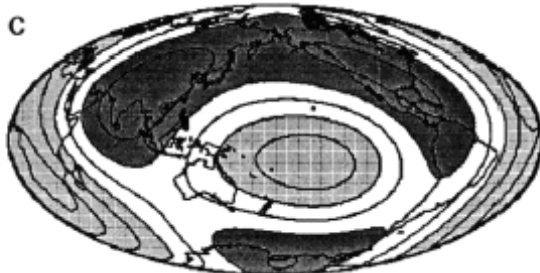
SLABS (DEPTH 2000 KM, DEGREES 1-15)



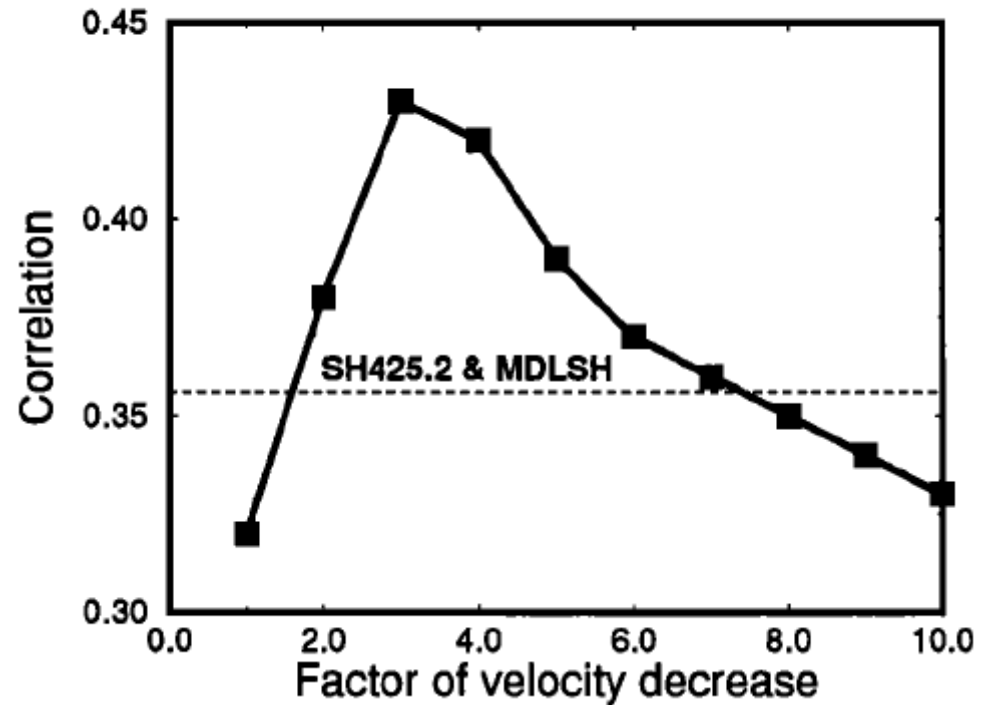
SLABS (DEPTH 2000 KM, DEGREES 1-3)



SH425.2 (DEPTH 2000 KM, DEGREES 1-3)

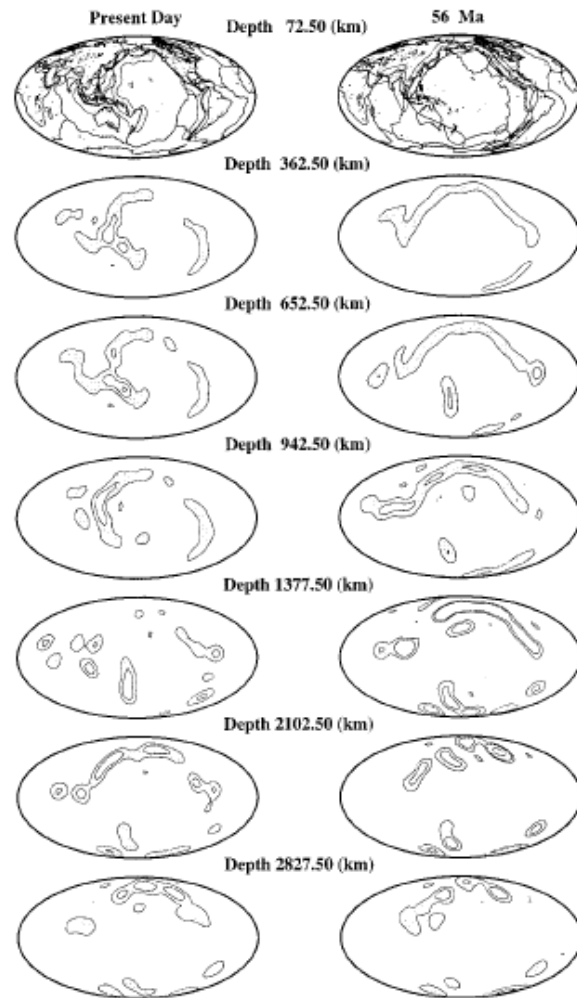
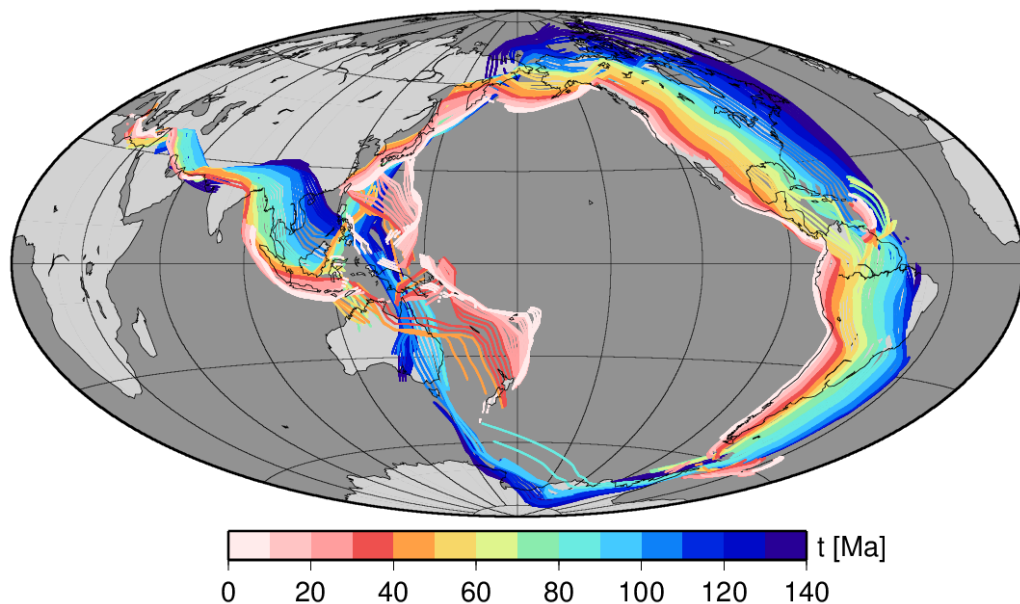


CORRELATION SLABS/SH425.2



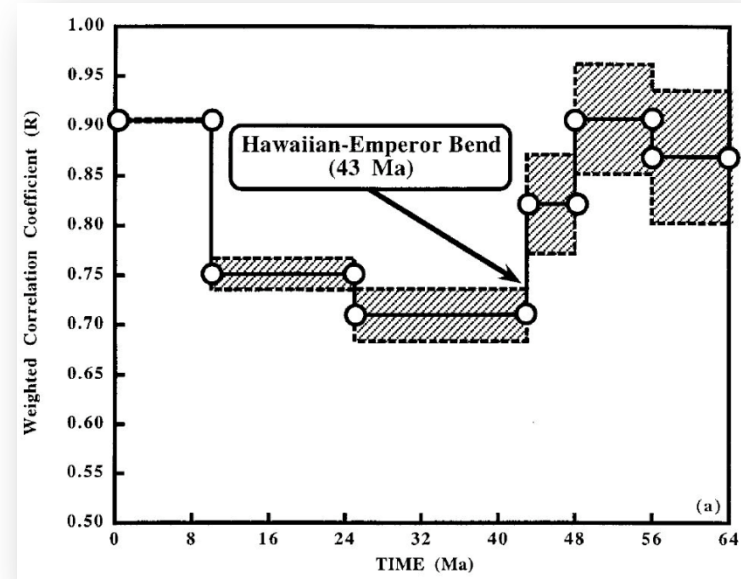
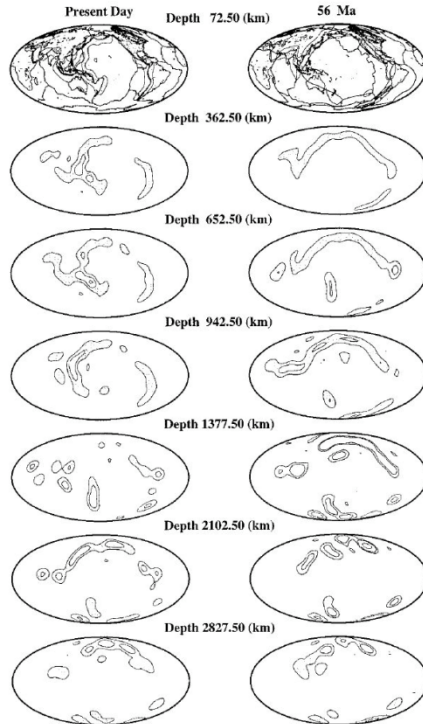
➤ constraint on upper/lower mantle viscosity increase

Mantle structure as $f(t)$ from subduction



Ricard et al. (1993); Lithgow-Bertelloni & Richards (1998);
Steinberger (2000); Spasojeovich et al. (2009); Steinberger and Torsvik (2010)
van der Meer (2010); Steinberger et al. (2014); Bower et al. (2015)

Time-dependent match to plate motions based on slablets

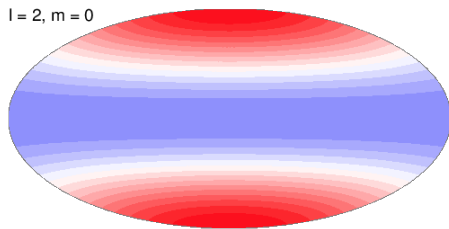


Lithgow-Bertelloni and Richards (1998)

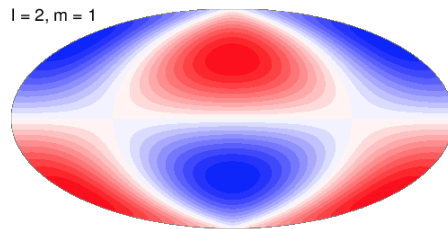
Spherical harmonics for analysis of global fields

increasing
degree /
(1/wavelength)

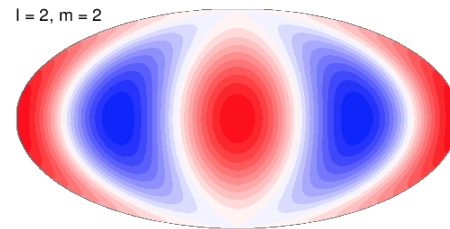
$l = 2, m = 0$



$l = 2, m = 1$

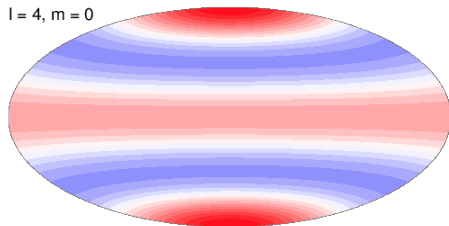


$l = 2, m = 2$

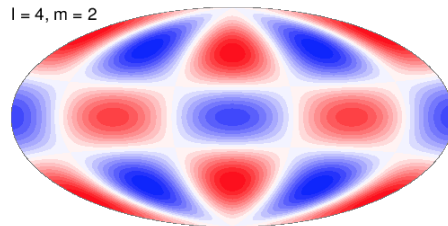


$$\delta v(\theta, \phi) \approx \sum_{l=0}^{\ell_{\max}} \left[a_{\ell 0} X_{\ell 0}(\theta) + \sqrt{2} \sum_{m=1}^{\ell} X_{\ell m}(\theta) \times (a_{\ell m} \cos m\phi + b_{\ell m} \sin m\phi) \right]$$

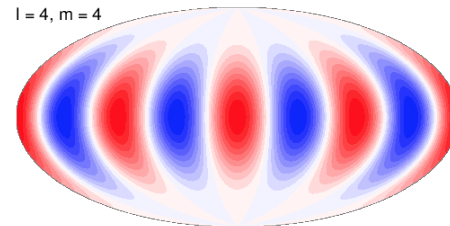
$l = 4, m = 0$



$l = 4, m = 2$

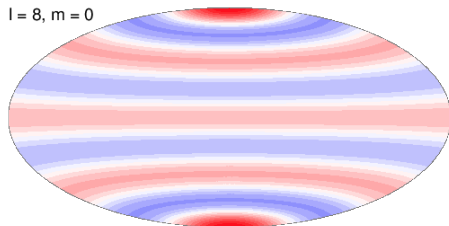


$l = 4, m = 4$

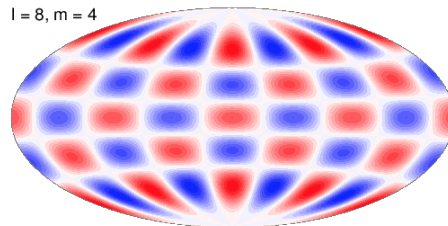


increasing order m ($0 < m < l$)

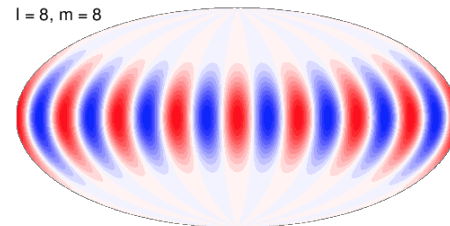
$l = 8, m = 0$



$l = 8, m = 4$

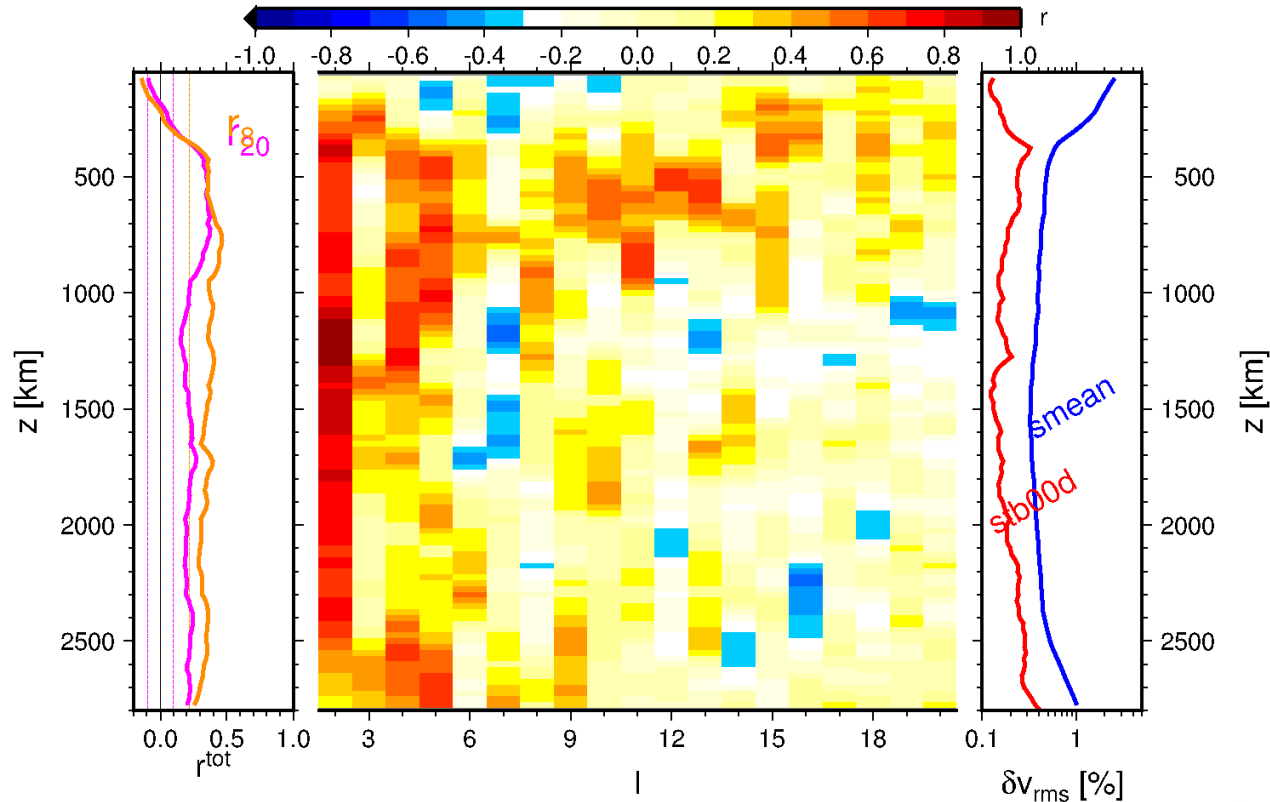


$l = 8, m = 8$

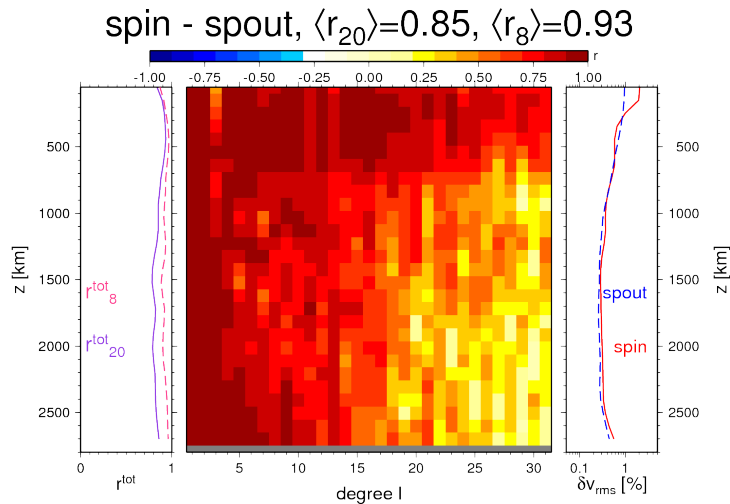


Correlation of global slab model with tomography (does not work that well...)

stb00d vs. smean, $\langle r_{20} \rangle = 0.21$, $\langle r_8 \rangle = 0.30$



Seeing the slab I



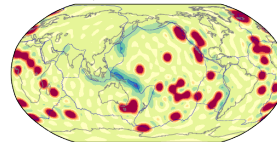
Note: can resolve structure up to degree $l_{max} \sim 15$
in wave theoretical framework

depth
↓

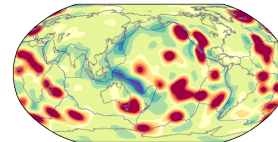
Input

Output

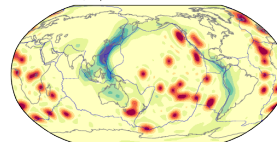
z = 500 km, input



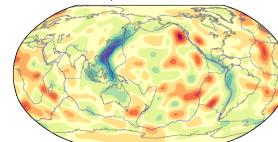
z = 500 km, output



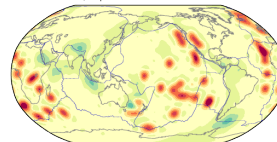
z = 1000 km, input



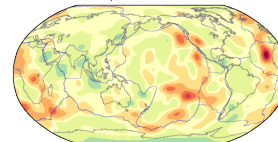
z = 1000 km, output



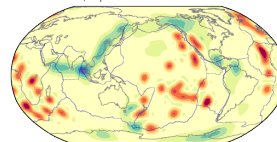
z = 1500 km, input



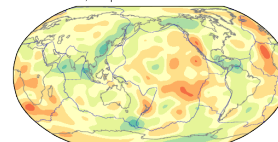
z = 1500 km, output



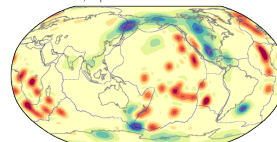
z = 2000 km, input



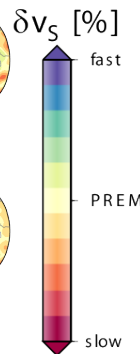
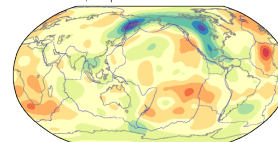
z = 2000 km, output



z = 2500 km, input

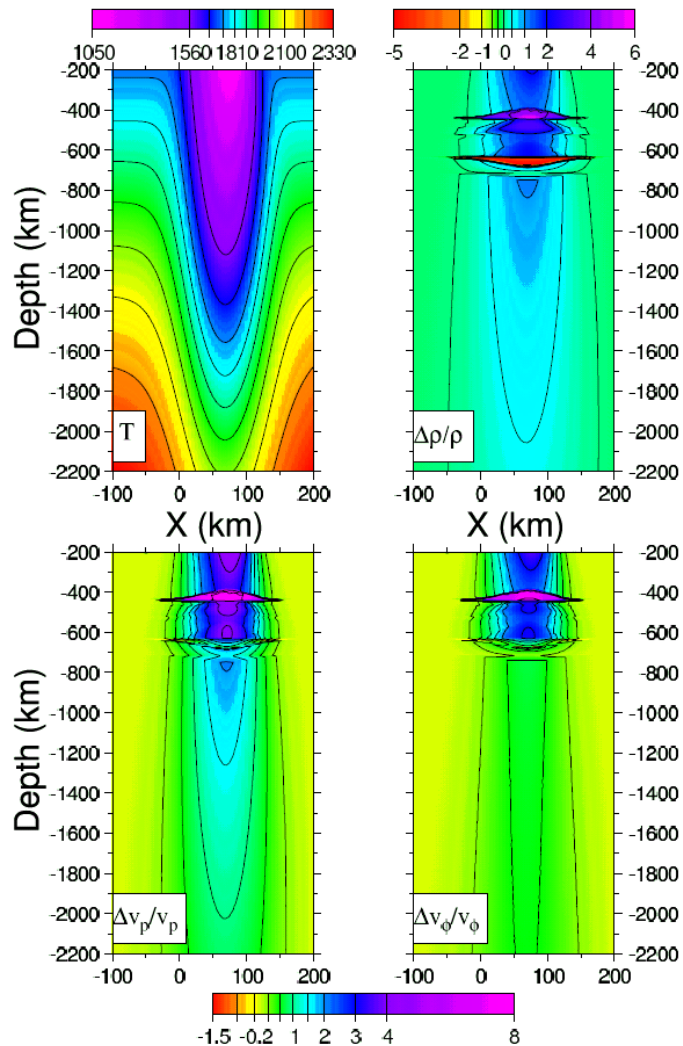


z = 2500 km, output



Seeing the slab II

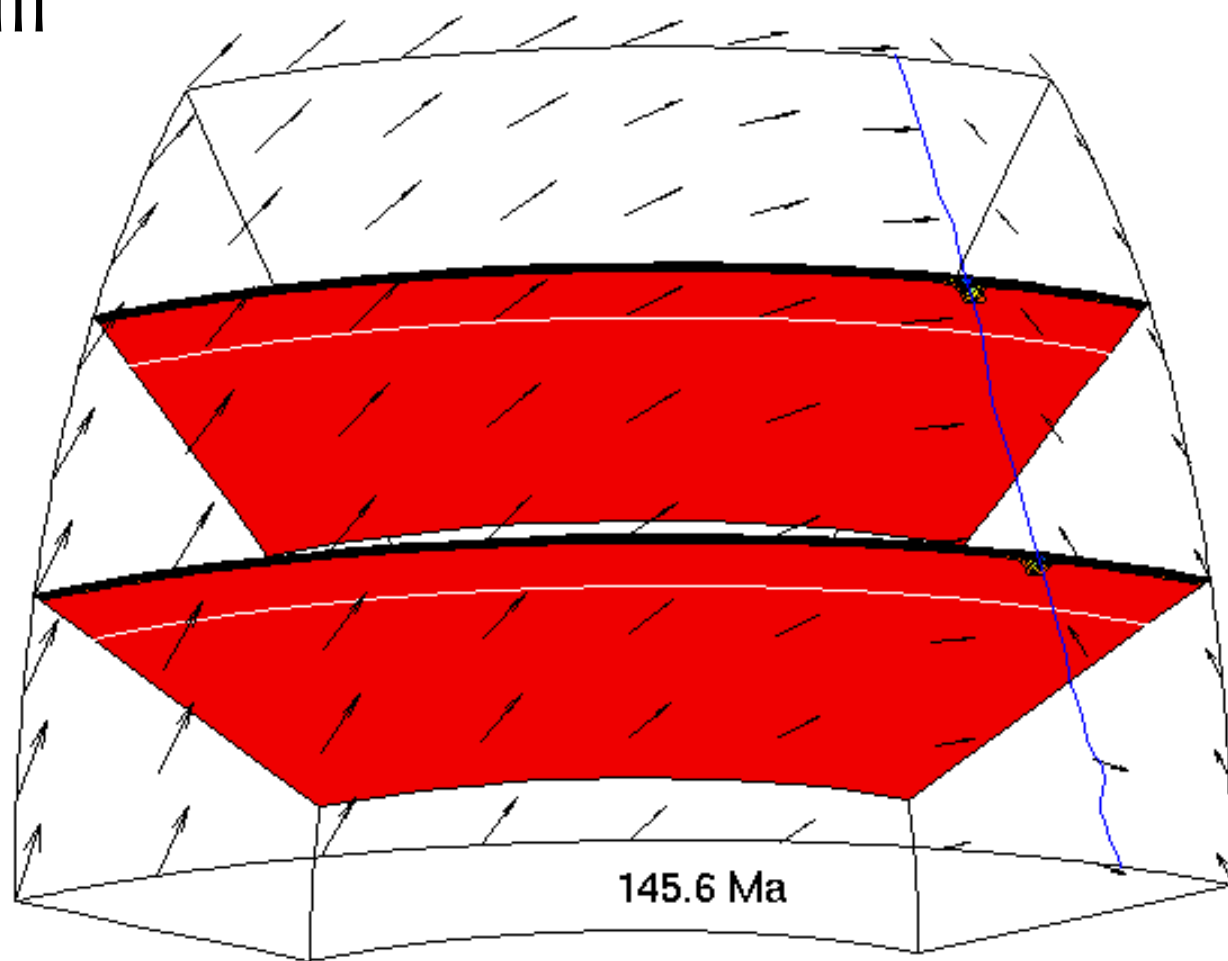
?



Ricard et al. (2005)

Seeing the slab III

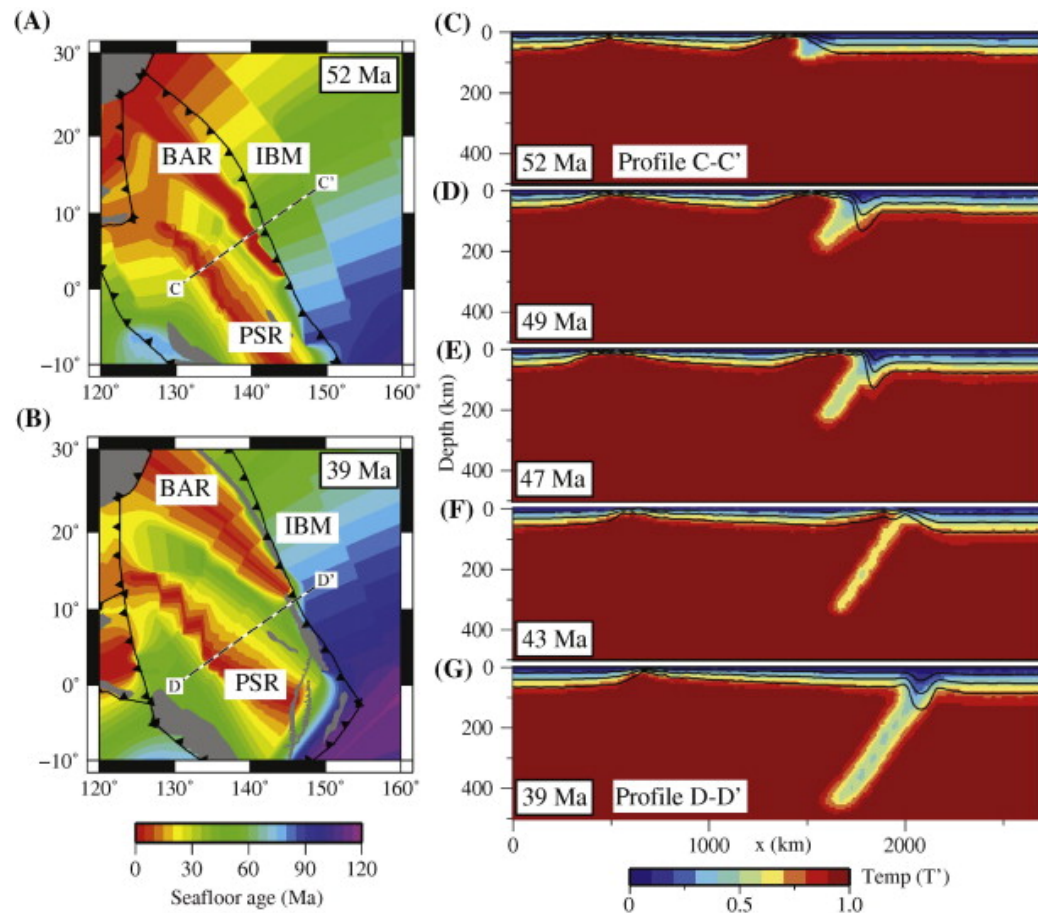
?



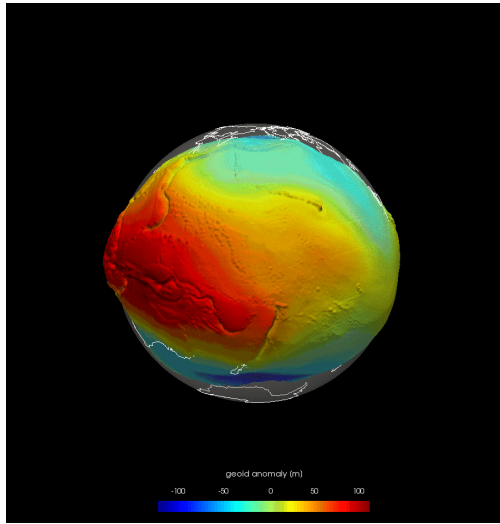
Slab models do not quite explain mantle structure, because of...

uncertainties about

- plate reconstructions
- slab and mantle rheology
- mineral physics and composition
 - mass flux through 660
- active upwellings and thermo-chemical piles



Other constraints on slabs



vertically exaggerated EGM360 geoid

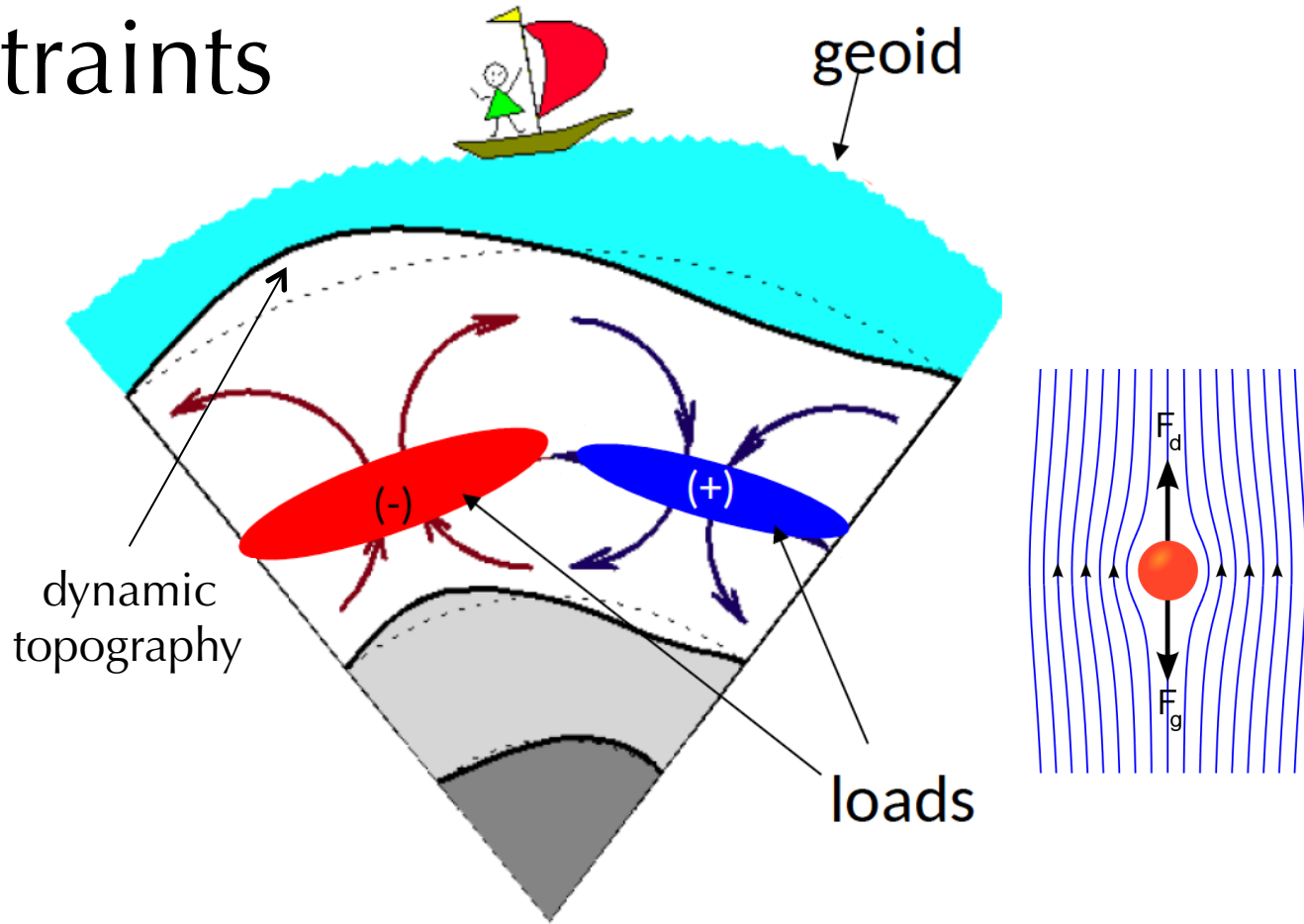
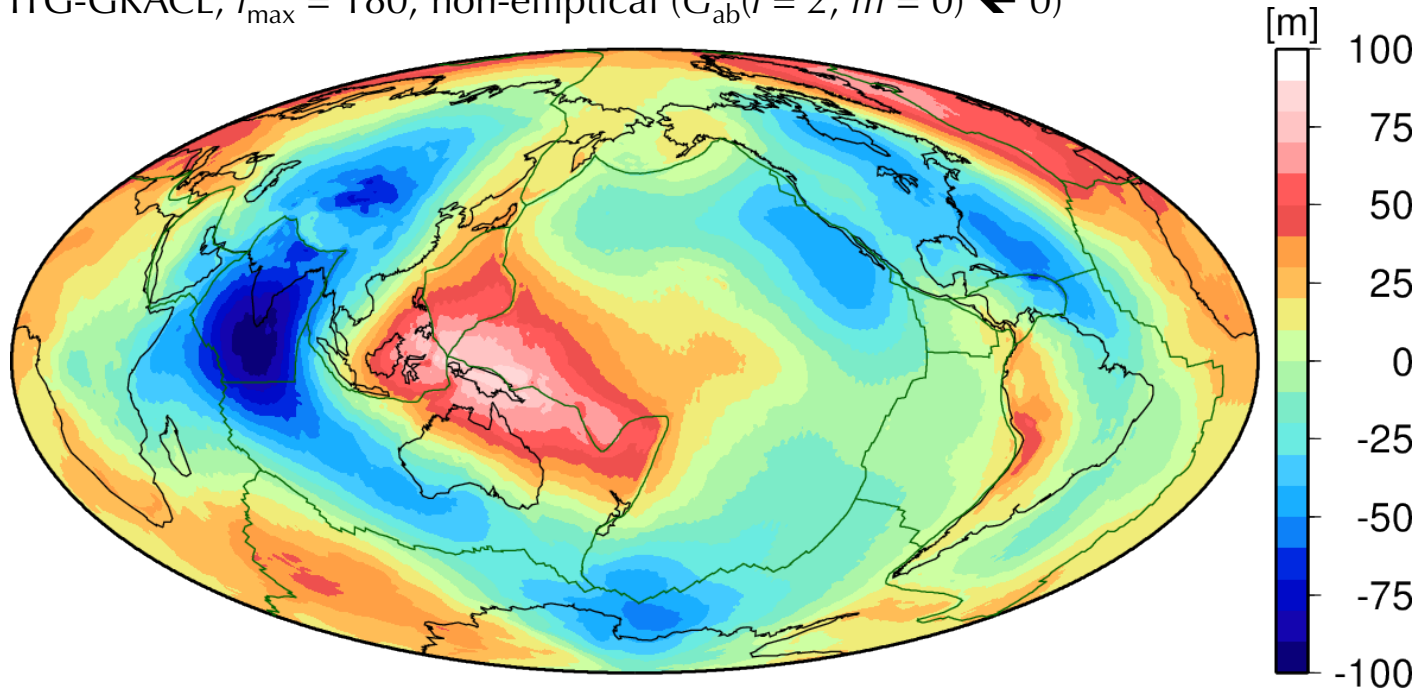


Figure from Y. Ricard

Geoid anomalies

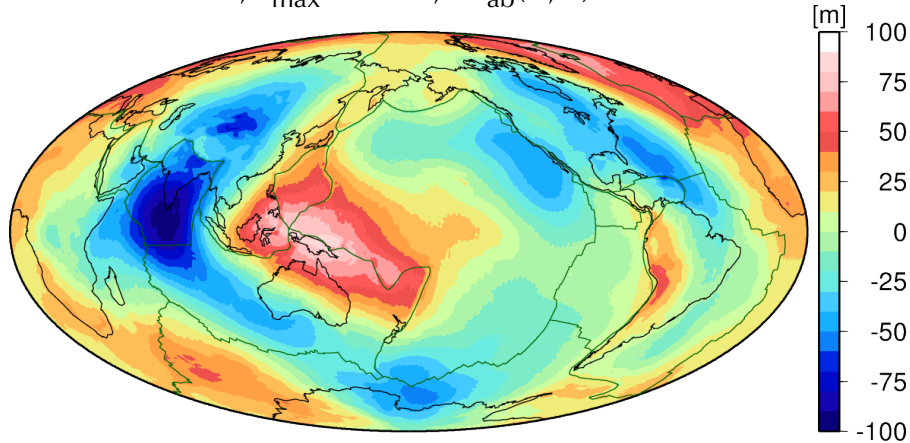
ITG-GRACE, $l_{\max} = 180$; non-elliptical ($G_{ab}(l = 2, m = 0) \leftarrow 0$)



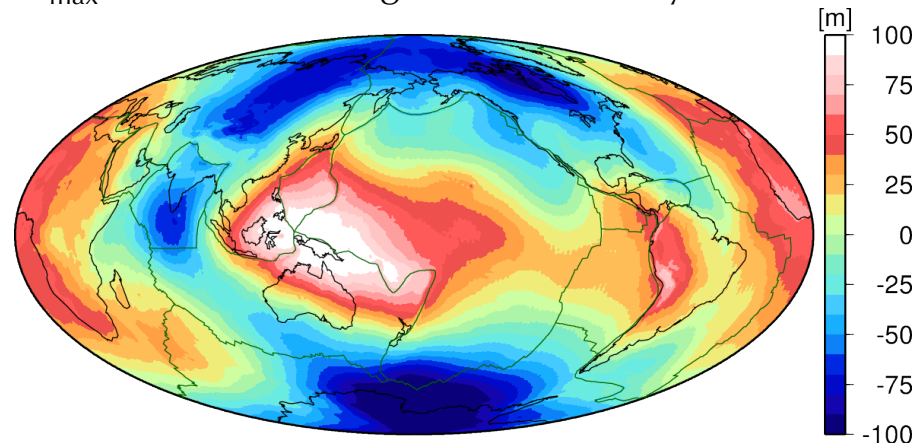
scale is saturated

Geoid anomalies

ITG-GRACE, $l_{\max} = 180$; $C_{ab}(2,0) = 0$



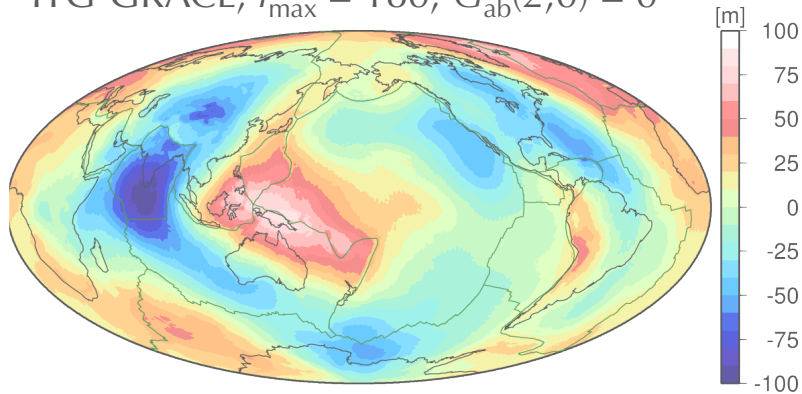
$l_{\max} = 180$ – Nakiboglu (1982) non-hydrostatic



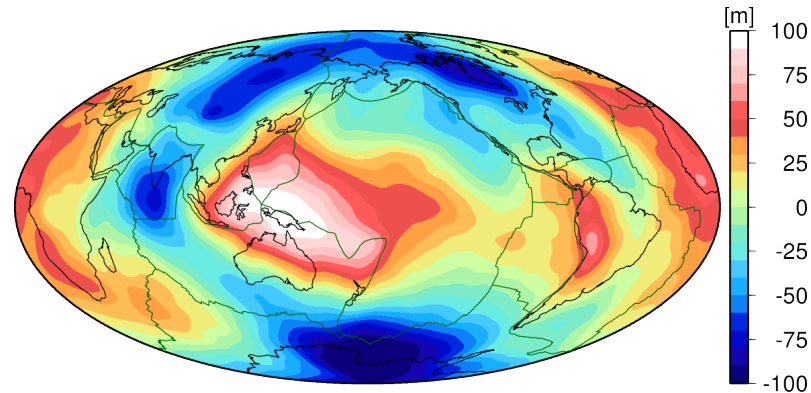
scales are saturated

Geoid anomalies

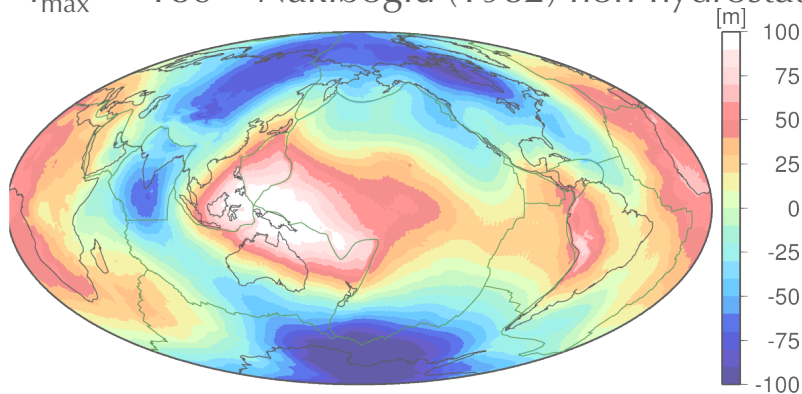
ITG-GRACE, $l_{\max} = 180$; $C_{ab}(2,0) = 0$



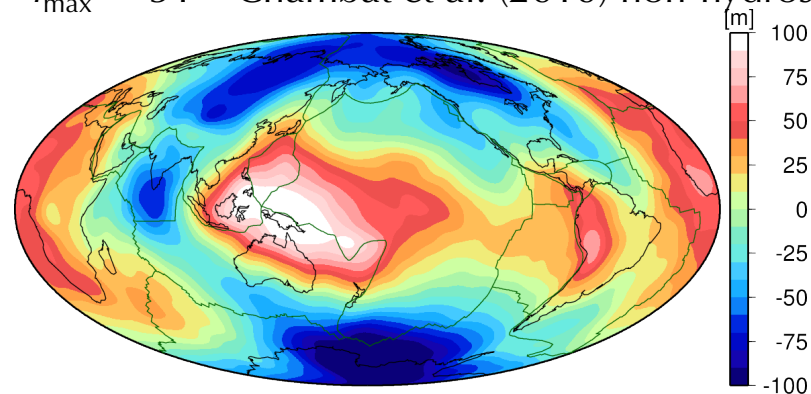
$l_{\max} = 31$ – Nakiboglu (1982) non-hydrostatic



$l_{\max} = 180$ – Nakiboglu (1982) non-hydrostatic

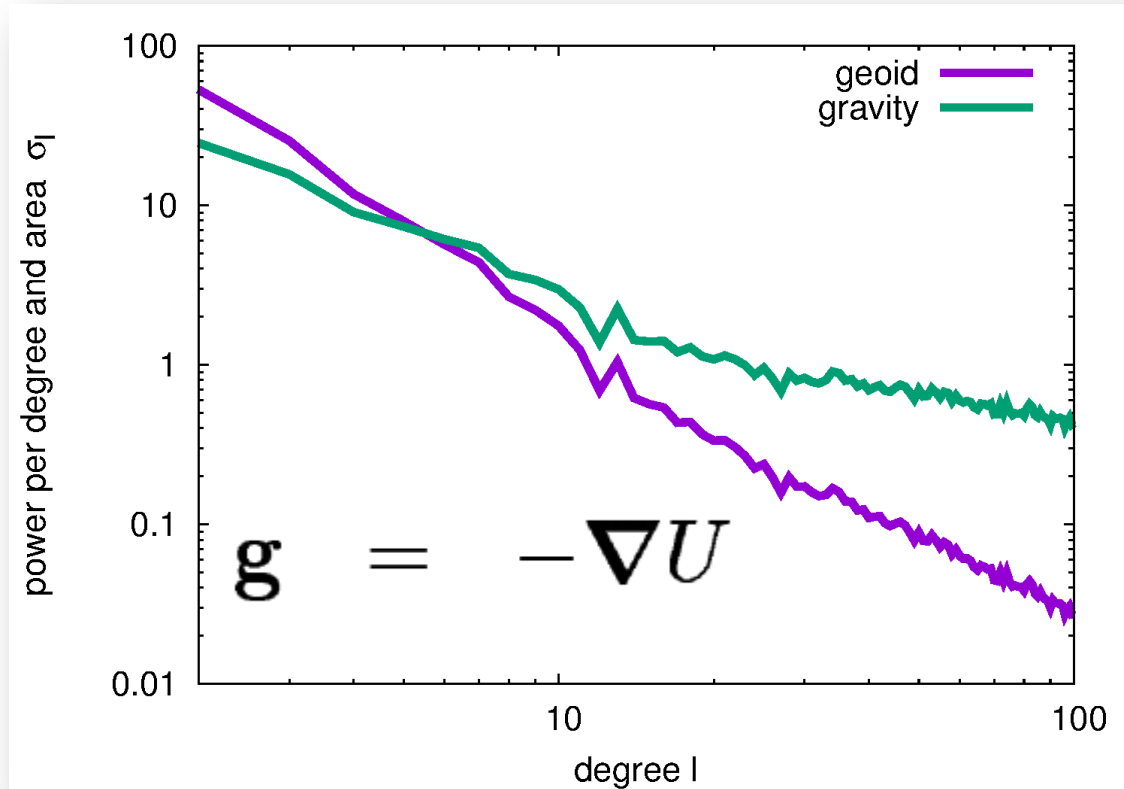


$l_{\max} = 31$ – Chambat et al. (2010) non-hydrostatic

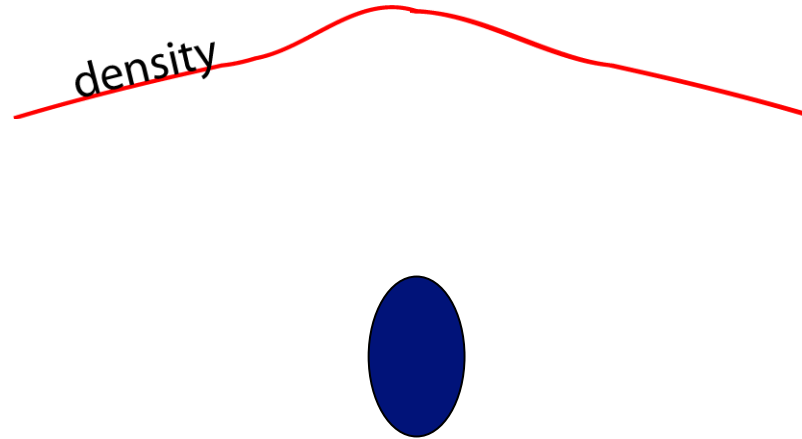


scales are saturated

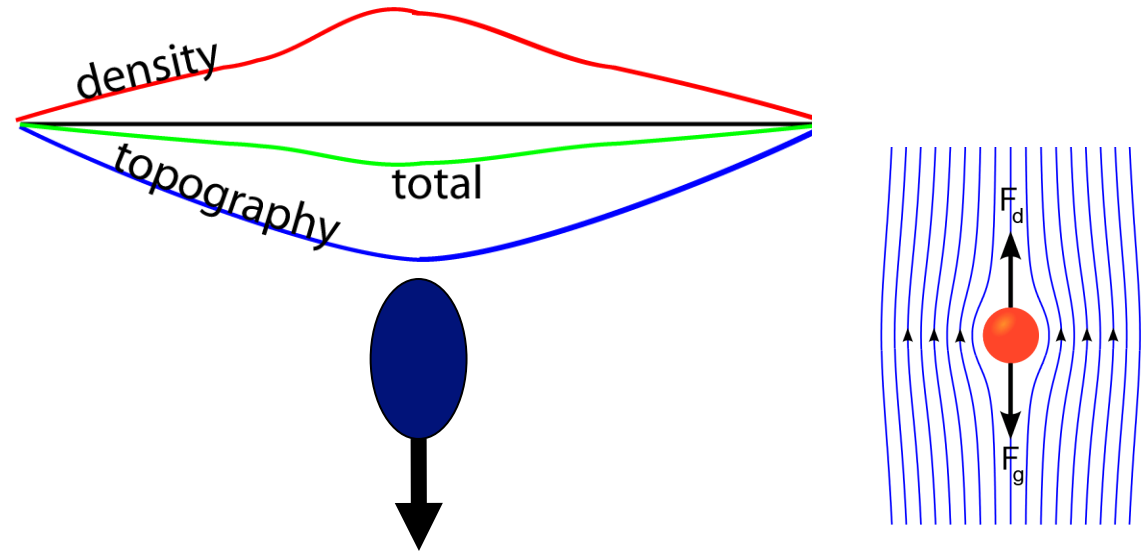
Potential field constraints: Bias in wavelength dependence of model fit



Let's fit the geoid: Static effect of dense anomaly

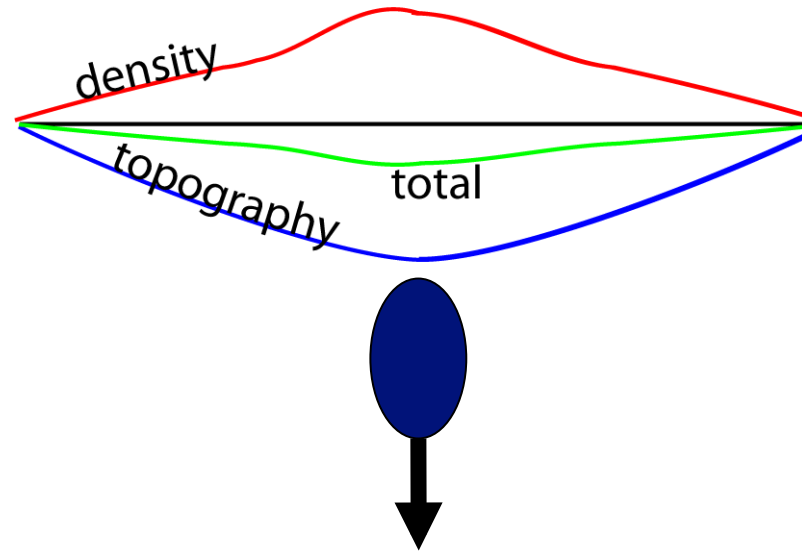
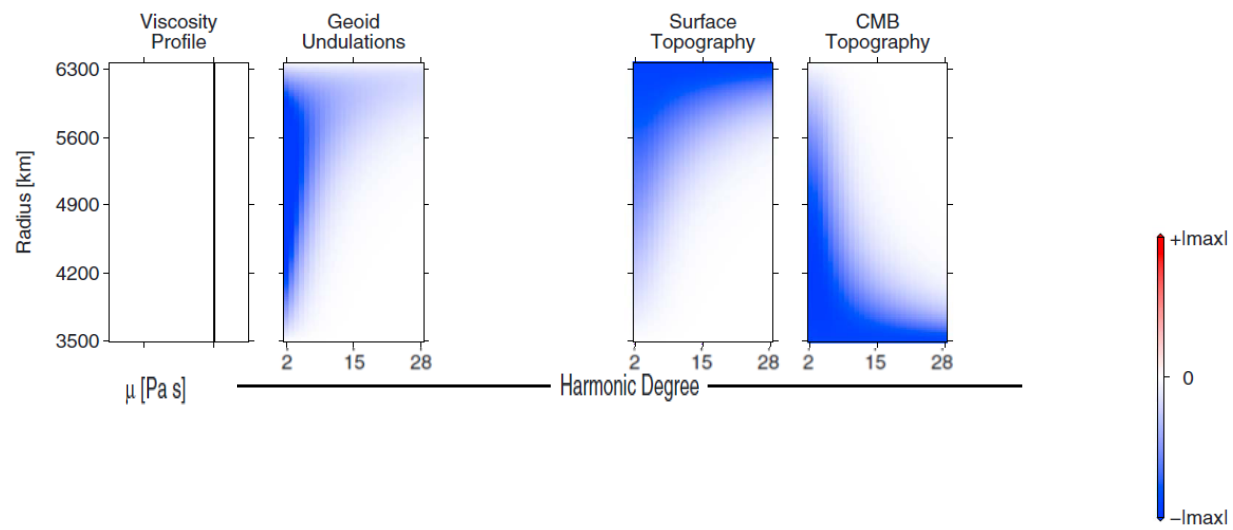


Combined static and dynamic effect of a slablet



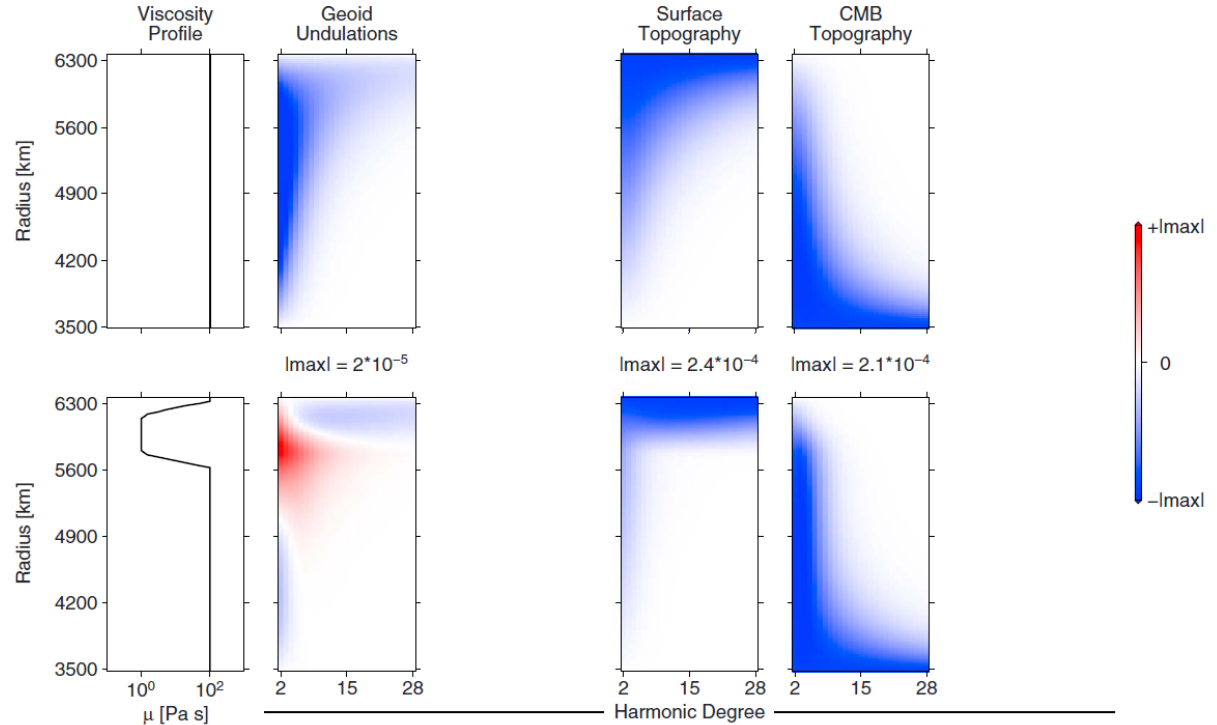
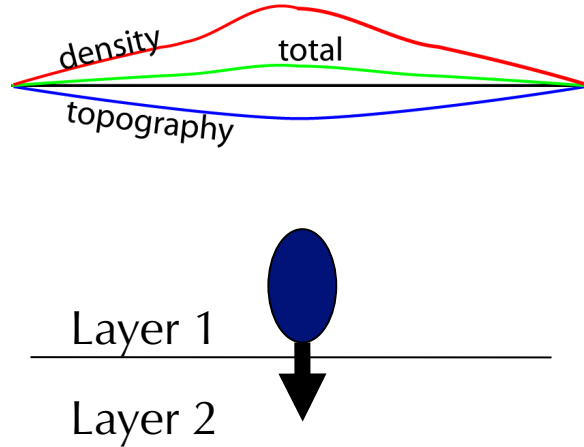
sketch from M. Billen's MYRES talk
cf. Hager (1984); Richards & Hager (1984); Ricard (1984)

response to density
anomalies as $f(z, l)$



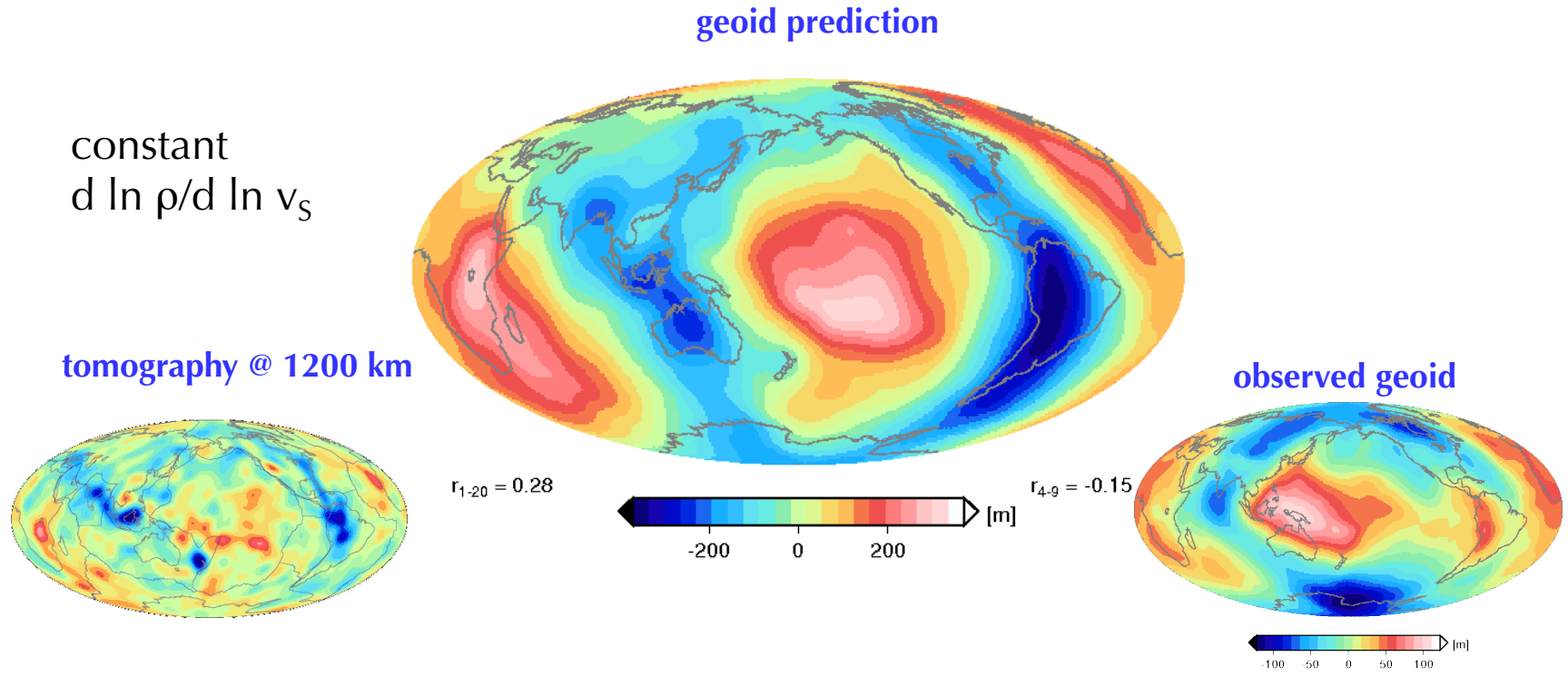
Kernels for layered viscosity mantle

Layered Viscosity



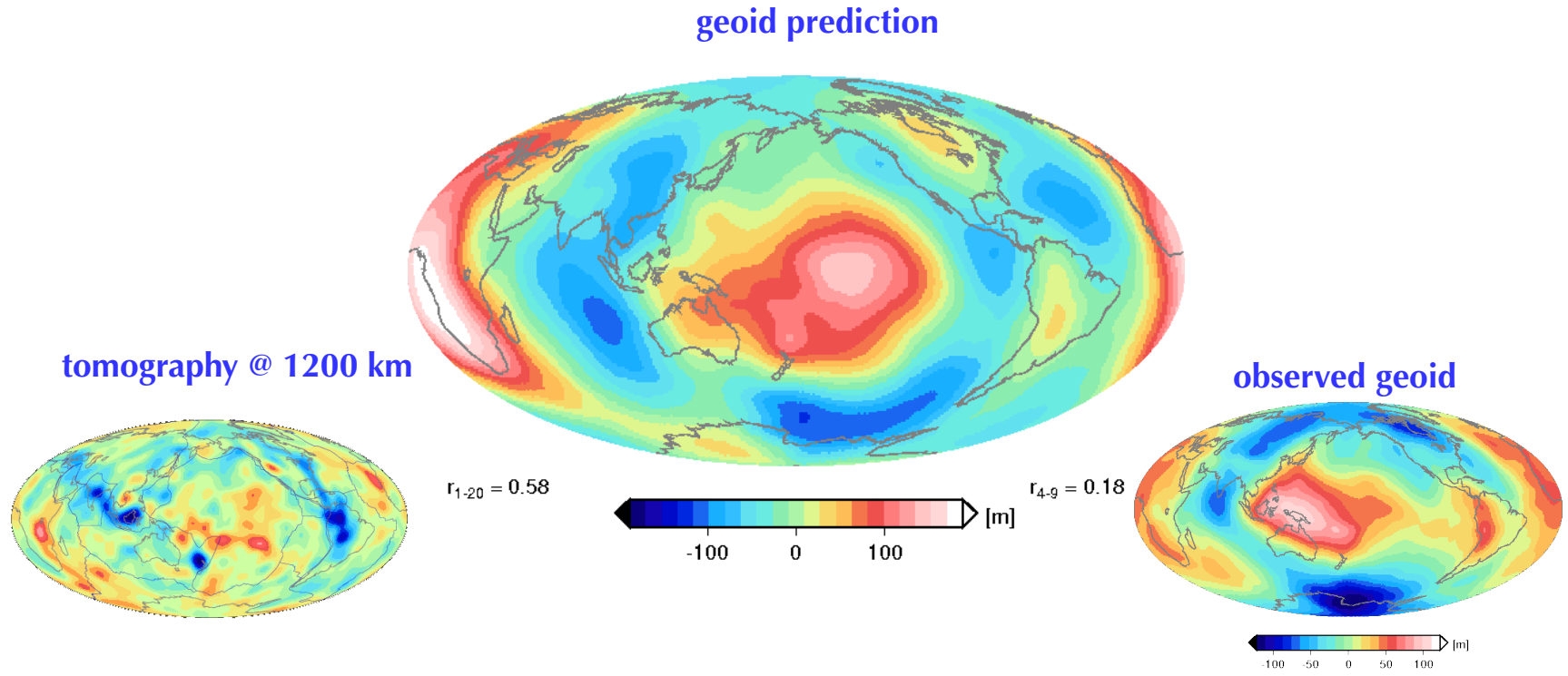
Geoid for tomography driven flow

- Isoviscous – free slip surface boundary condition



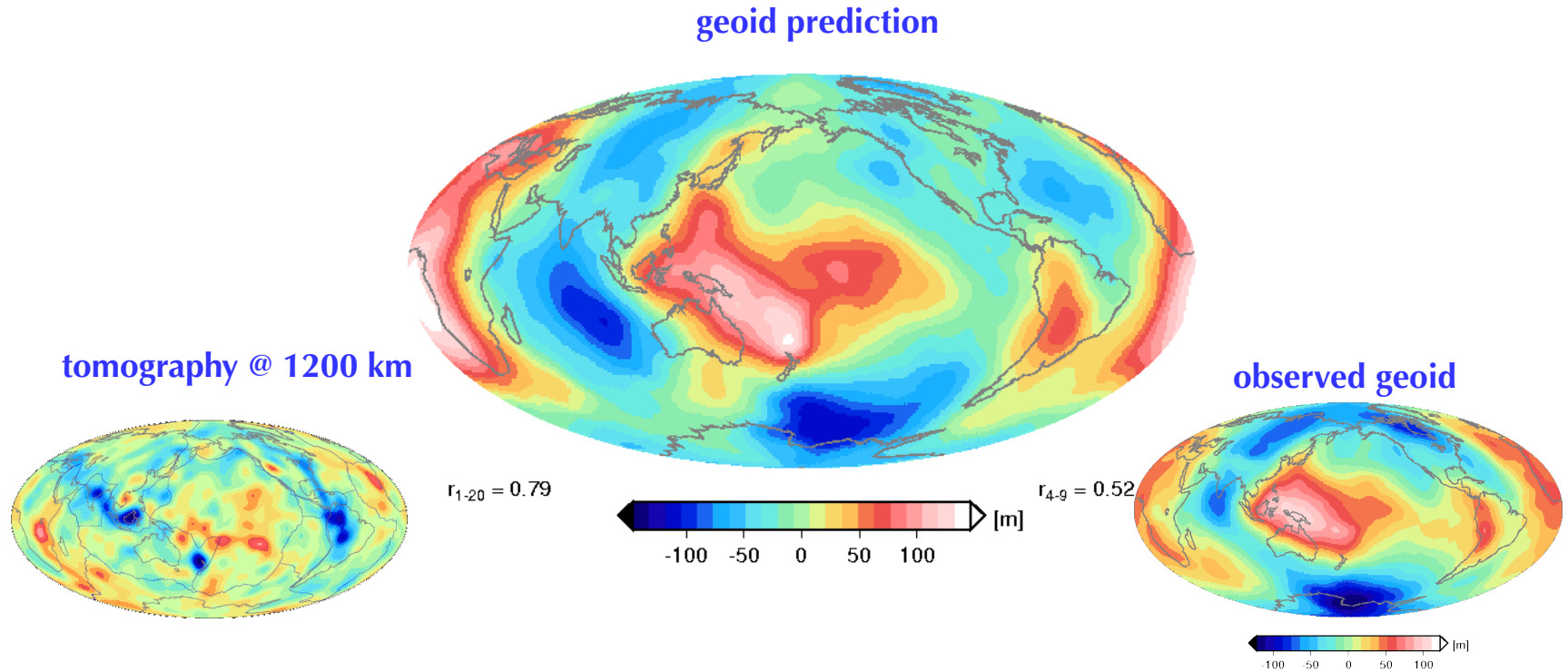
Geoid for tomography driven flow

- Lower mantle viscosity increase



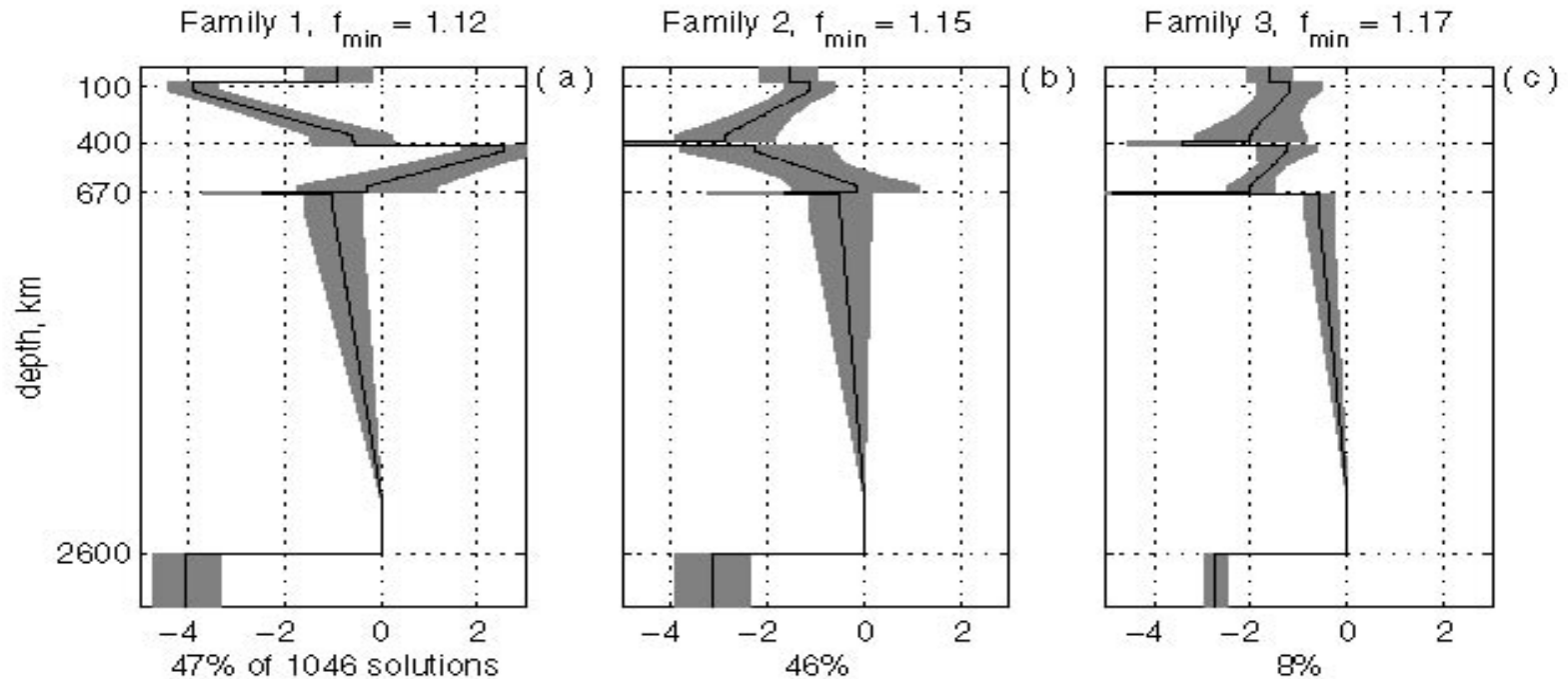
Geoid for tomography driven flow

- Four layer model



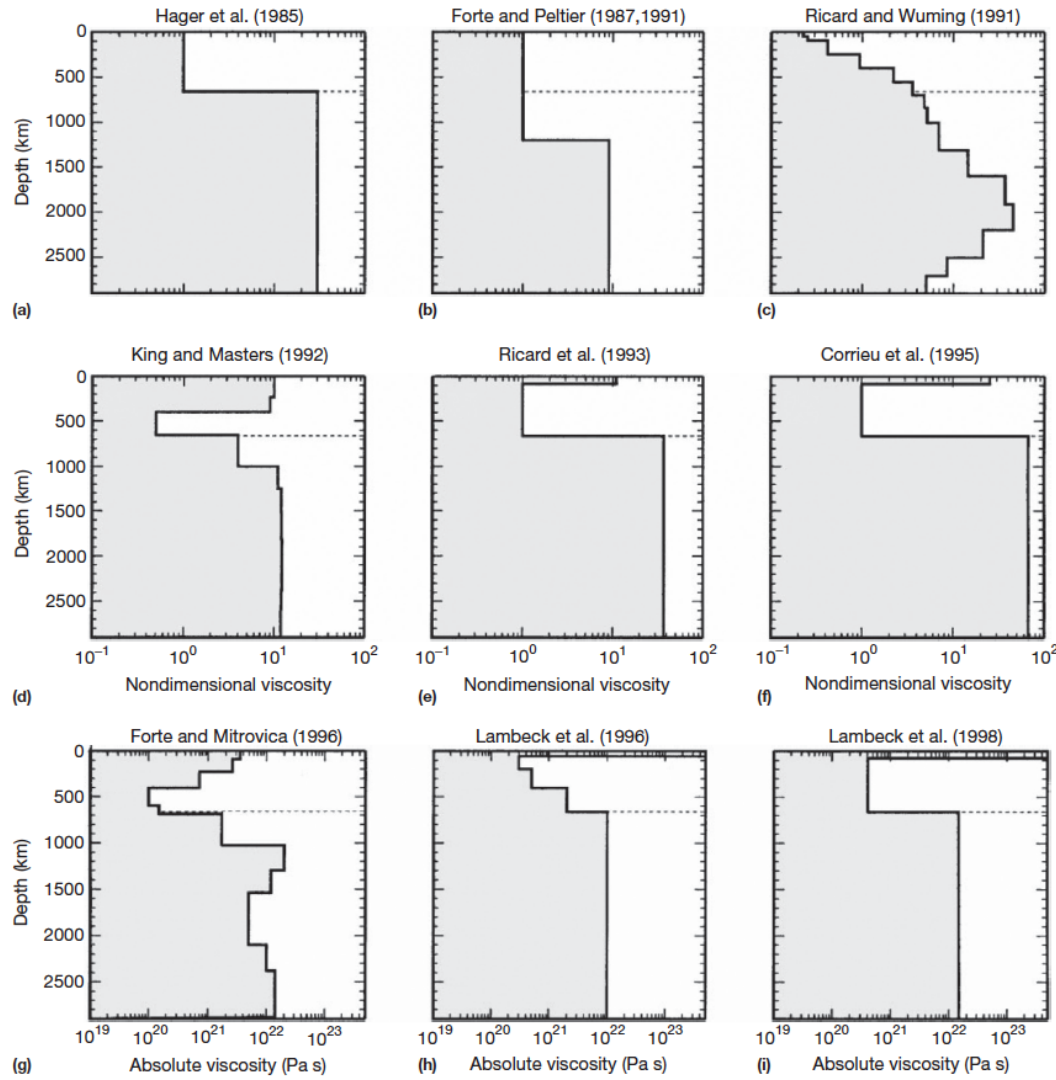
Viscosity inversions are non-unique

(Monte Carlo approach, based on geoid and surface dynamic topography)



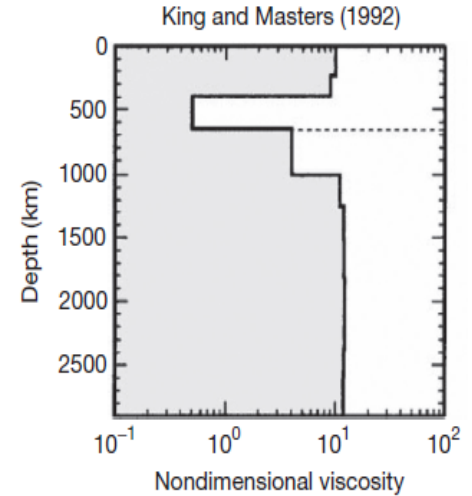
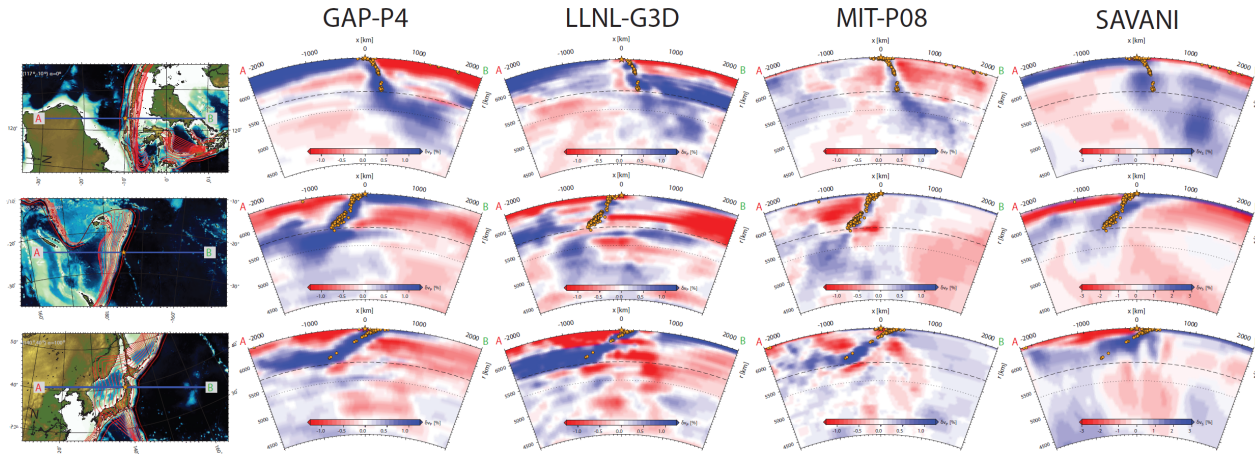
Geoid vs. GLA inversions

- still need to reconcile ice load model viscosities



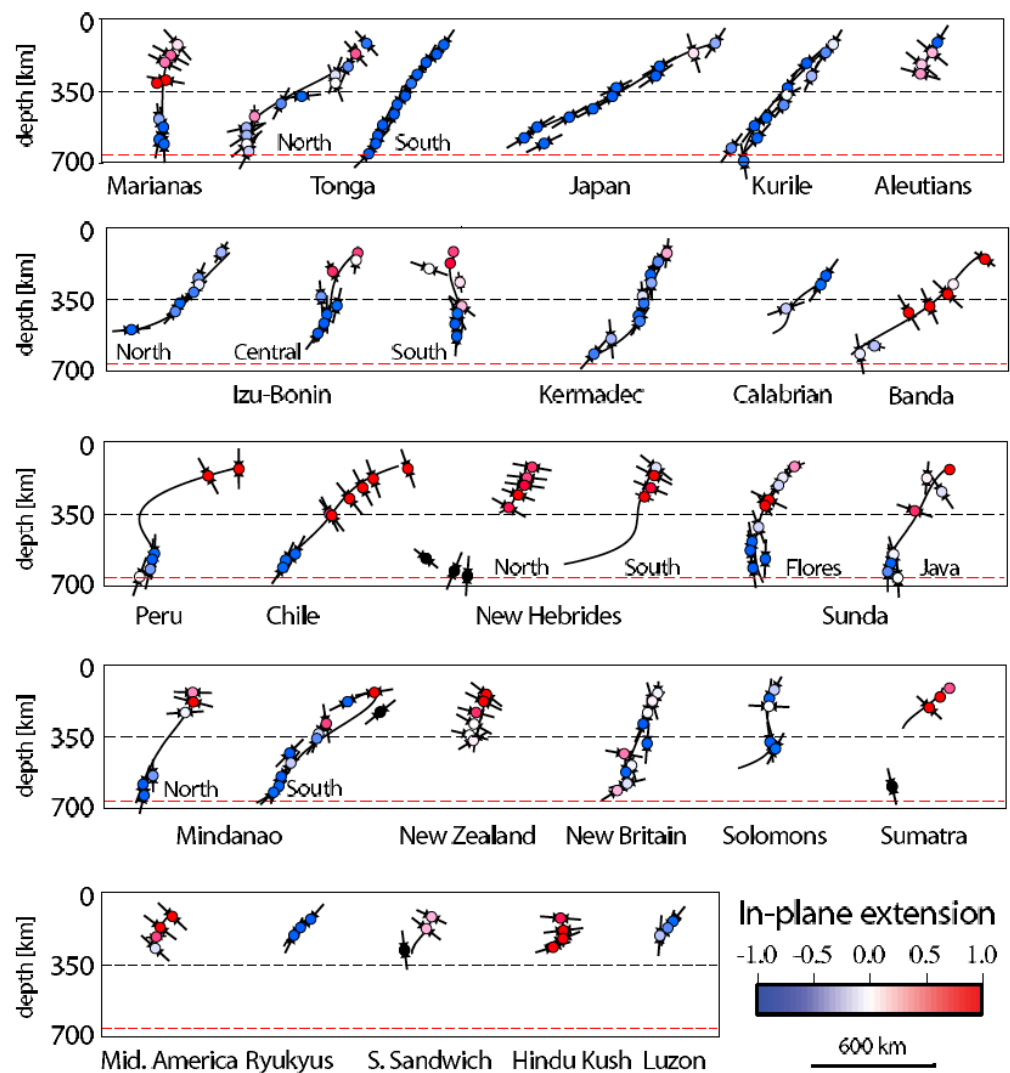
Forte (2015)

Slab ponding diversity and viscosity stratification in transition zone

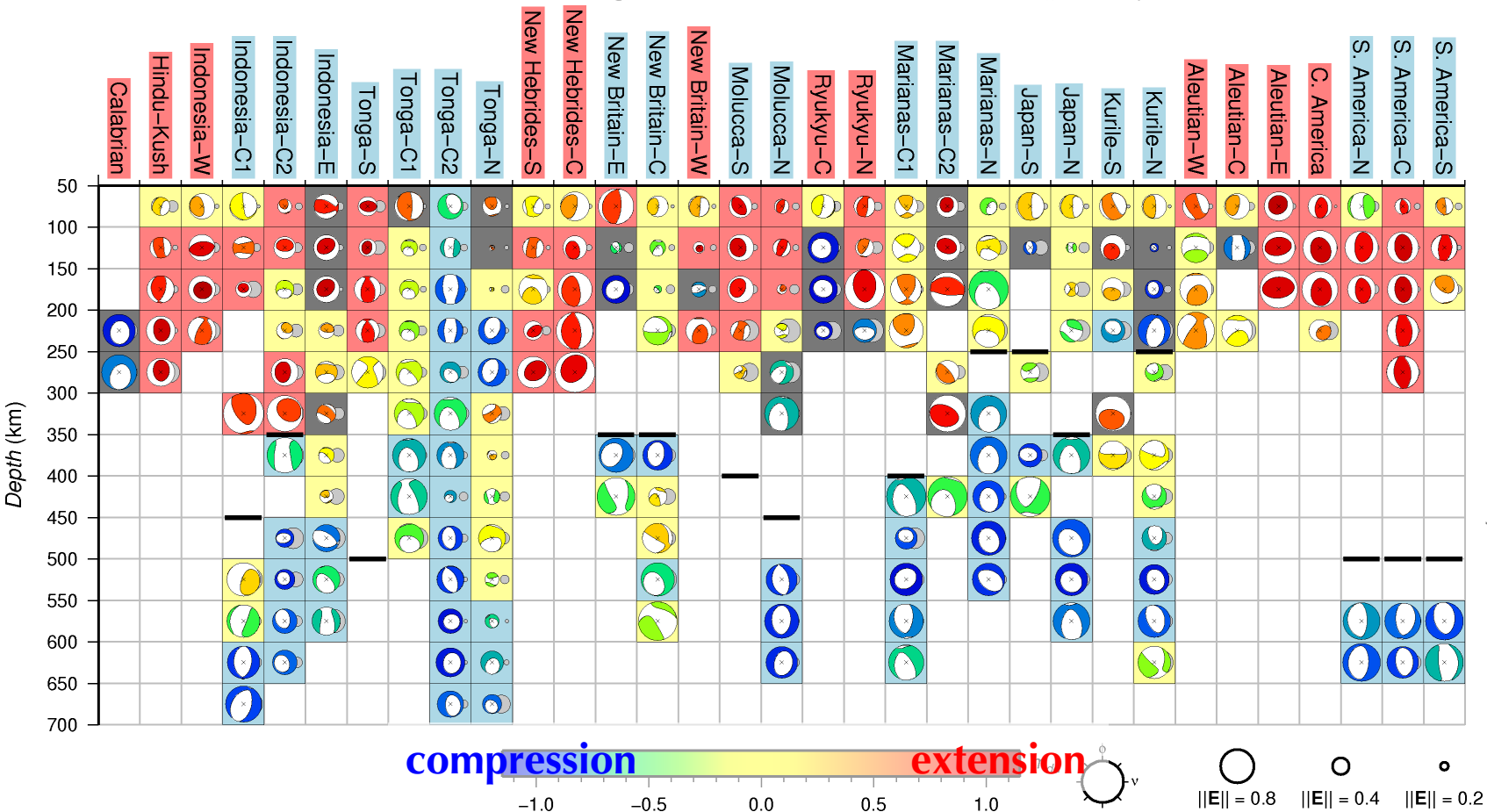


Earthquake constraints for slab and mantle viscosity

Alpert et al. (2010)
cf. Isacks and Molnar (1971)



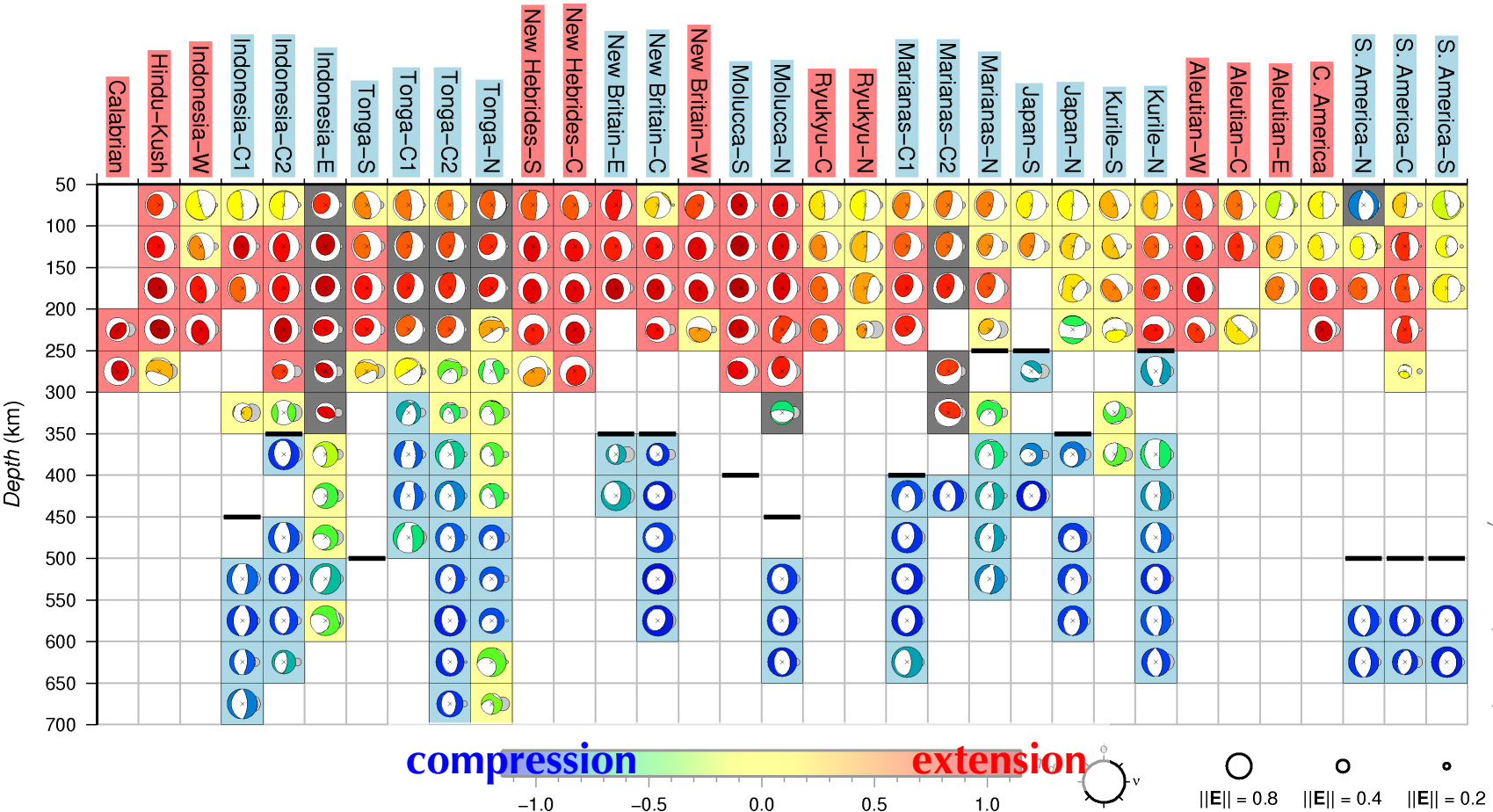
Summed gCMTs (in-slab projection)



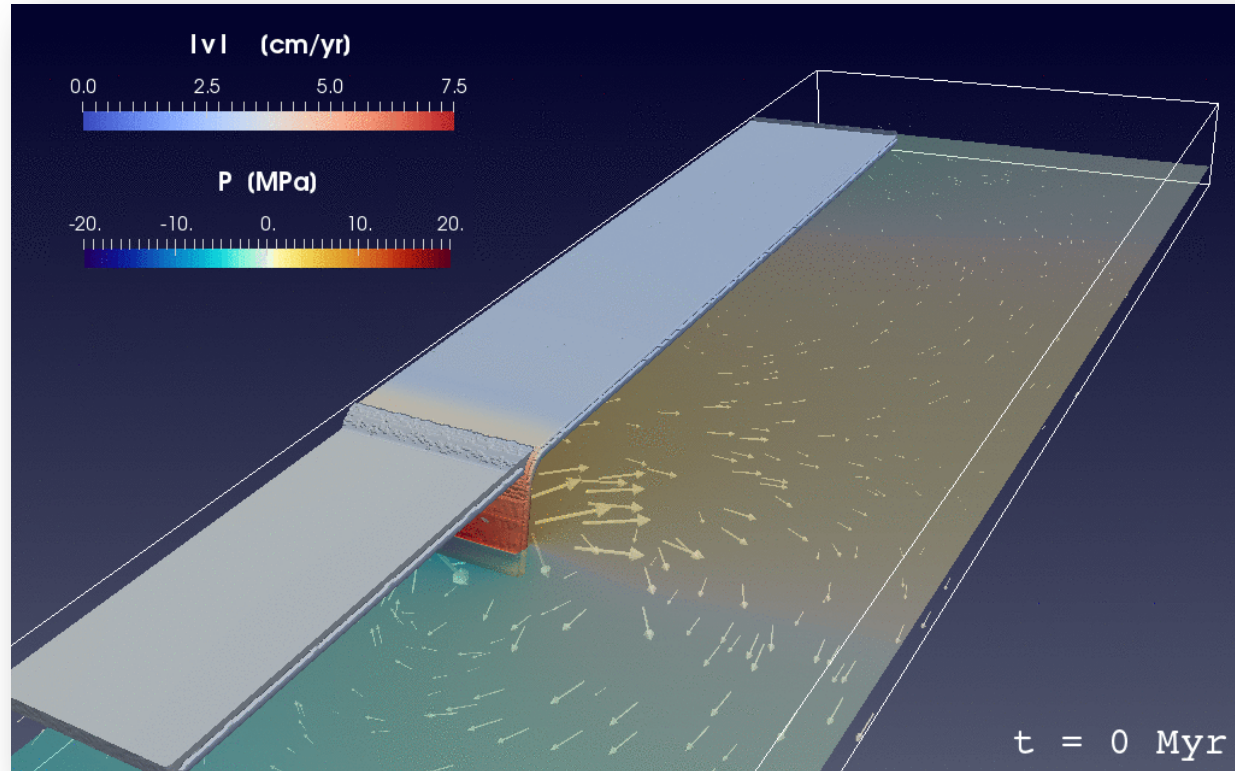
Bailey et al. (2012)

cf. Vassiliou and Hager (1998), Billen and Gurnis (2003),
 Alsic et al. (2010), Alpert et al. (2010)

Bailey et al. (2012)

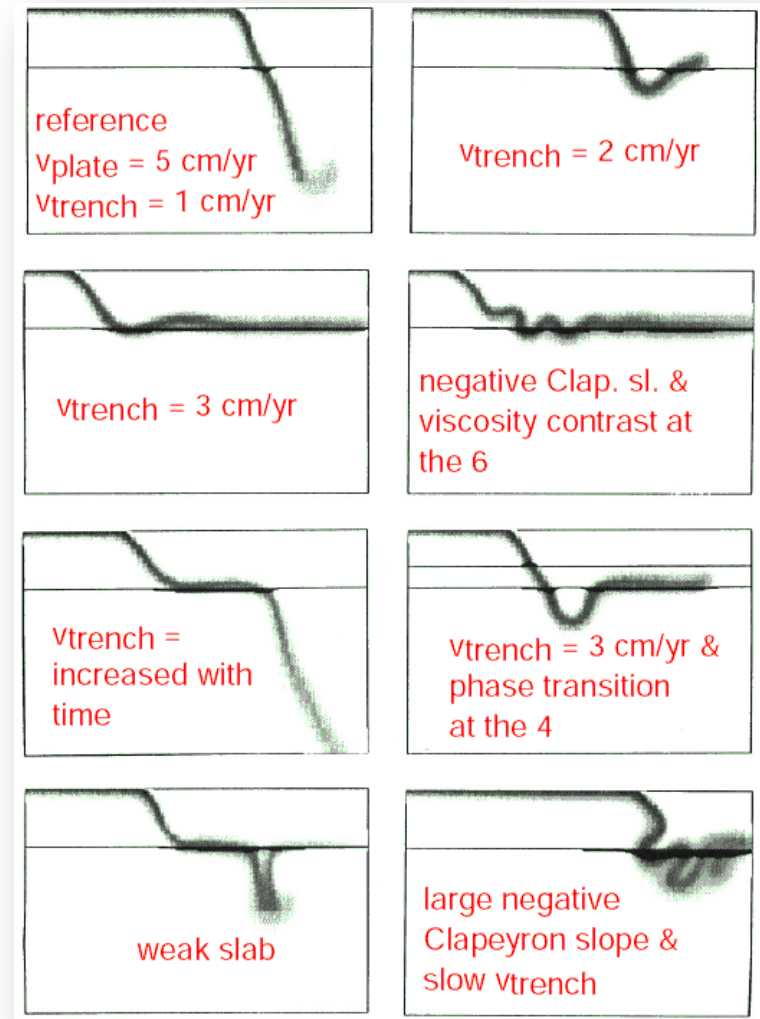
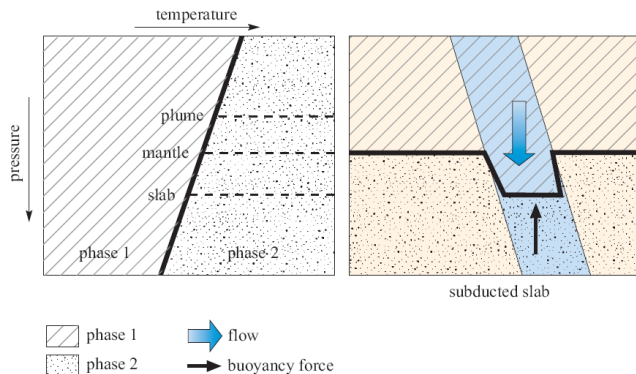


Regional subduction dynamics (let's do our own experiments)

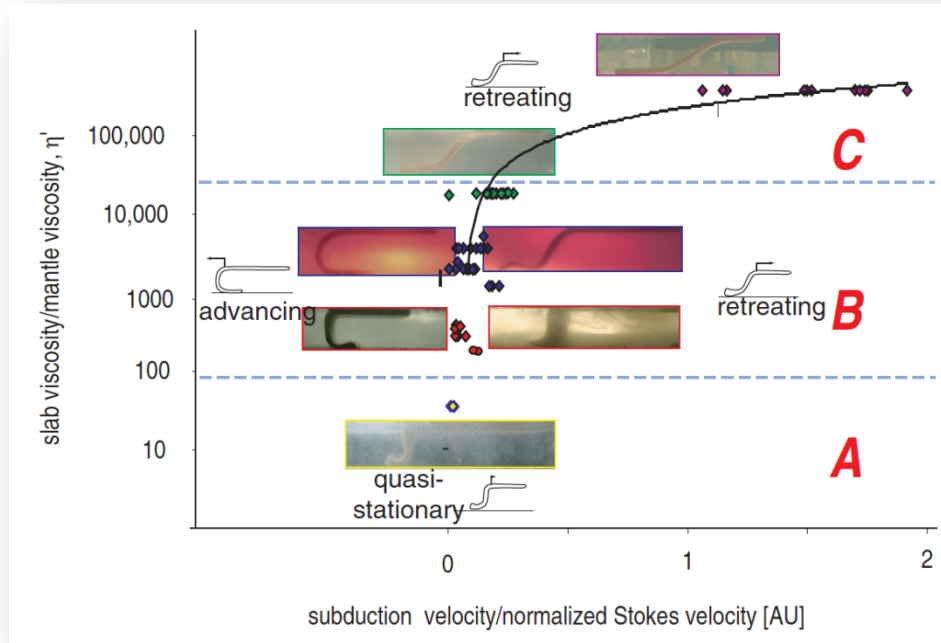


Trench rollback and slab penetration

- kinematic (prescribed velocities)
model varying plate/trench motion
partitioning and importance of phase
transition (v_{trench} vs. v_{Stokes})
- Δp due to deeper phase transition for
negative Clapeyron slope

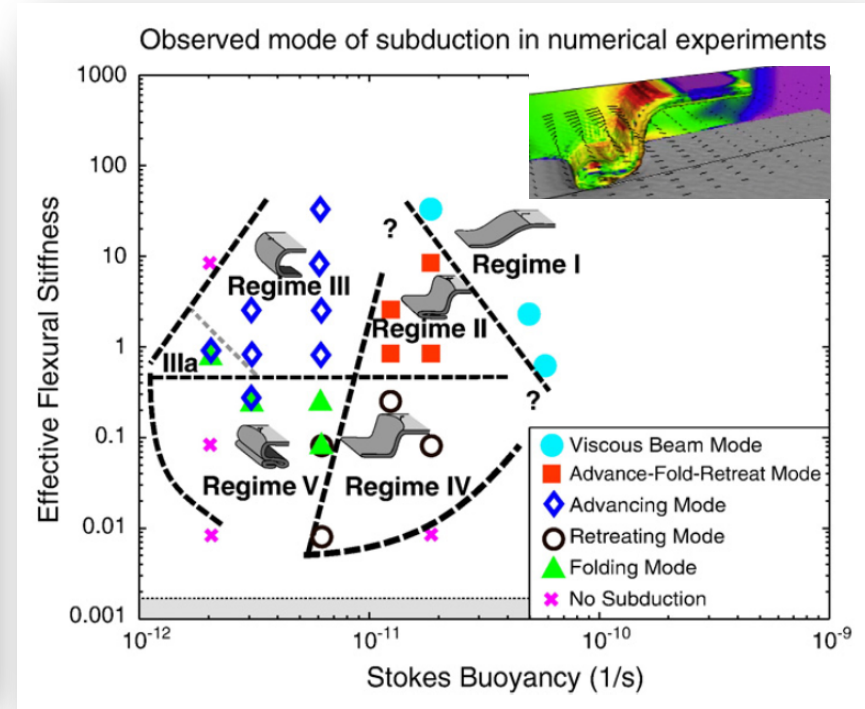


Rollback dynamics phase diagrams



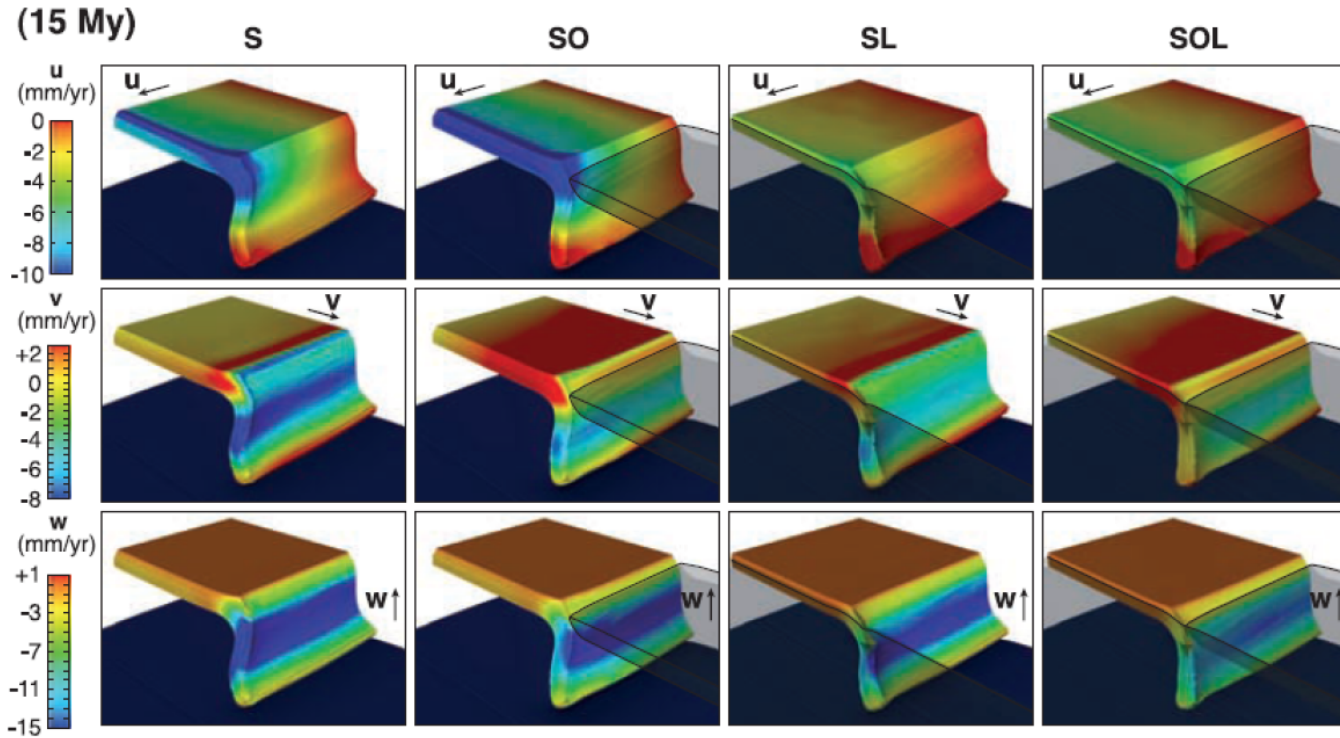
Funiciello et al. (2007)

cf. Ribe (2010)

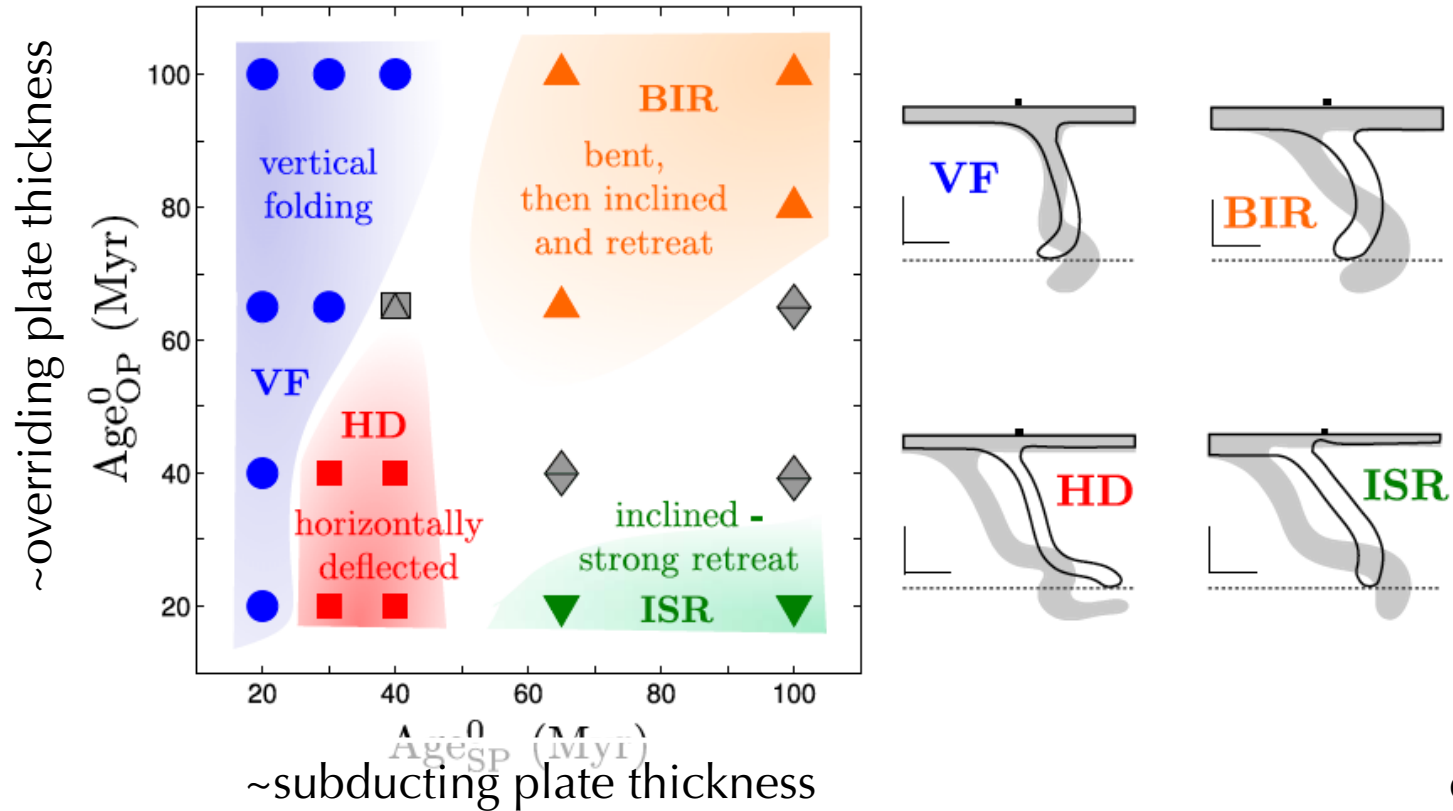


Stegman et al. (2010)

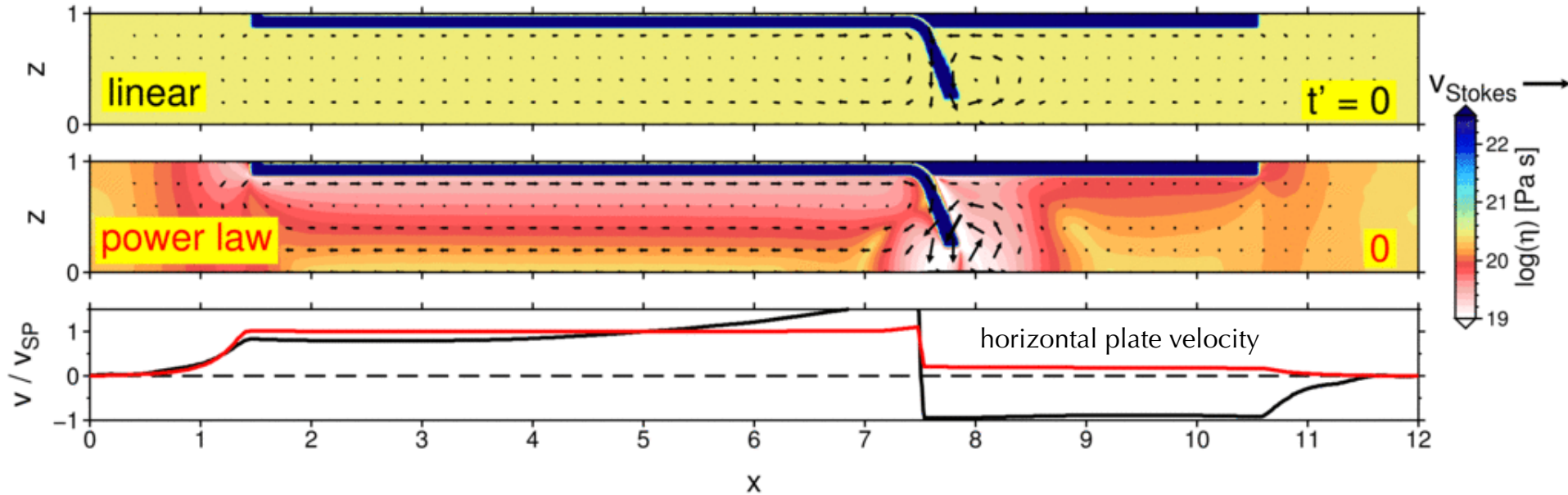
Annoyances I: The overriding plate



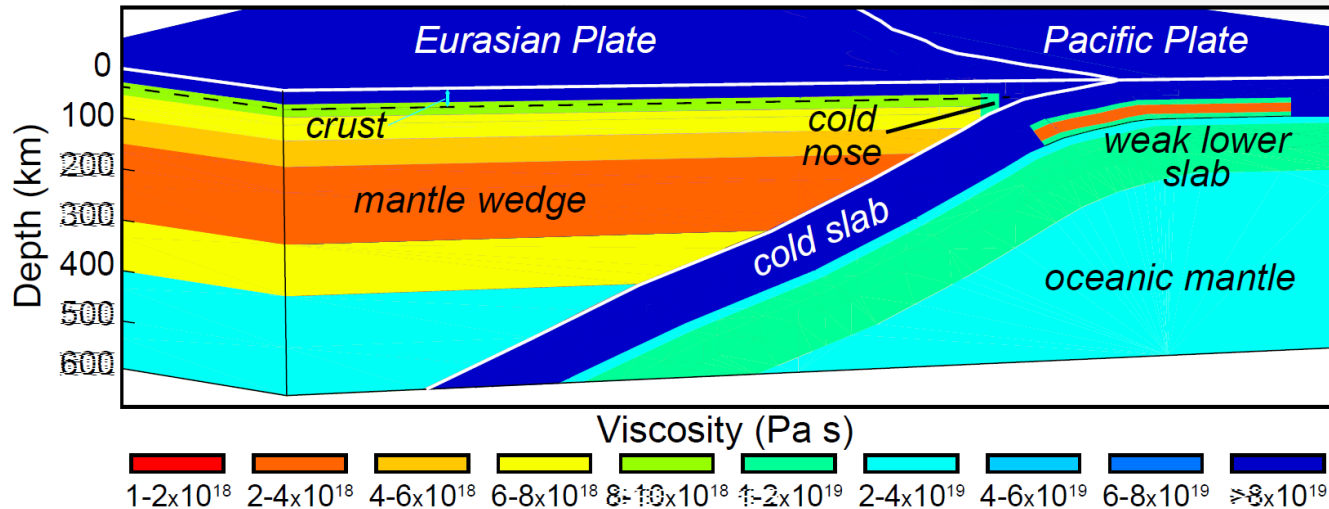
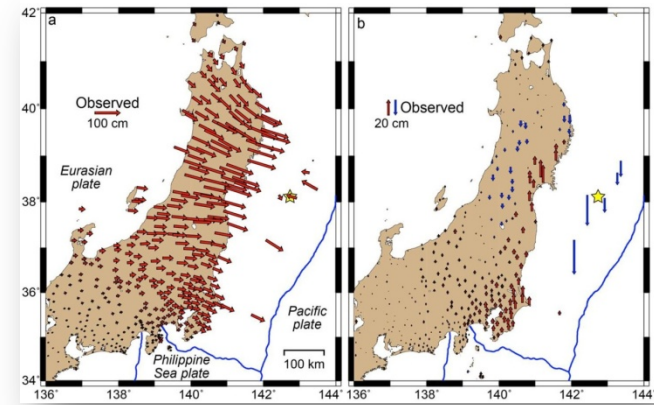
New regime diagrams for trench motions and ponding



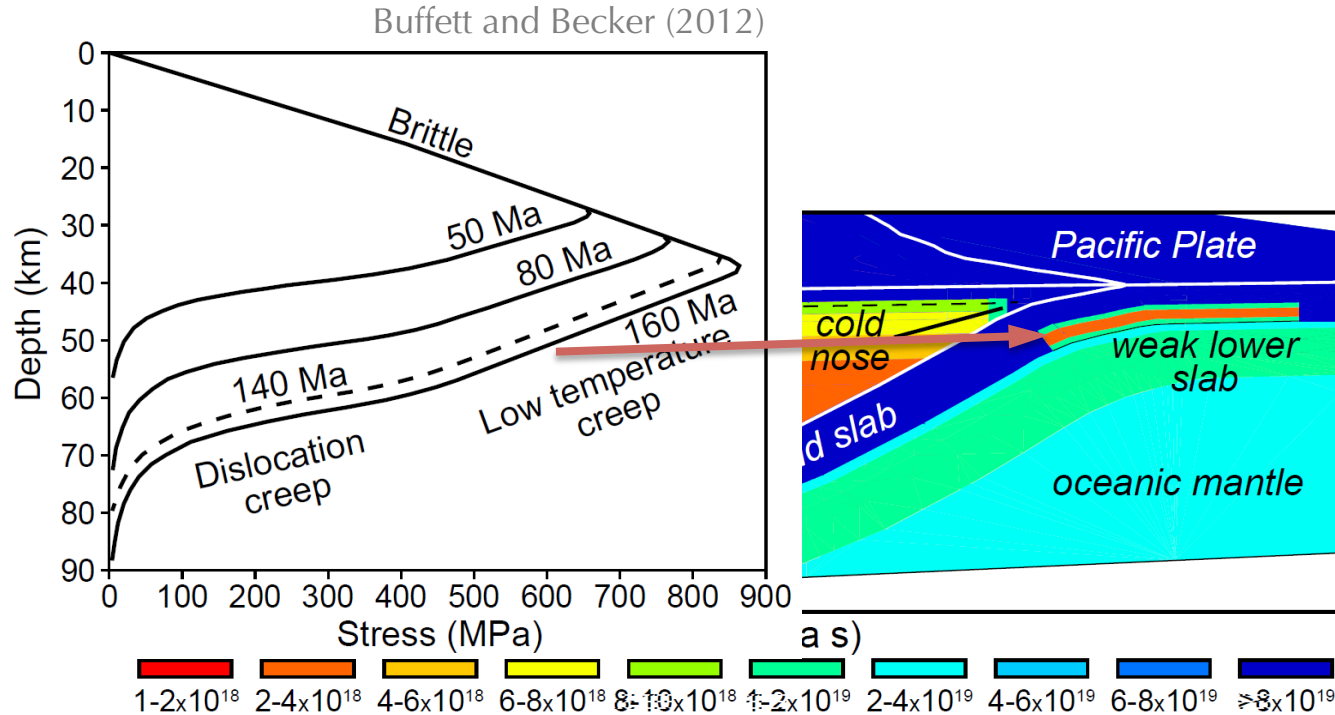
Annoyance II: Trench motions for nonlinear mantle rheology



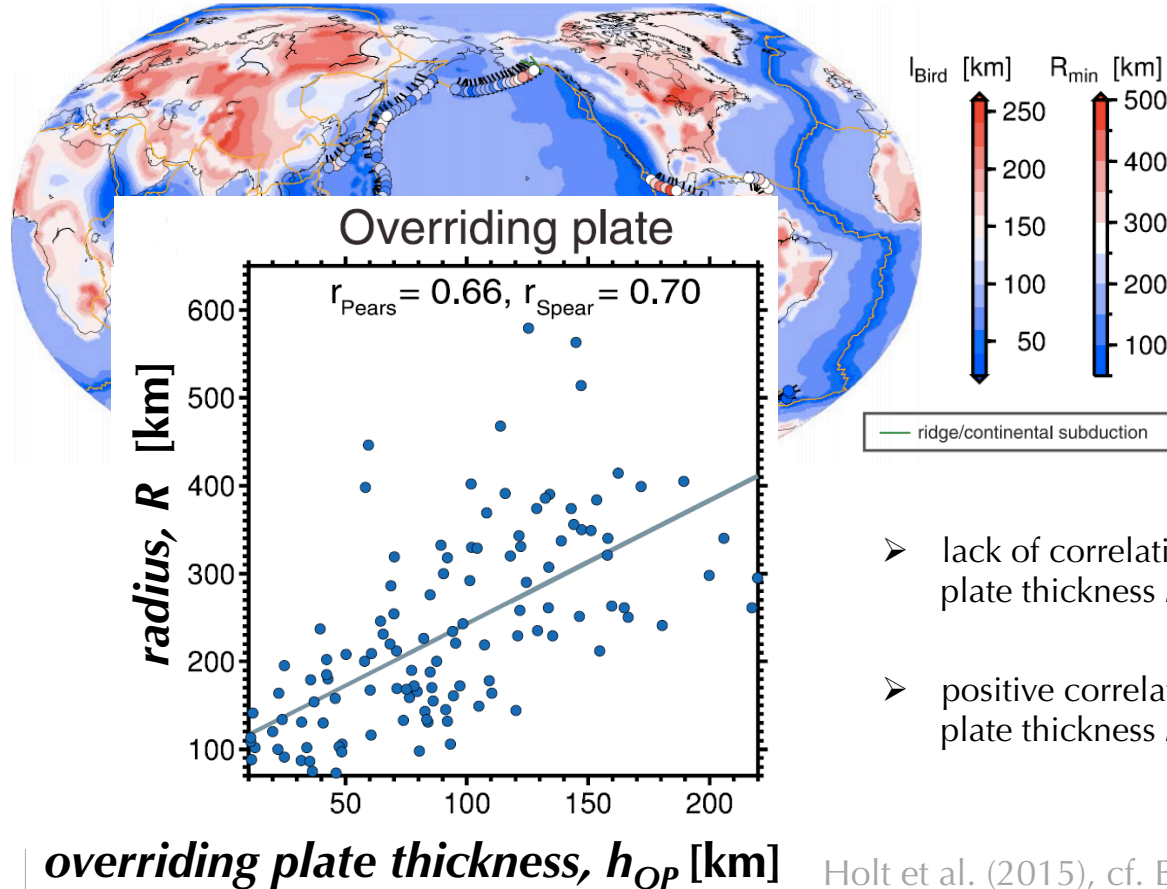
Large-scale rock mechanics experiment: Infer slab and mantle rheology from post-seismics (Tohoku-oki 2011 M9)



Plastic slab I: Weakening of oceanic plate by Peierls creep?



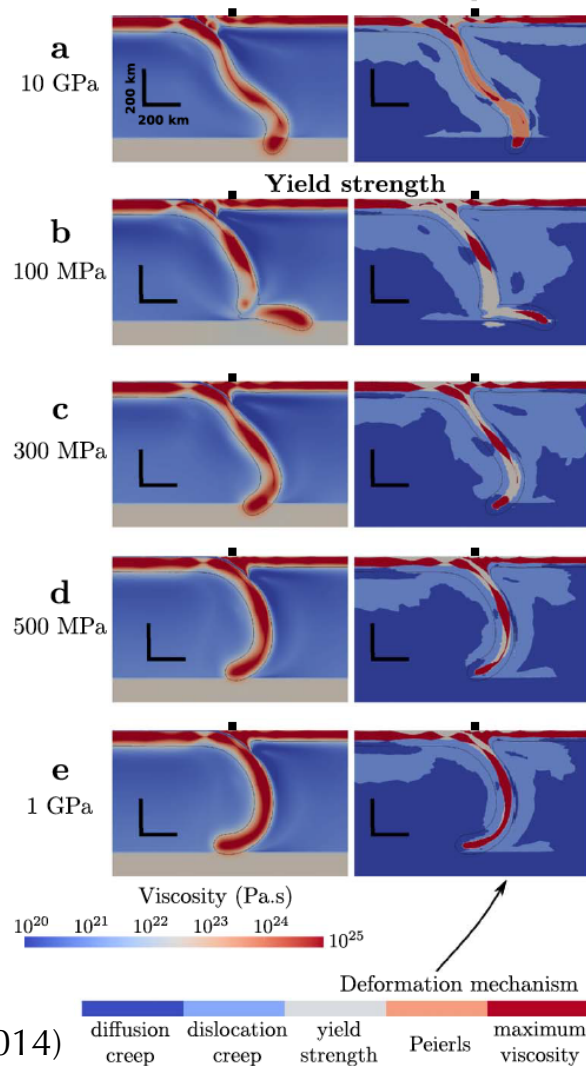
Plastic slab II: Overriding plate thickness controls bending radius for plastic, not for viscous plate – Earth behaves like that



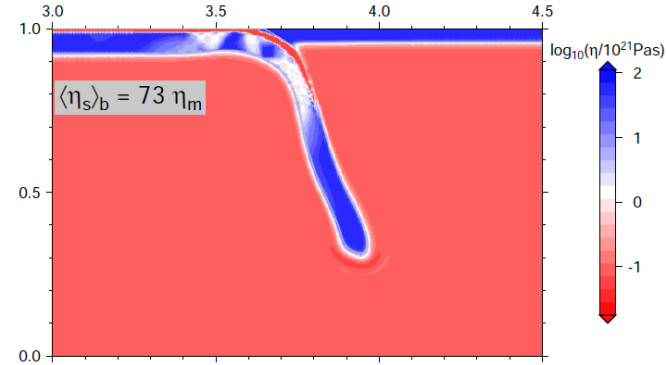
- lack of correlation between subducting plate thickness h_{SP} and radius R
- positive correlation between overriding plate thickness h_{OP} and radius R

Annoyance III: Subduction with plastic rheology

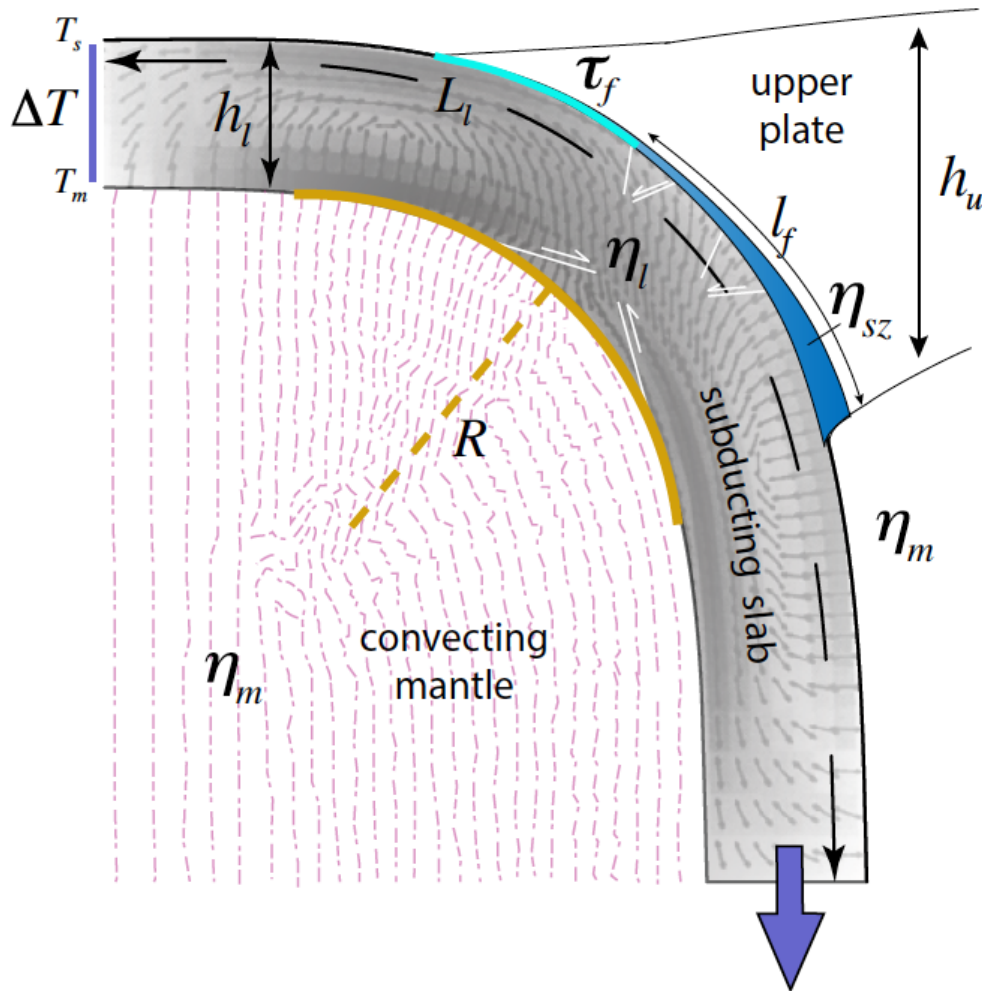
- significant oceanic plate viscosity reduction, perhaps $\eta_{\text{slab}} \sim 100 \eta_{\text{upper mantle}}$



Garel et al. (2014)



Holt et al. (2015)



Work balance: viscous dissipation
vs. potential energy release

slab sinking due to thermal density contrasts

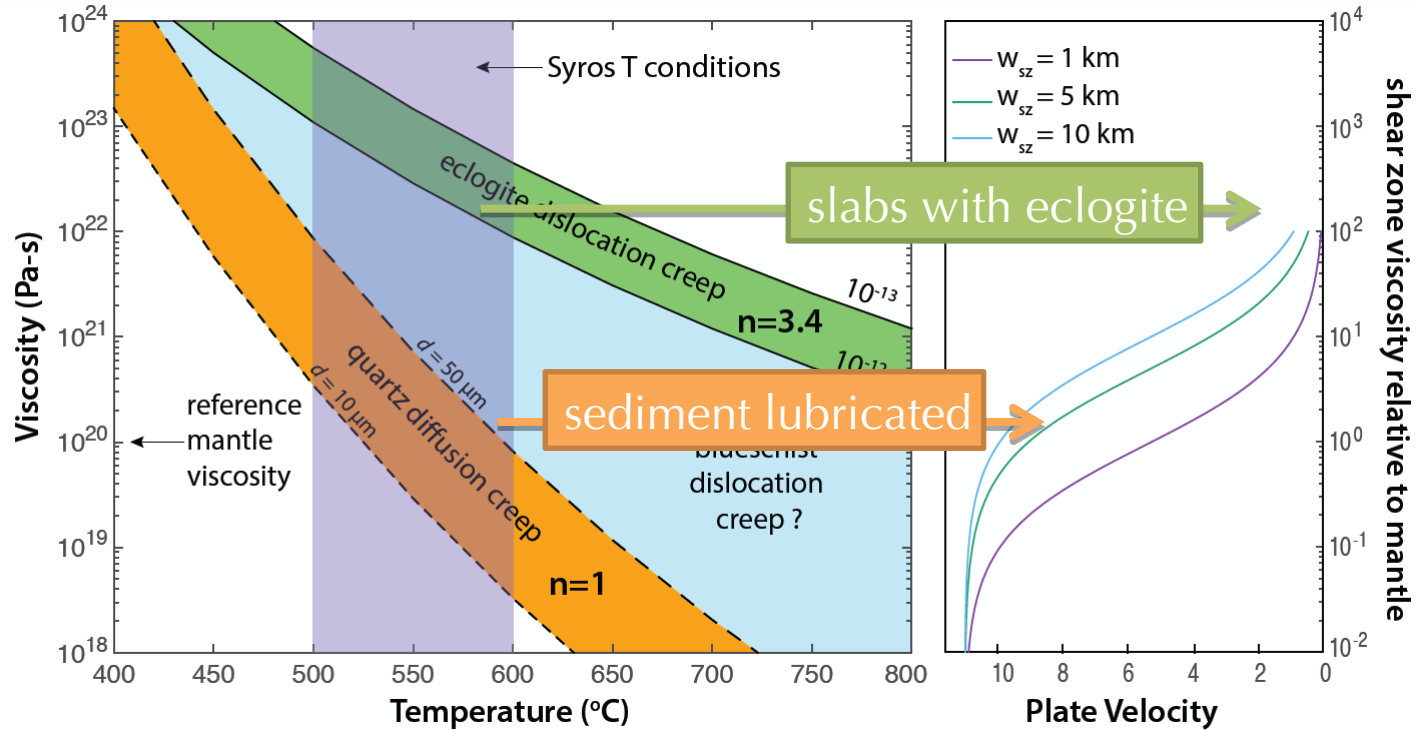
shear resistance along slab interface

$$V_p = \frac{C_s \rho g \alpha \Delta T l_s h_s - C_f \tau_f l_f}{3\eta_m (A + C_m) + C_l \eta_l (h_l / R)^3}$$

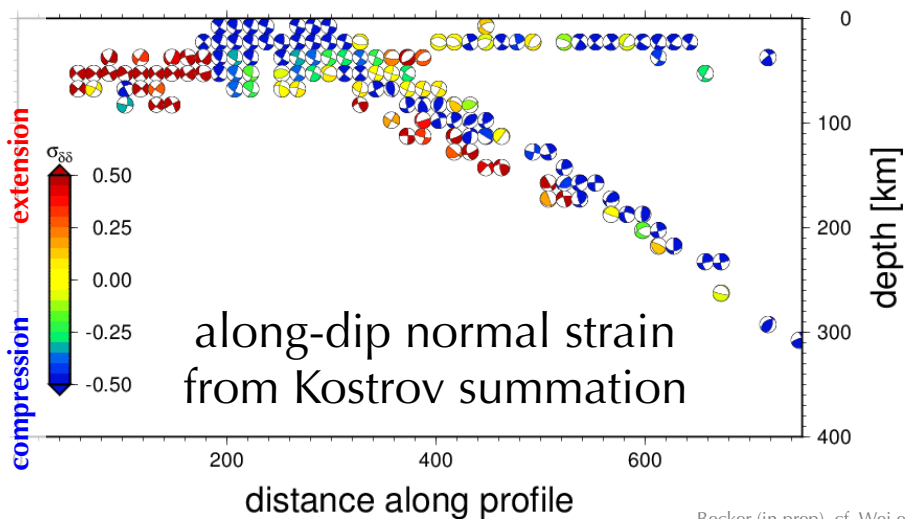
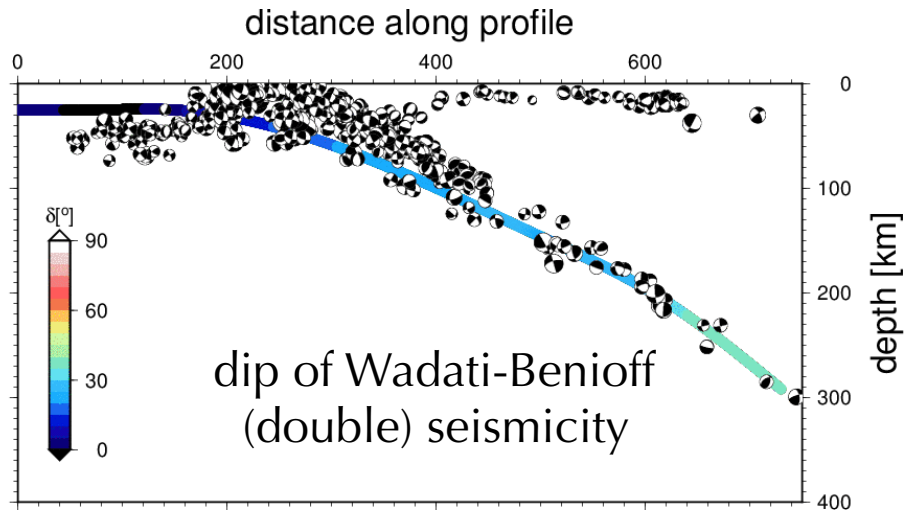
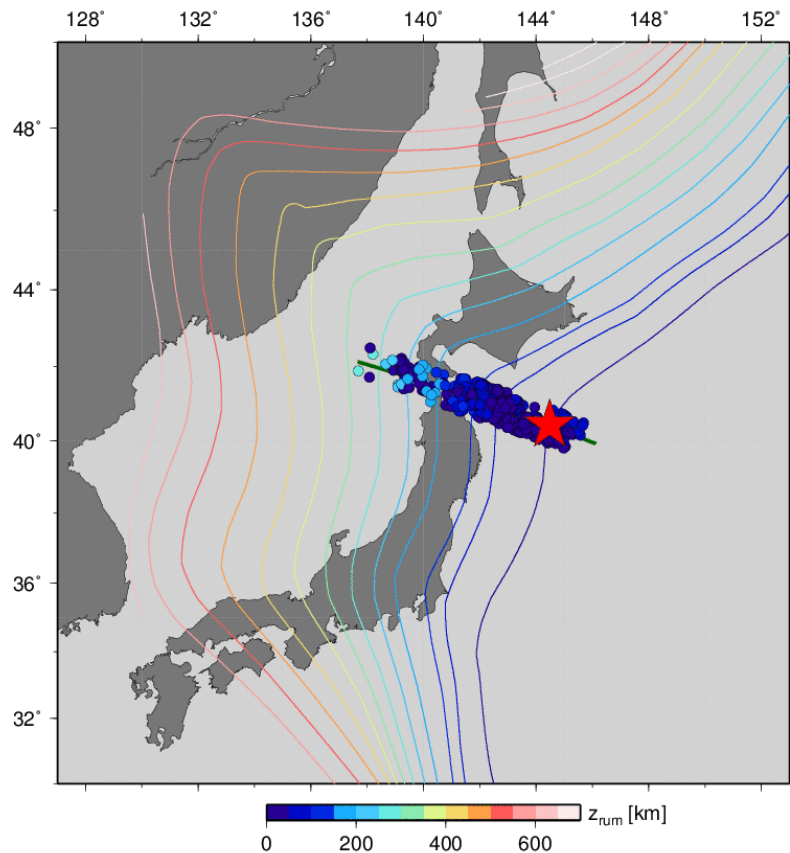
dissipation into convecting mantle

dissipation into slab due to bending

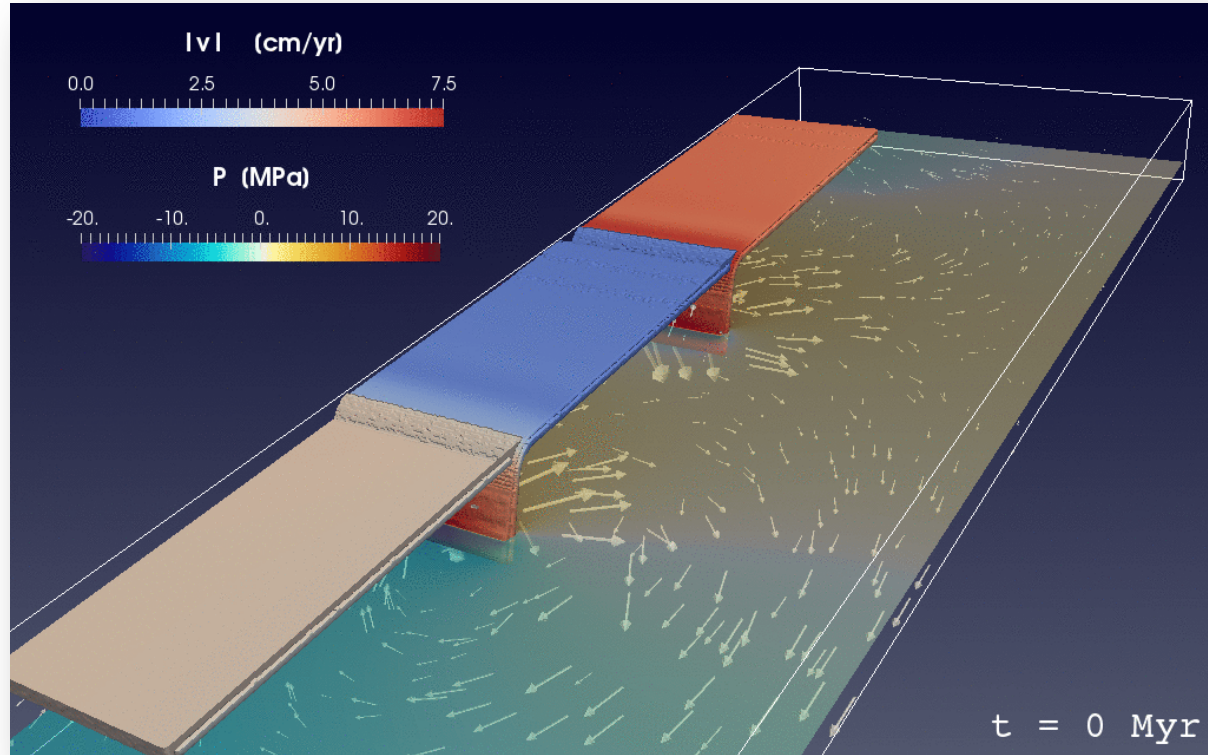
Annoyance IV: Interface (i.e. geology) control on plate velocities (continental erosion, long term cycles)?



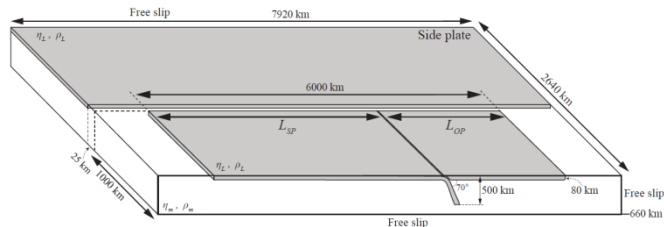
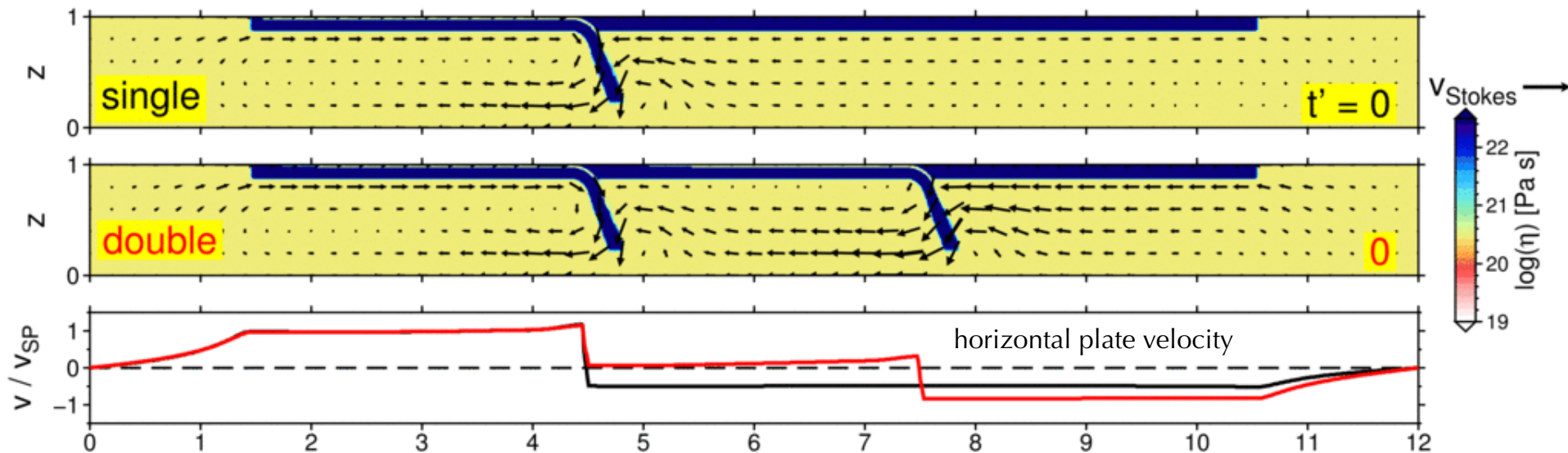
Regional CMTs (F-Net)



Slab – slab interactions

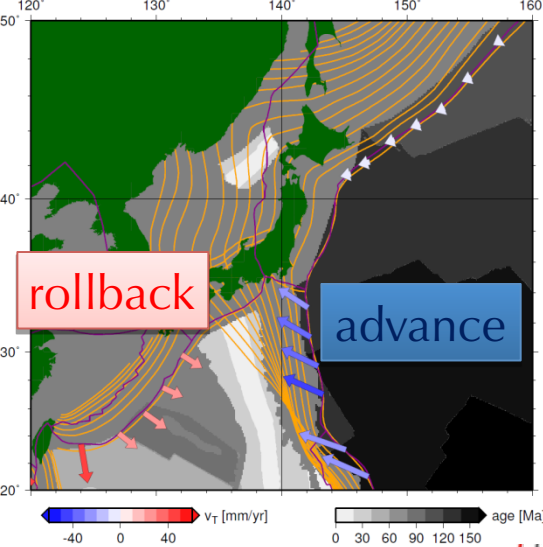


Trench motions for slab – slab interactions

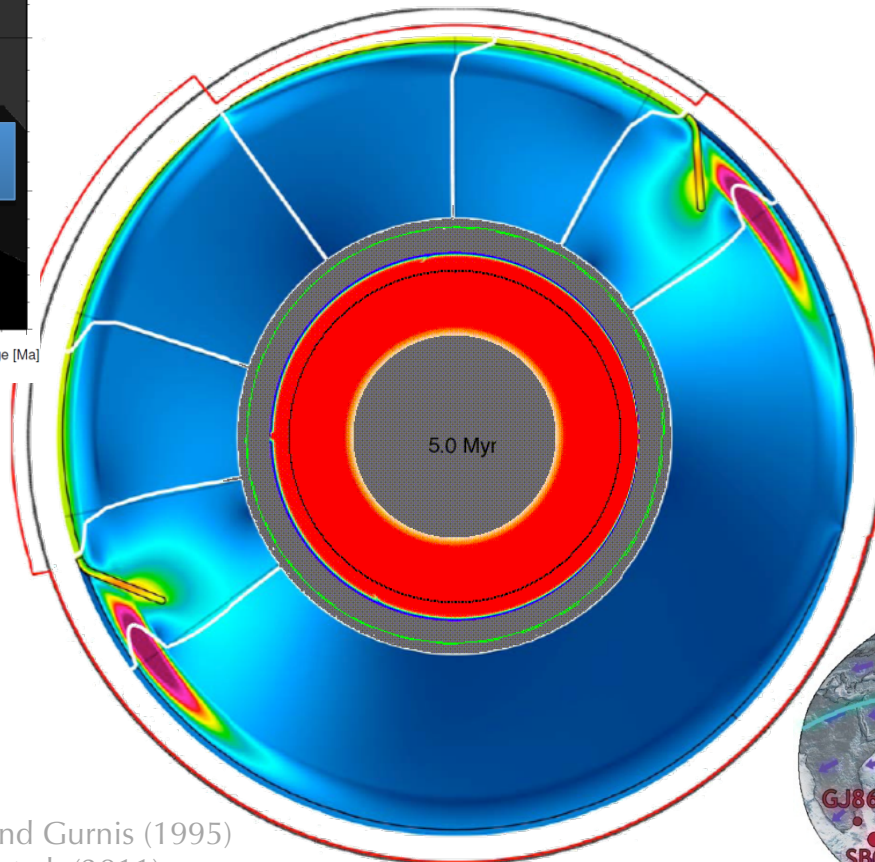


$$1 + 1 \sim 2$$

Global interactions: Trench motions and reference frames

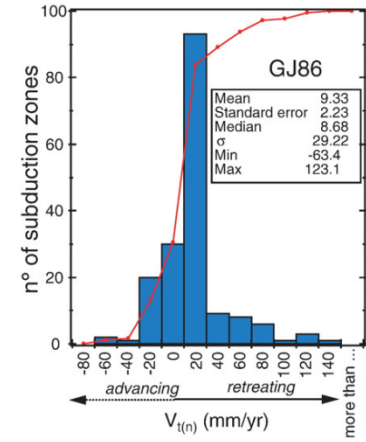


Becker et al. (2015)

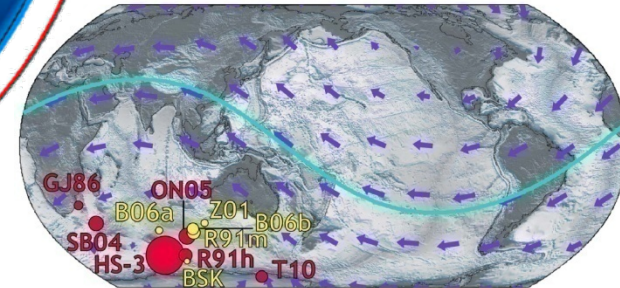


Zhong and Gurnis (1995)

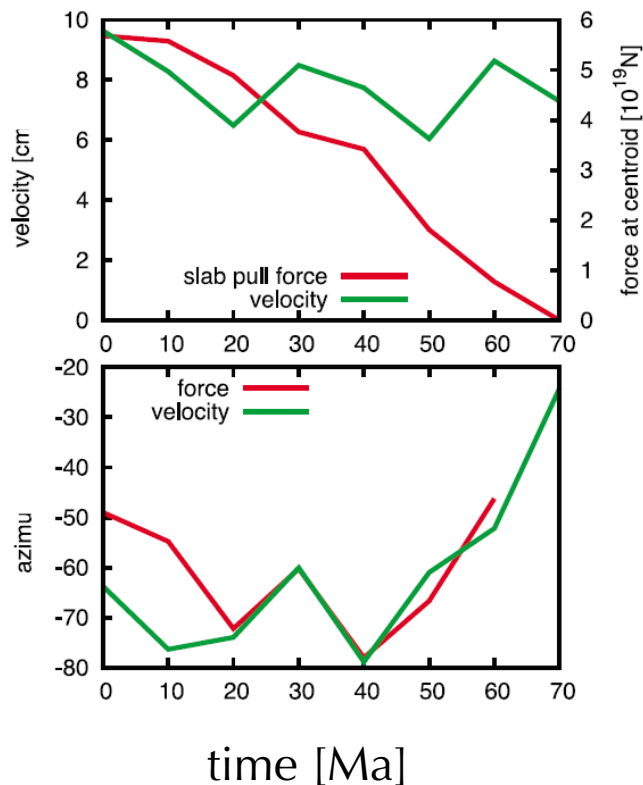
Gerault et al. (2011)



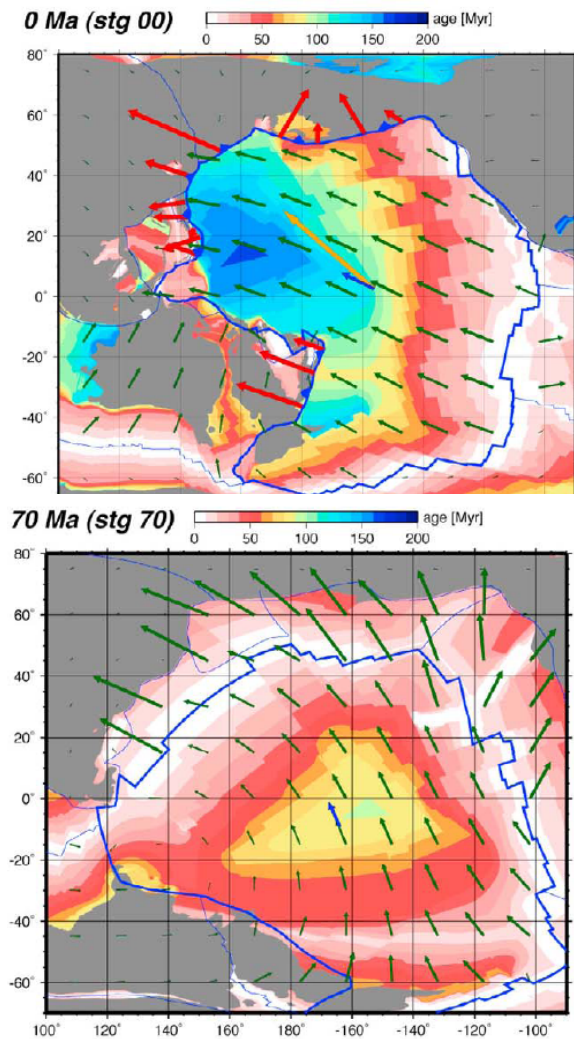
Funiciello et al. (2008)



Upper
mantle
slabs are
not the
whole
story...

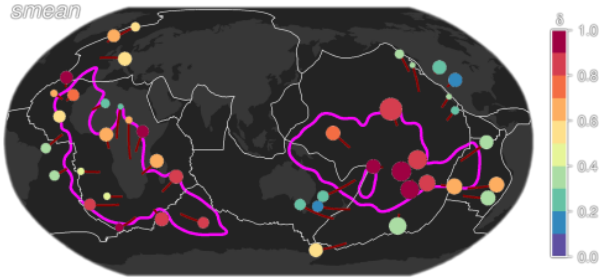


Faccenna et al. (2011)
cf. Gordon

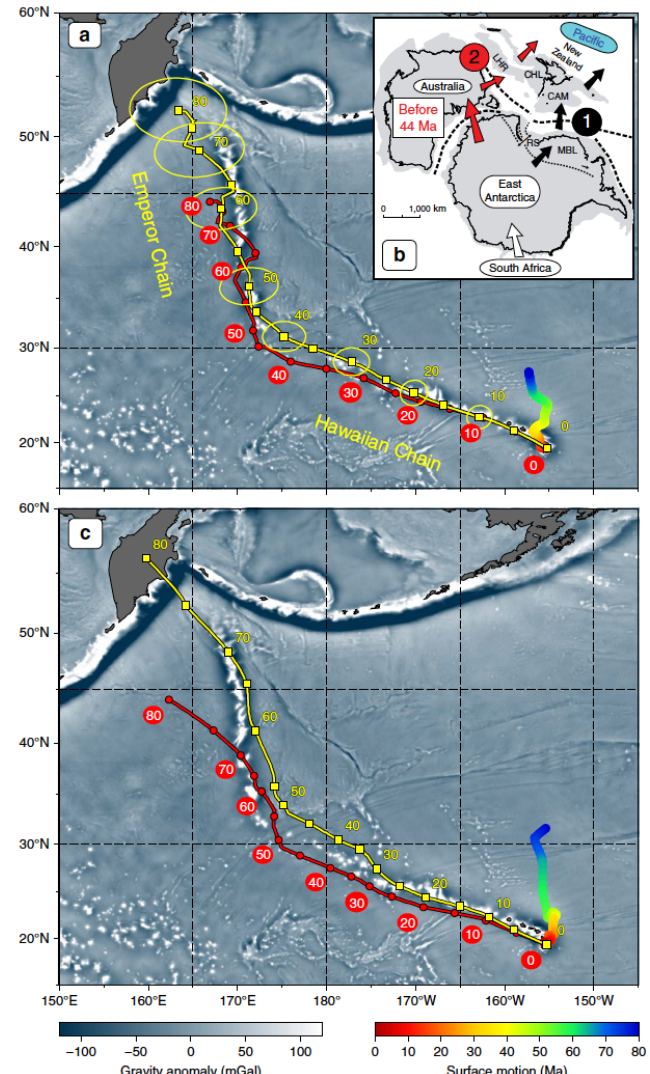


How about plumes?

- Did the plumes or the plates move to cause hotspot track bends?
- Can we predict plume conduits?



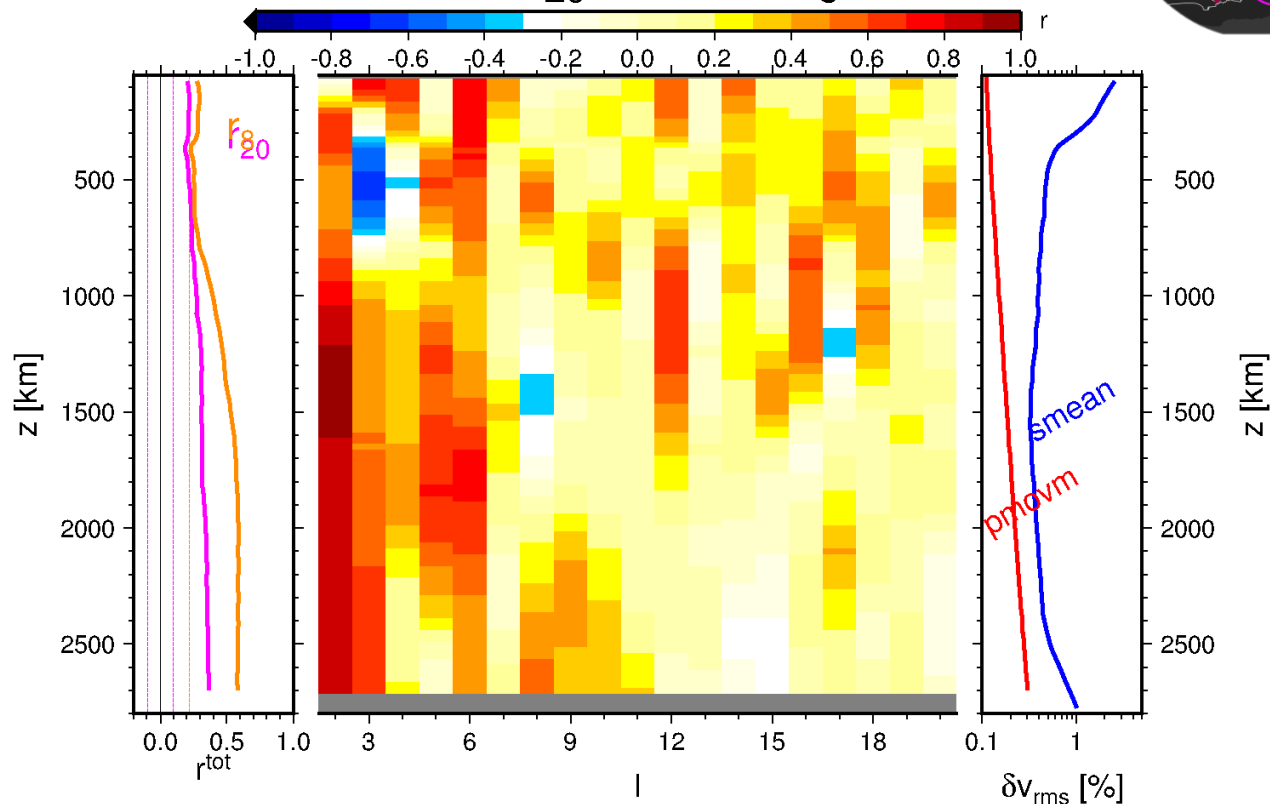
Torsvik et al. (2017)



cf. Steinberger (2000); Tarduno et al. (2009); Hassan et al. (2016)

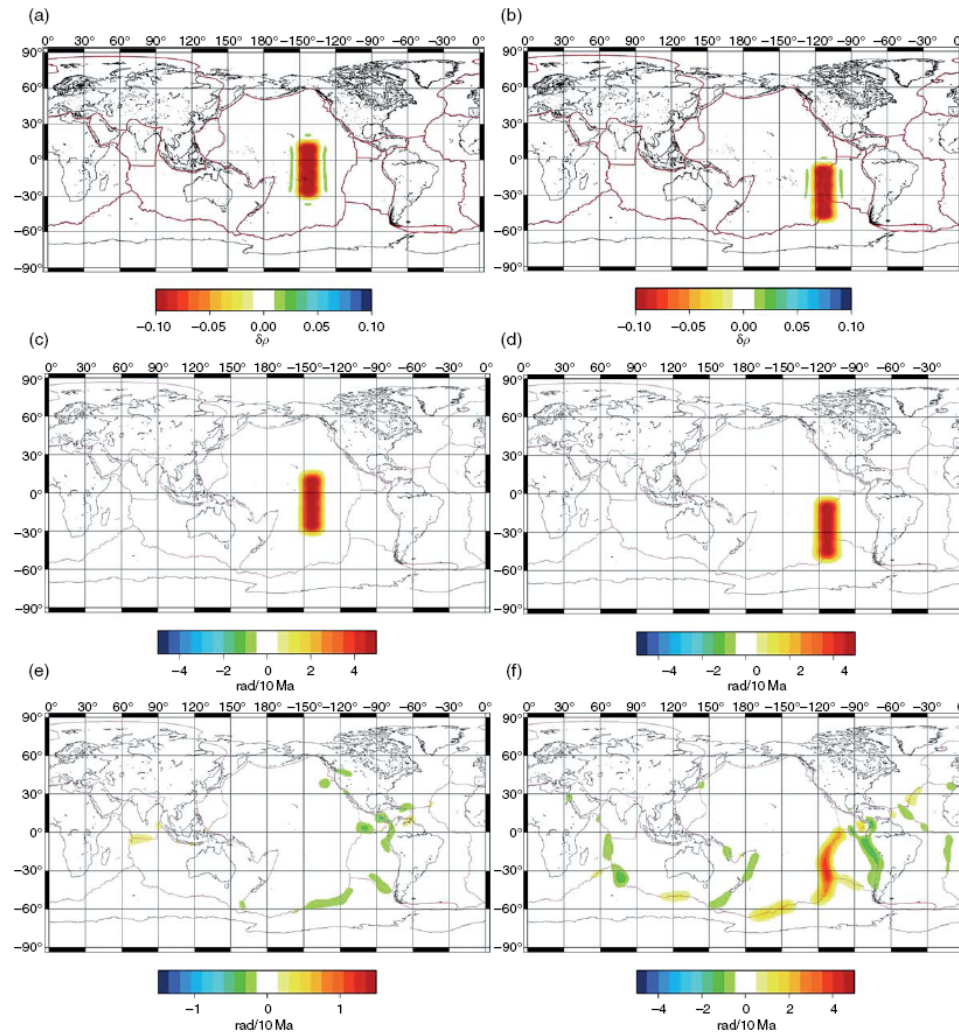
Advected plumes vs. tomography

pmovm vs. smean, $\langle r_{20} \rangle = 0.28$, $\langle r_8 \rangle = 0.42$



- correlation with *moving* plume conduits ~better than with slabs ($r_8 \sim 0.5$ for slabs and plumes)

Plumes as plate driving forces

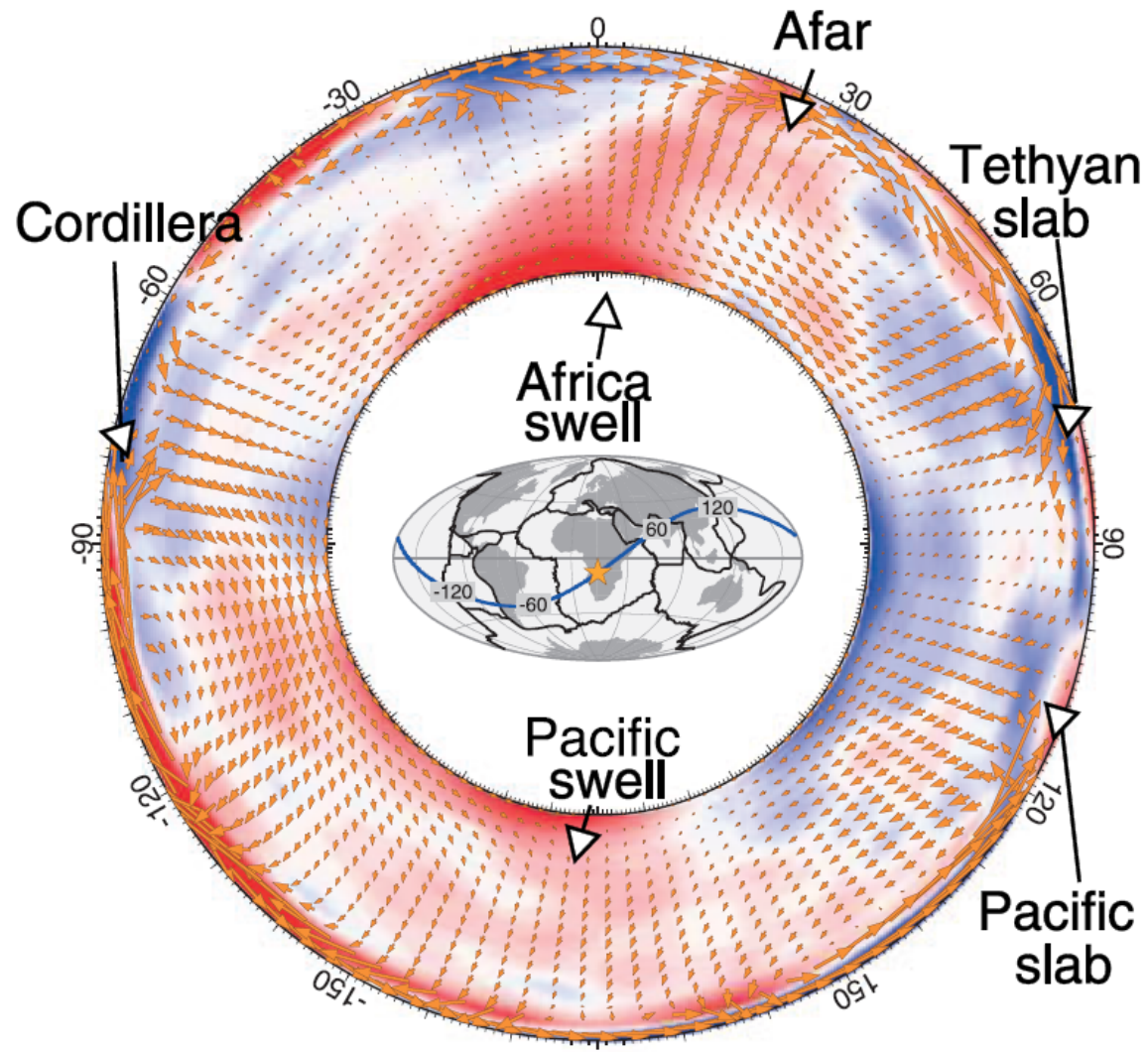


Two density
sources

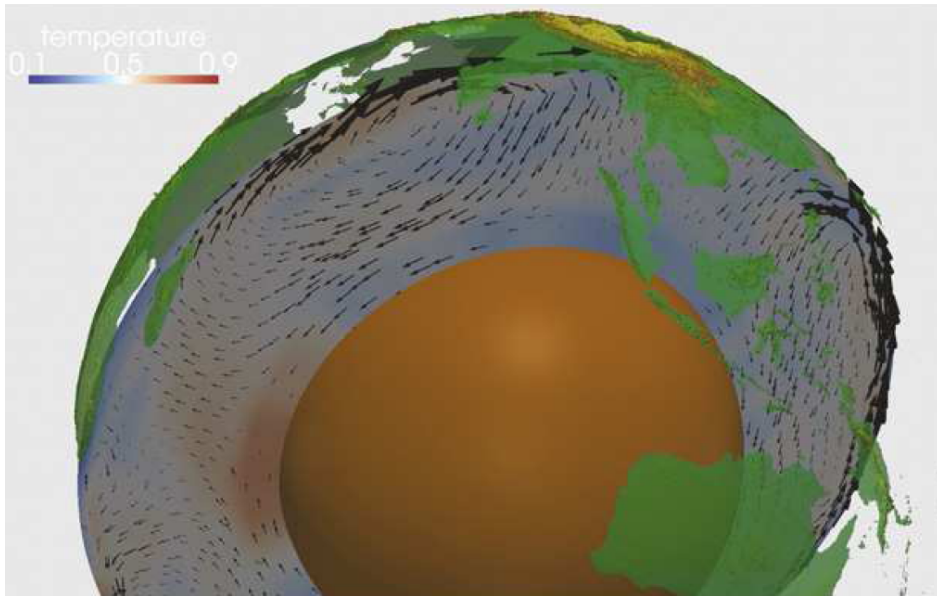
Free-slip solution
(poloidal only)

Plate solution
(poloidal only)

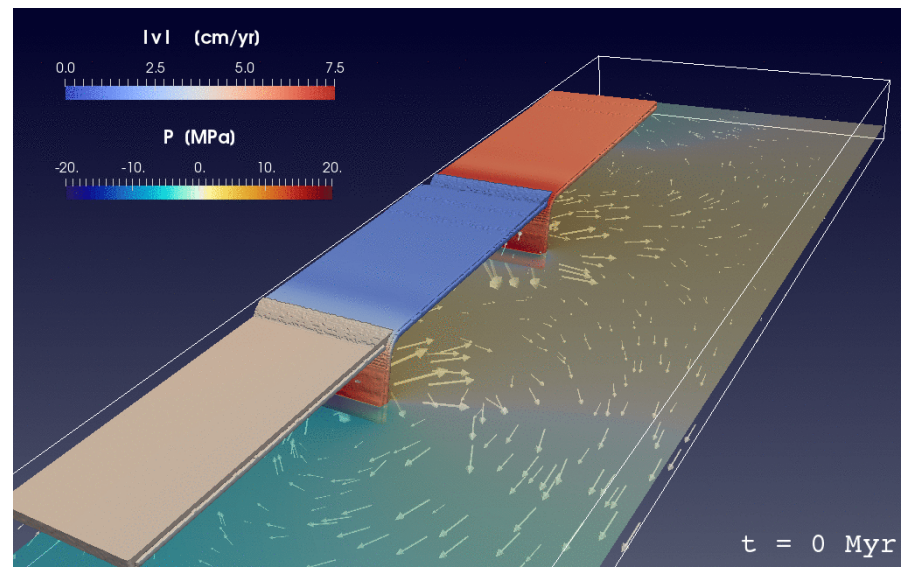
Mantle conveyor belts



India's motion at present:
Mantle conveyor belt,
broad-scale upwelling push?



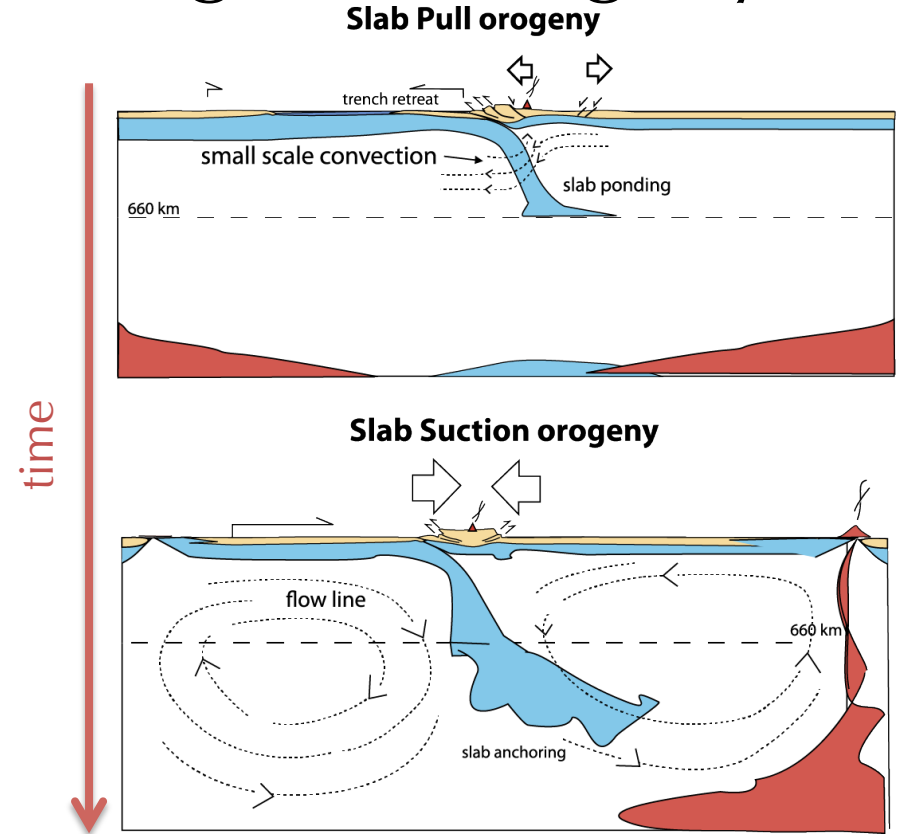
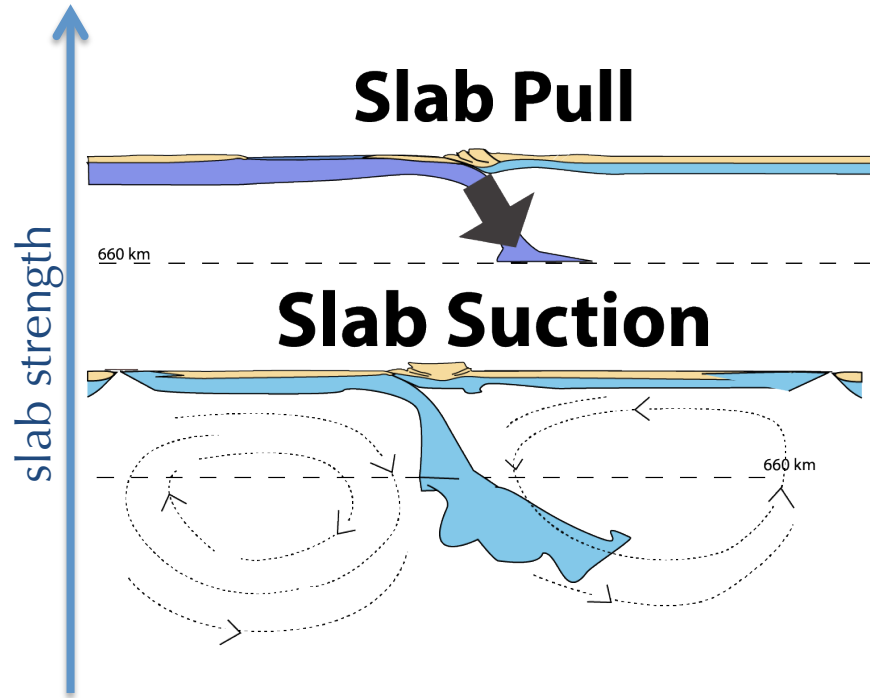
Becker and Faccenna (2011)



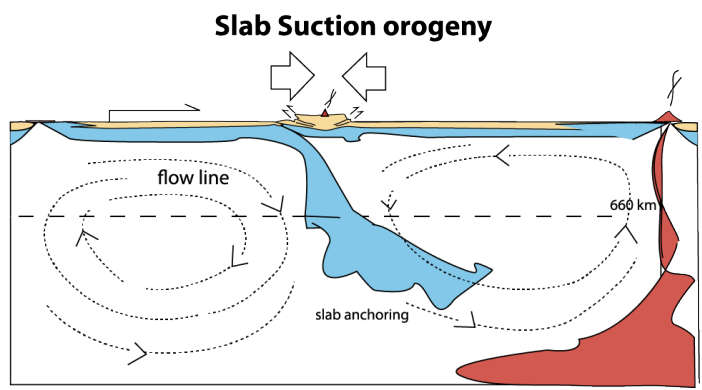
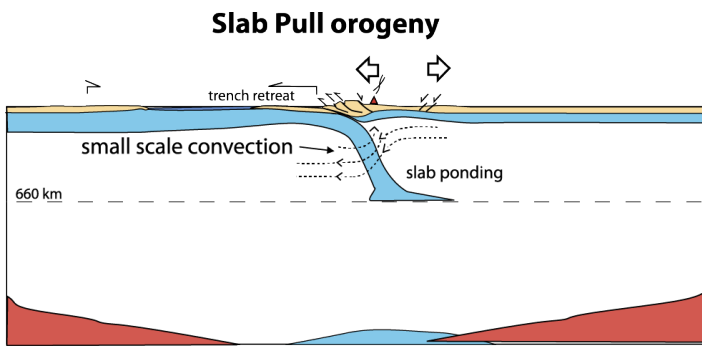
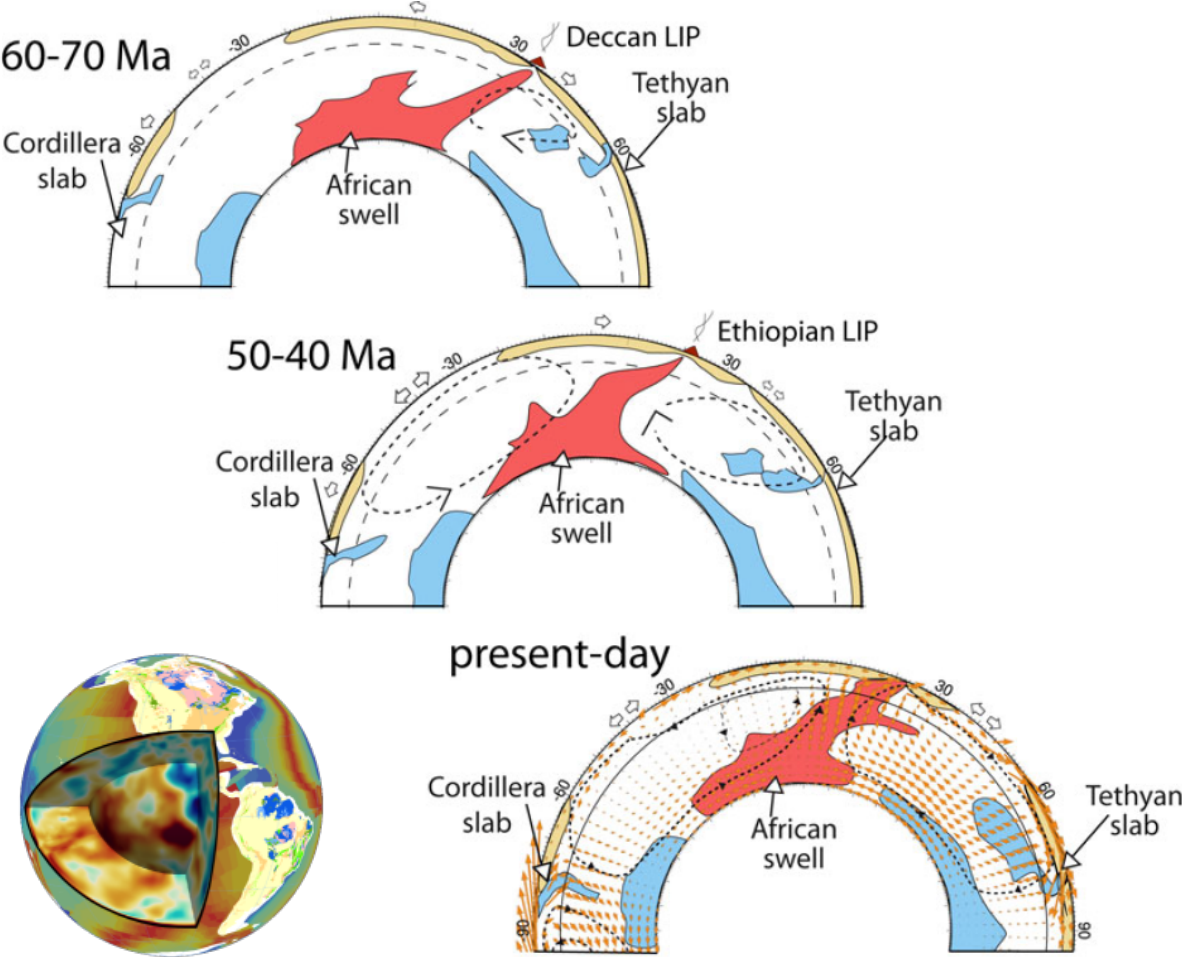
Jagoutz et al. (2015); Holt et al. (2017)

India's motion in the past:
Double slab dynamics?

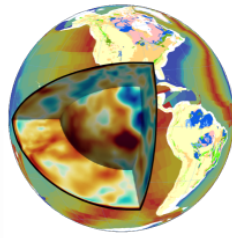
Link between mantle forcing and orogeny



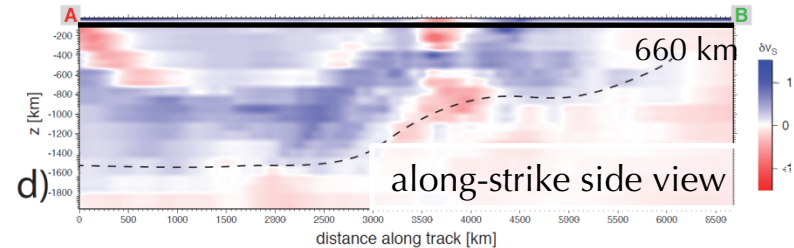
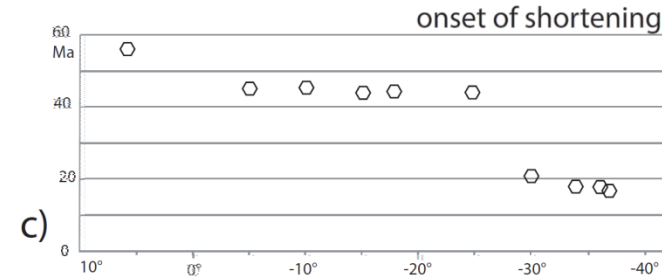
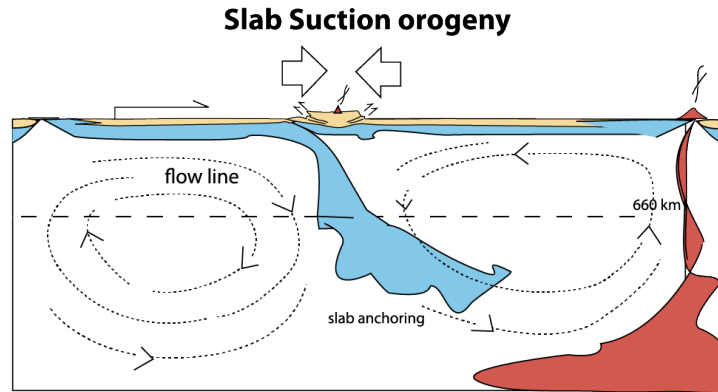
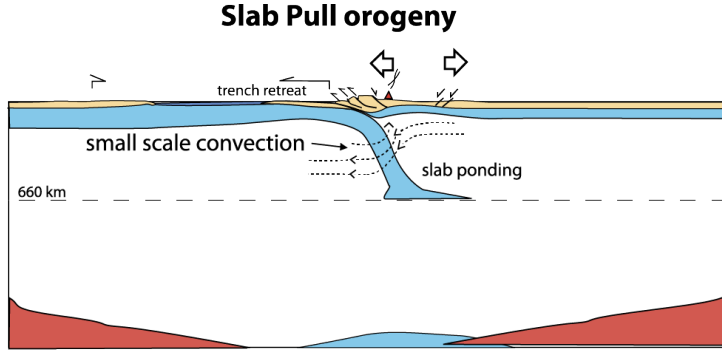
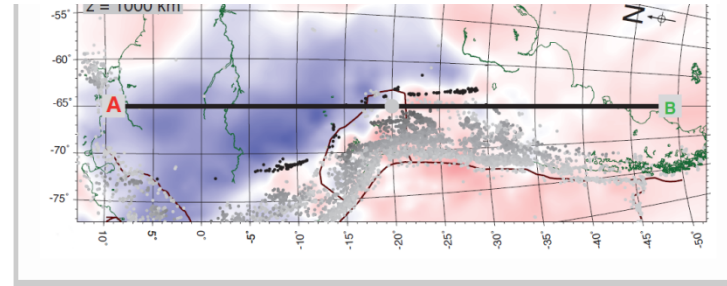
Tibetan orogeny linked to establishment of whole mantle convection cell



Subduction and orogeny: links between convection and continental geology (Andes)



map view of seismic tomography under Andes



Mantle convection

- energy equation governs planetary heat loss
- introduces time-dependence and non-linearity (coupling between velocity and temperature)
- can time reverse advection, but not diffusion

conservation of momentum $\frac{\partial \sigma_{ij}}{\partial x_j} = -Ra\tilde{T}\delta_{ir} - Ra_C\tilde{C}\delta_{ir}$

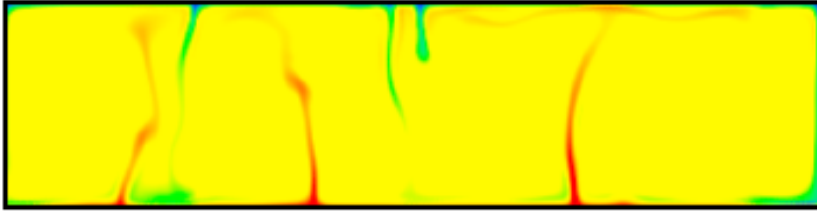
constitutive law $\sigma_{ij} = -p\delta_{ij} + 2\eta(\sigma, T, d, H_2O, \varepsilon)\dot{\varepsilon}_{ij}$

conservation of mass $\frac{\partial v_i}{\partial x_i} = 0$

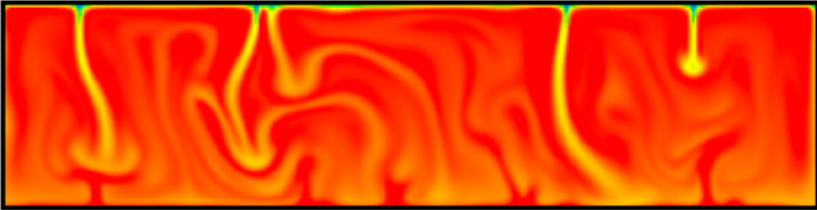
conservation of energy $\frac{\partial T}{\partial t} = \underbrace{-v_i \frac{\partial T}{\partial x_i}}_{\text{advection}} + \underbrace{\kappa \frac{\partial^2 T}{\partial x_j \partial x_j}}_{\text{diffusion}} + \underbrace{H}_{\text{heat generation}}$

Slabs are part of heat transport – effect of heating mode
isoviscous fluid

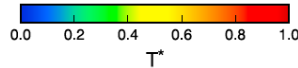
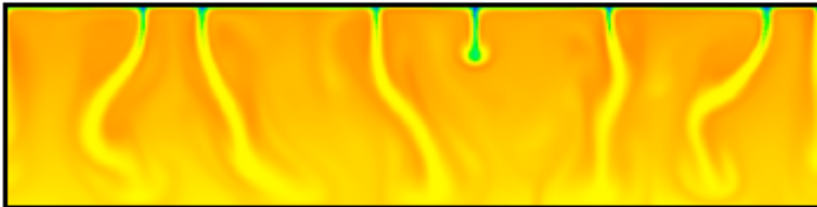
basal



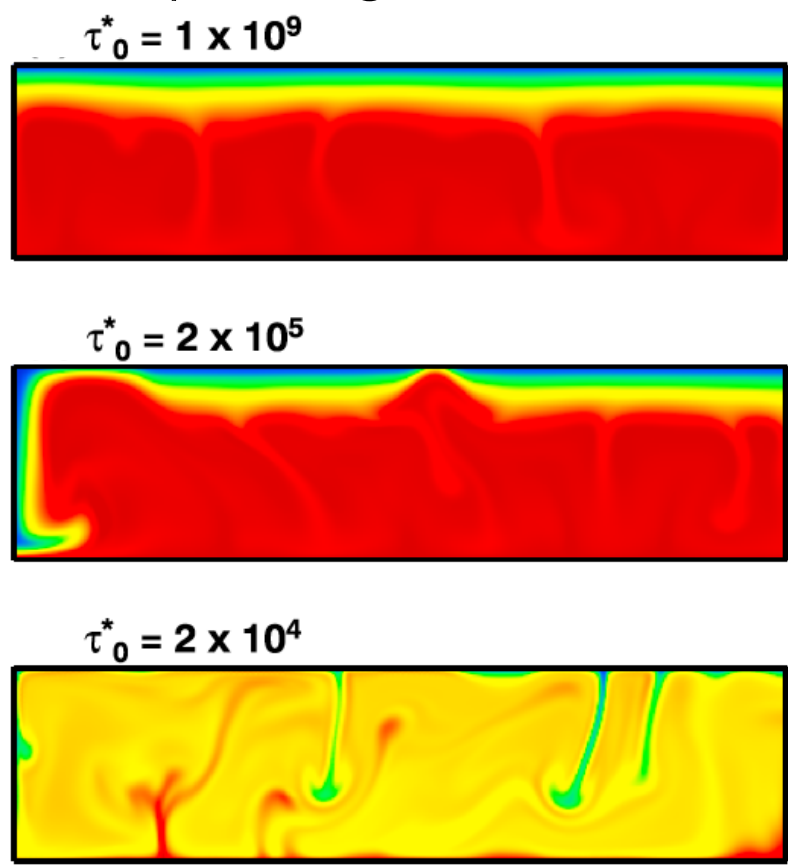
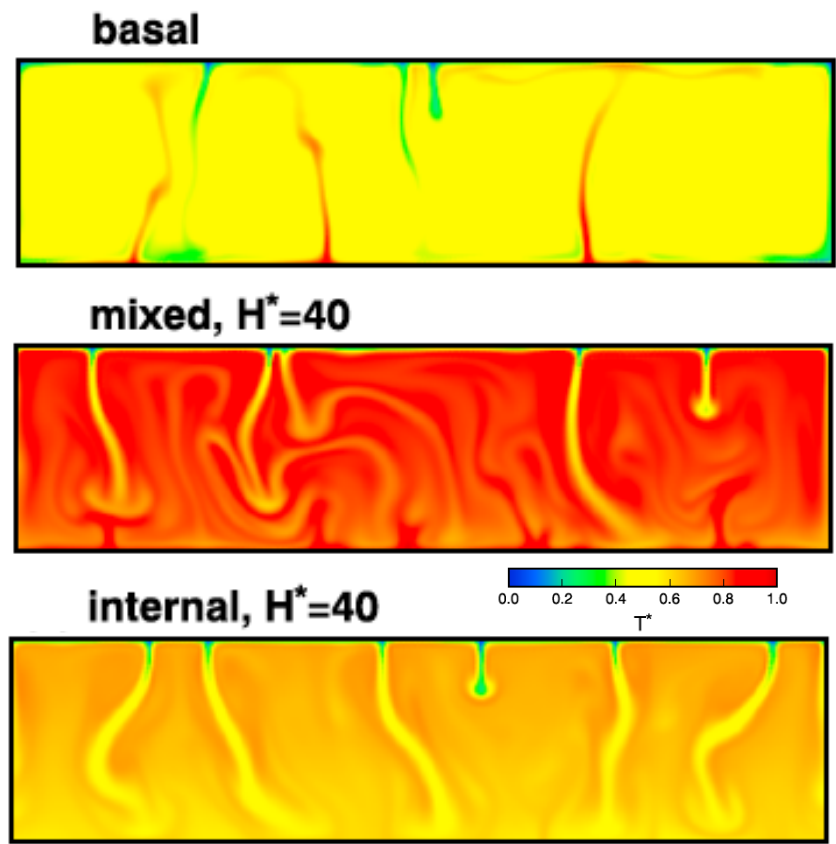
mixed, $H^*=40$



internal, $H^*=40$

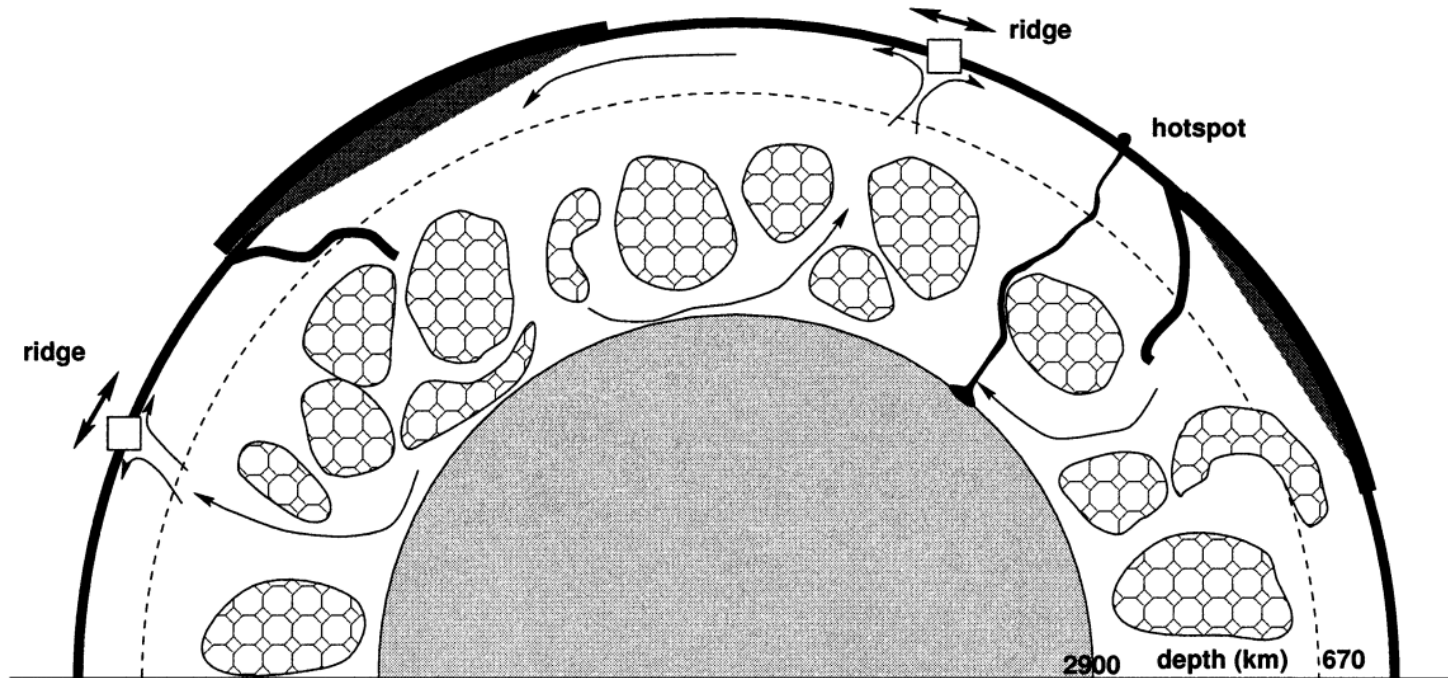


Slabs are part of heat transport – effect of heating and plate mode
isoviscous $\eta(T)$ + yielding (mixed heating)

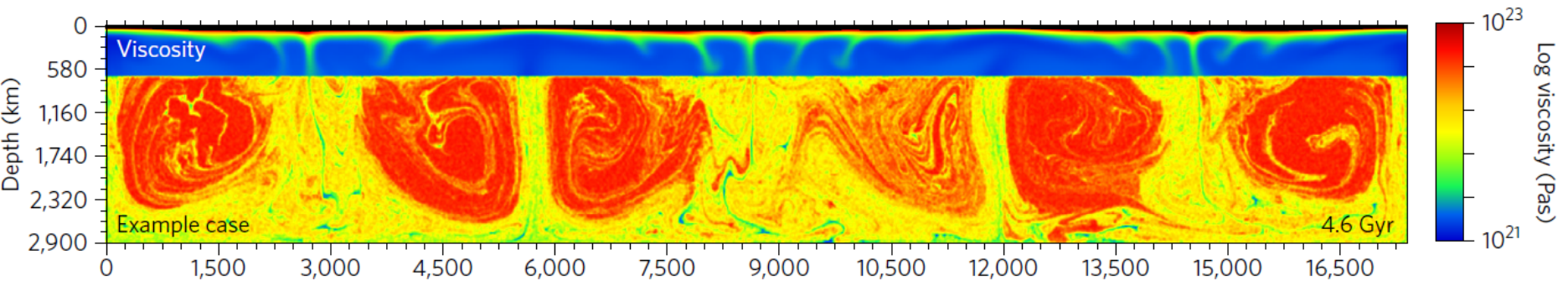
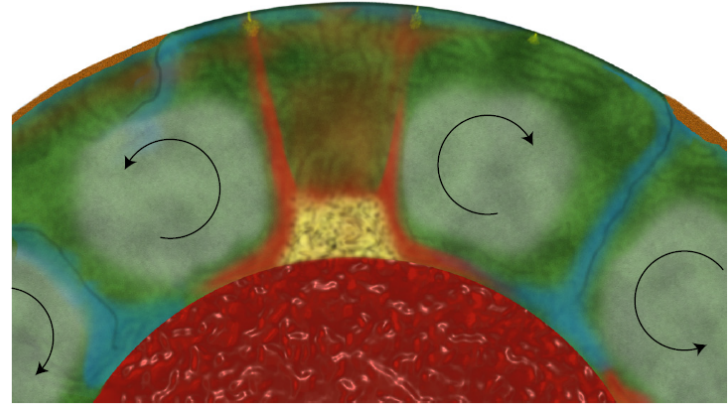
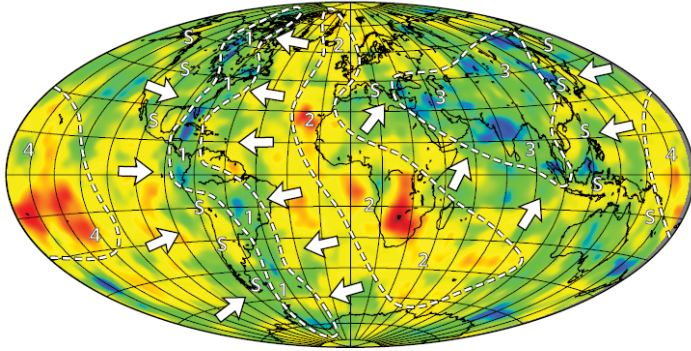


Korenaga (2017), cf. Foley and Becker (2009)

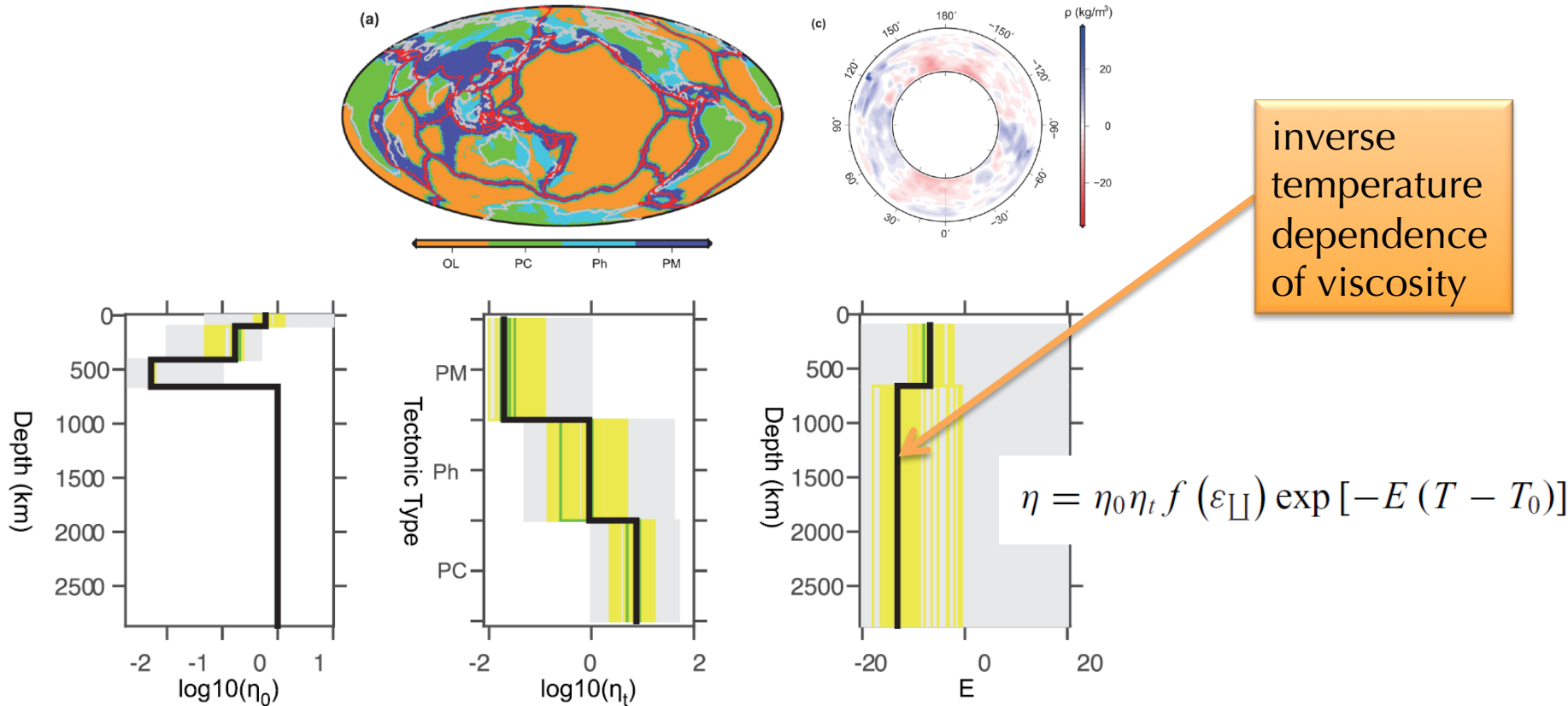
The role of viscous blobs and other (e.g. LPO induced) mechanical anisotropy for convective stability



BEAMs stabilizing conveyor belts?



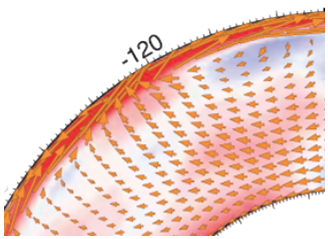
Inversions for lateral viscosity variations



Conclusions

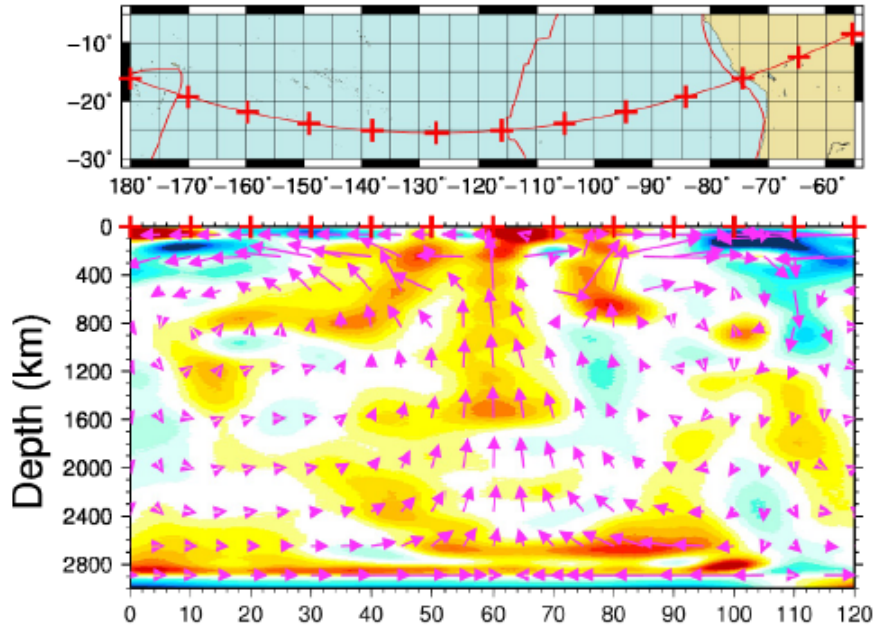
- subduction controls tectonics (including orogeny) and heat loss (by setting mantle convection's spatio-temporal scales)
 - ✧ subduction interface may affect plate velocities in a weak slab world
- potential links between continental dynamics and deep mantle may help decipher planetary evolution
- make progress by integrating diverse datasets in inverse models, e.g. to better constrain mantle and slab rheology
- use tools that capture global, multi-scale interactions of mantle flow for regionally realistic subduction models

Additional slides

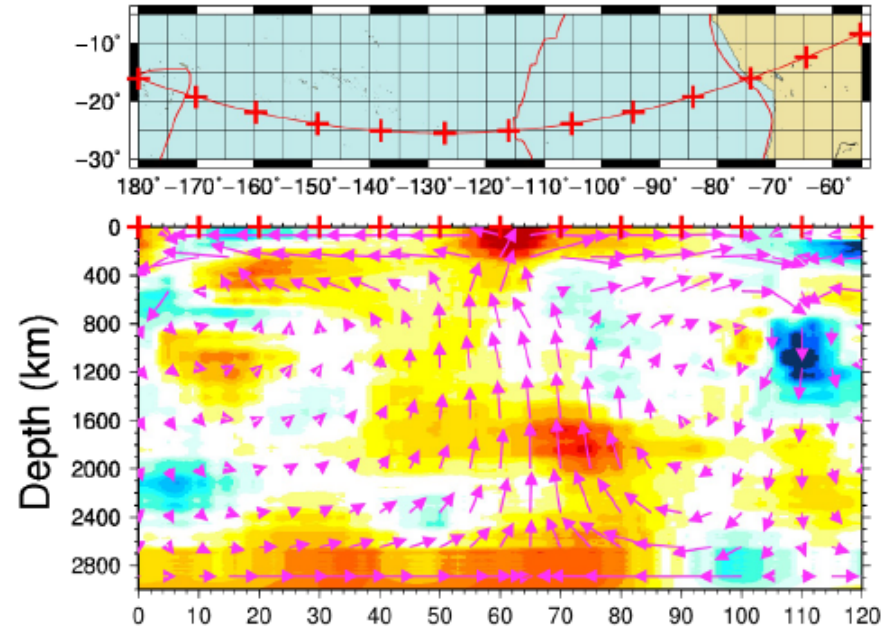


Persistence of upwellings?

A $t = -55$ Ma (B.P.)



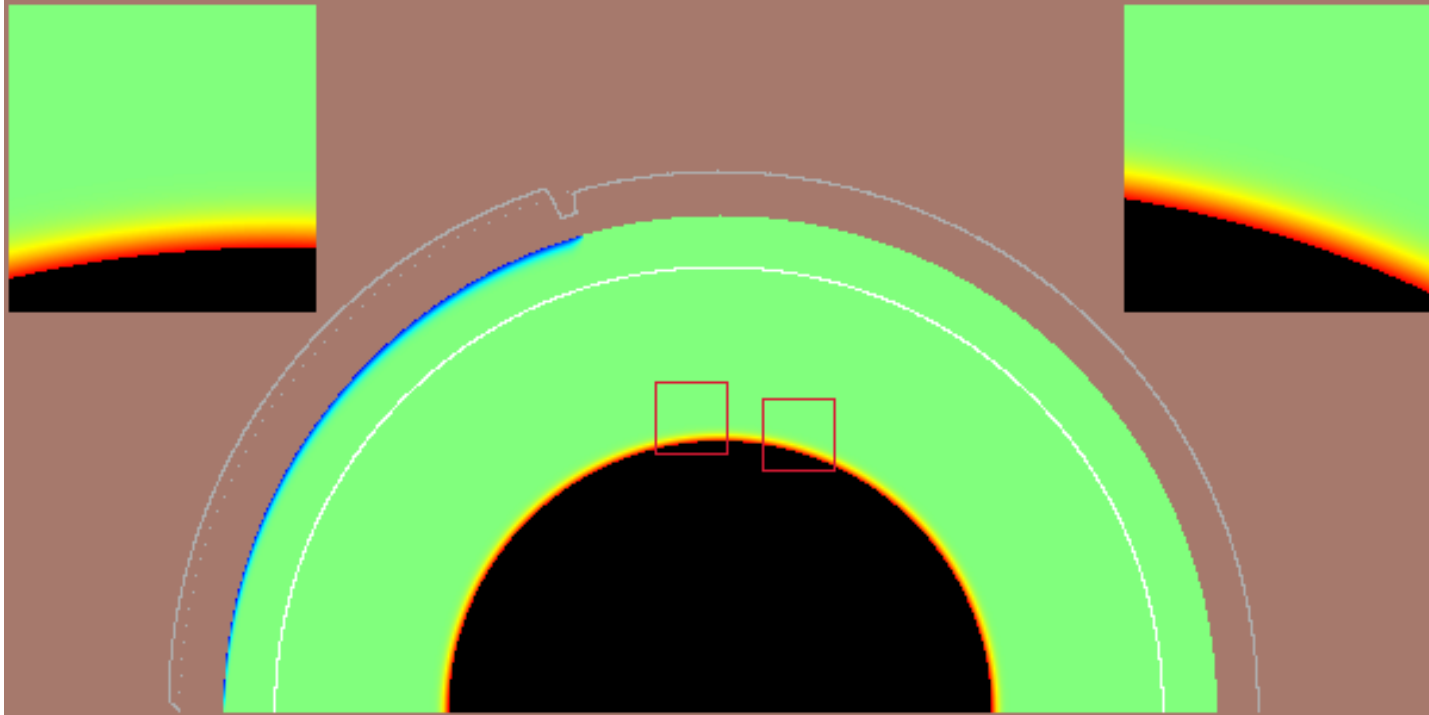
B Present day ($t = 0$ Ma)



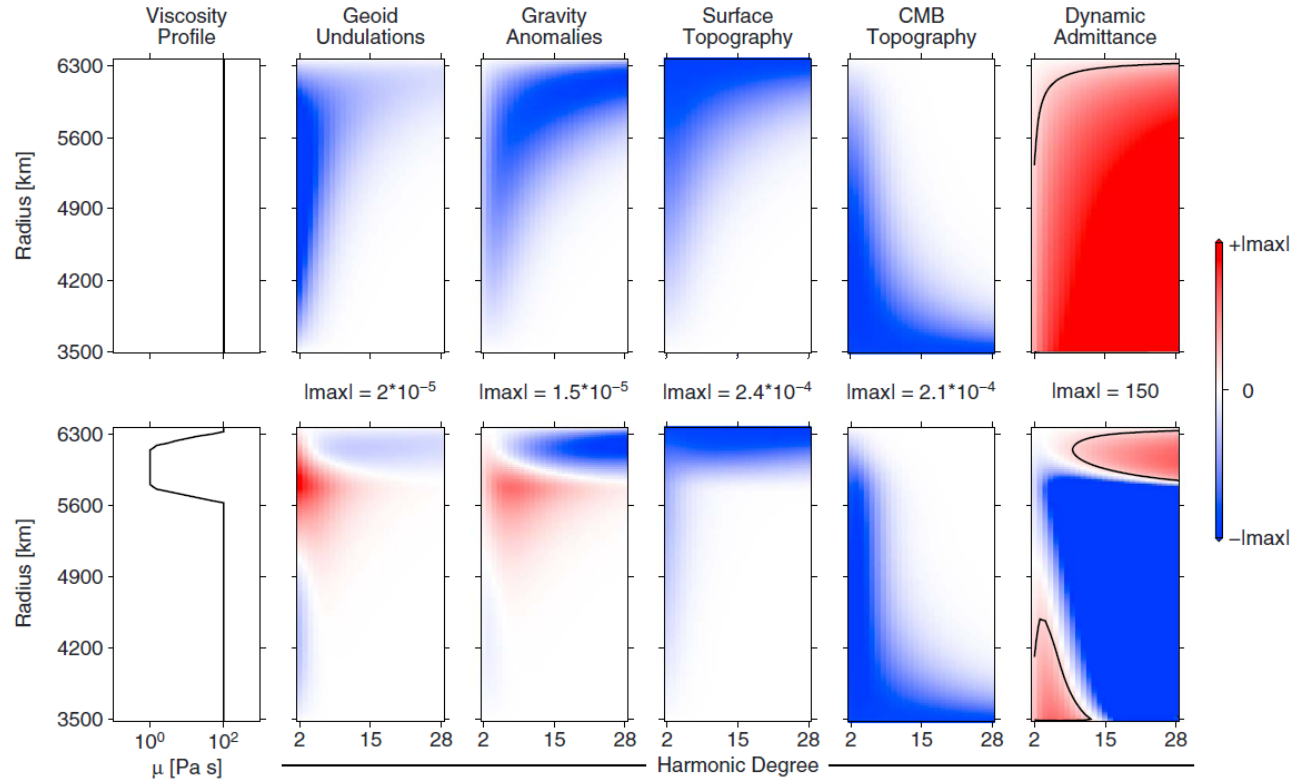
➤ active upwelling underneath East Pacific Rise since 50 Ma?

Rowley et al. (2016)

Ying and yang of up and downwellings

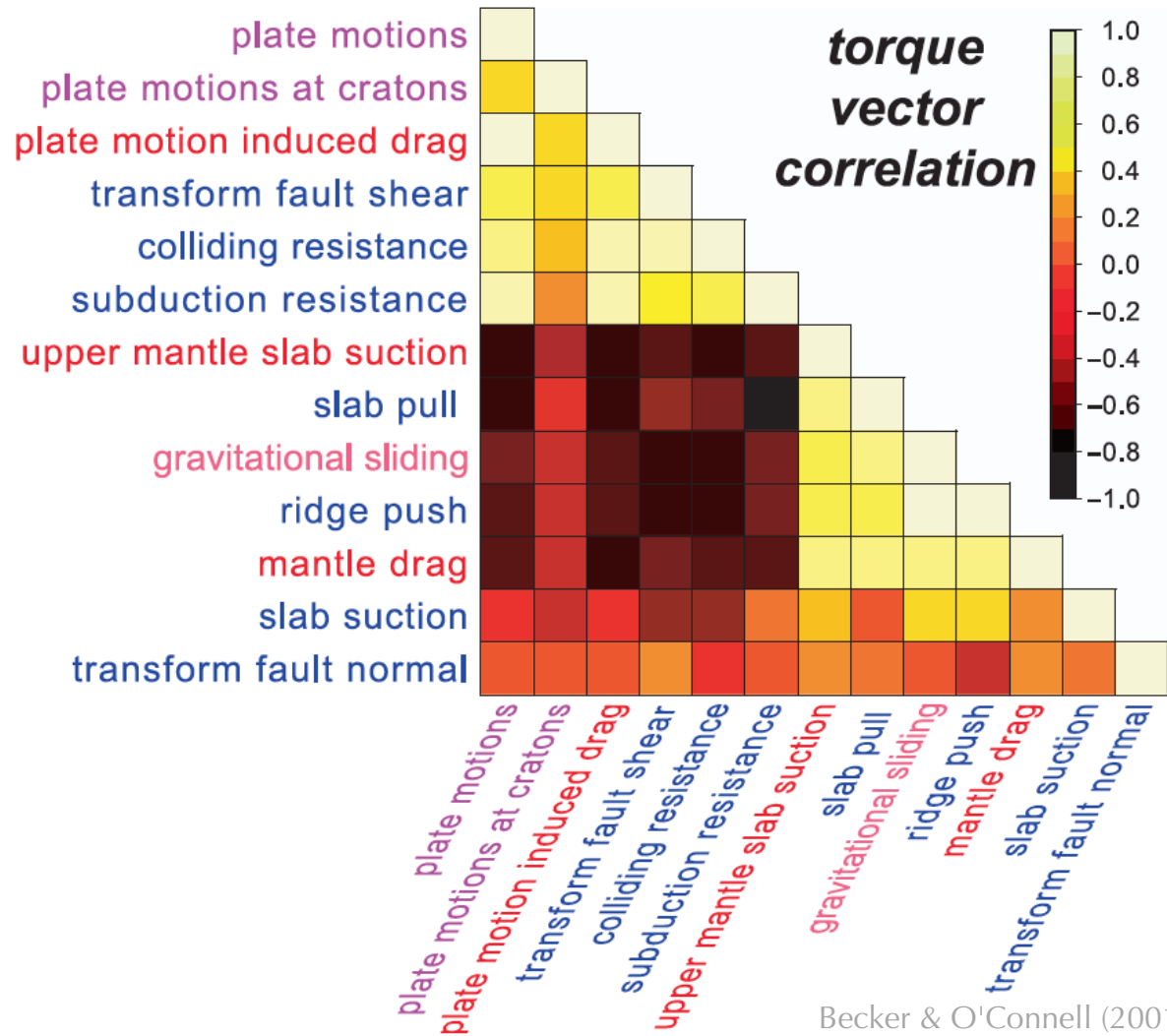


Kernels for layered viscosity mantle



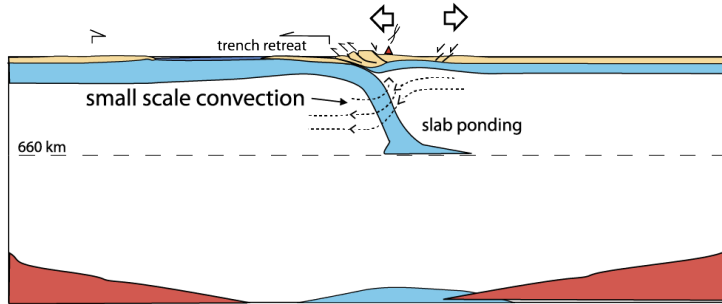
edge forces

mantle tractions

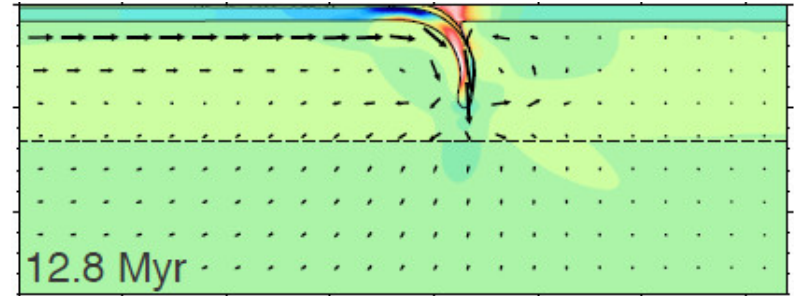
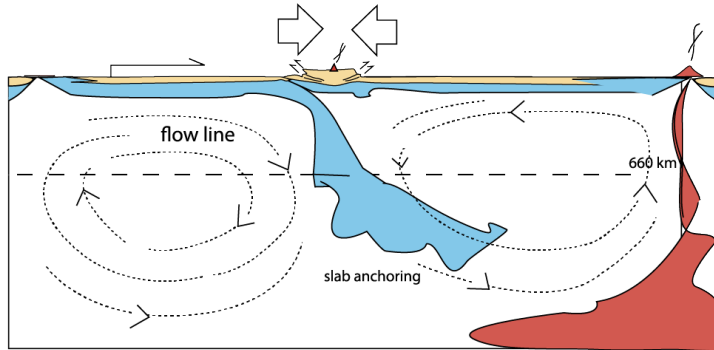


Time-dependent compressional stresses during subduction penetration related to onset of crustal shortening?

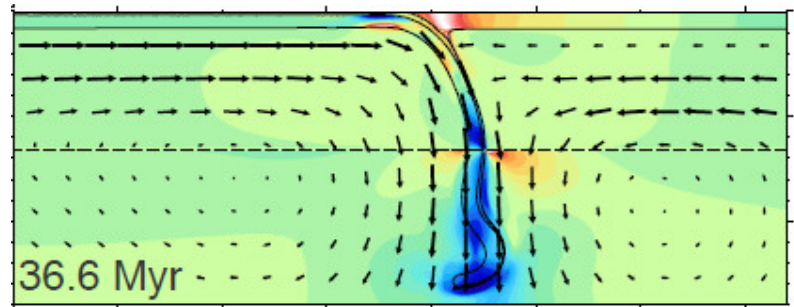
Slab Pull orogeny



Slab Suction orogeny

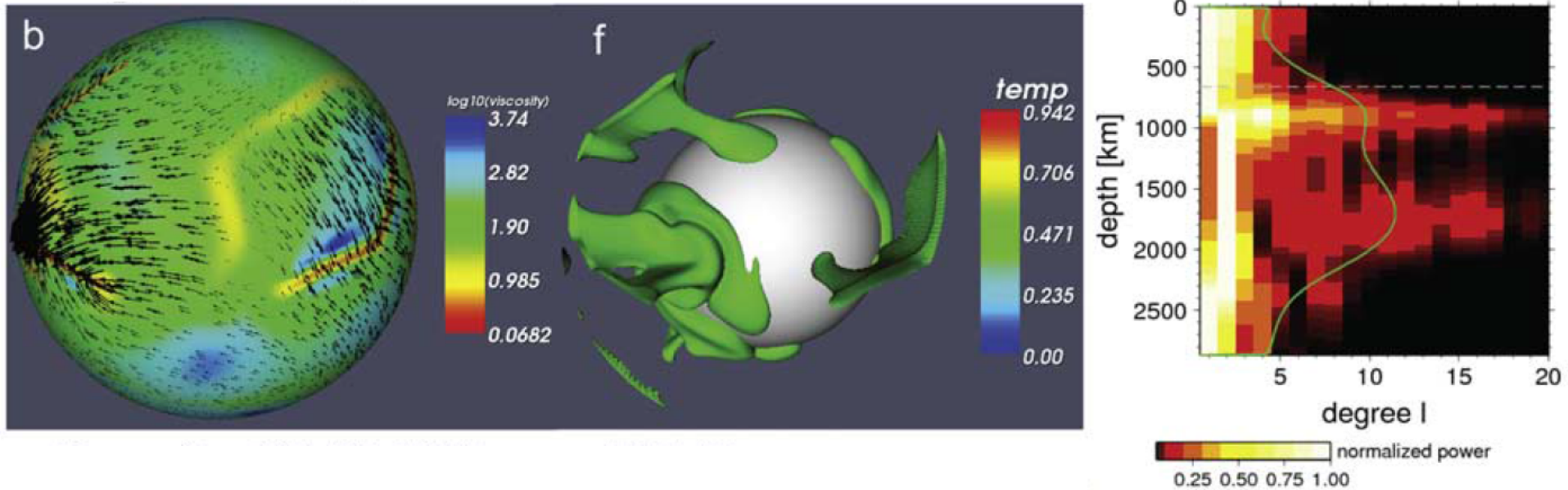


red = horizontal compression
blue = horizontal extension

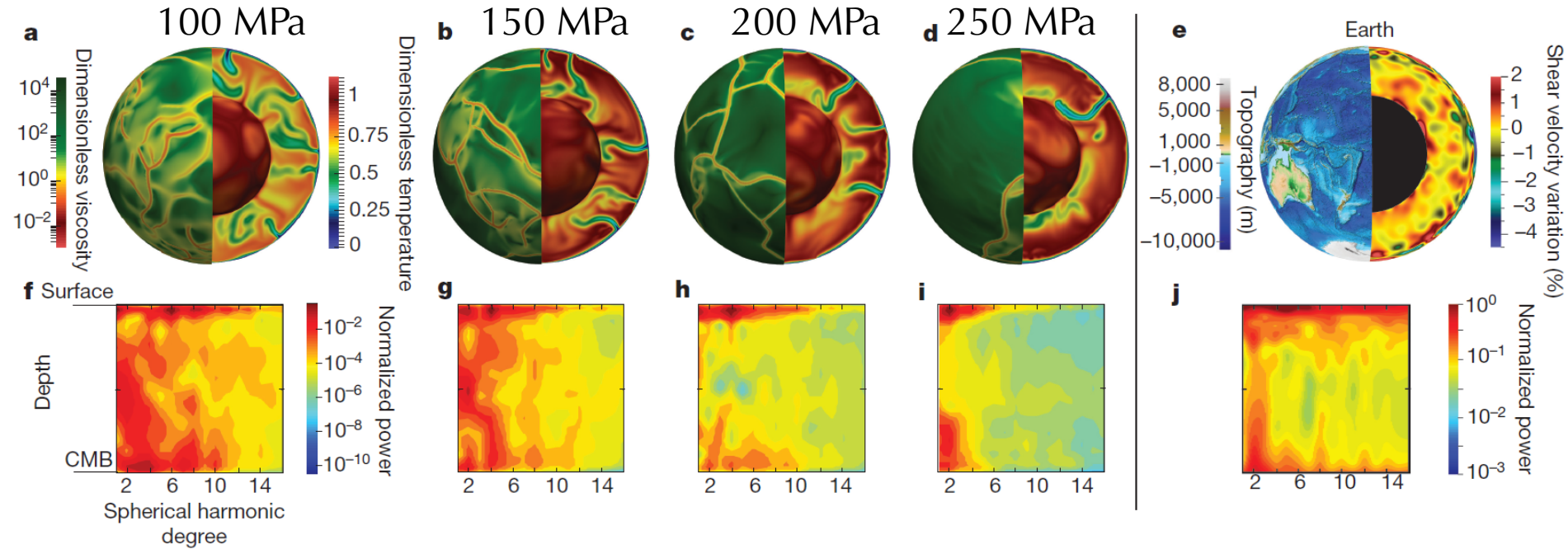


How are the plates made?

$$\eta(T) \text{ \& } \tau_{\text{yield}}, \eta(z)/f_{\text{melt}}, \Delta\rho_c, \eta(\phi)$$



Slabs control convection: Effect of yield stress on plateness



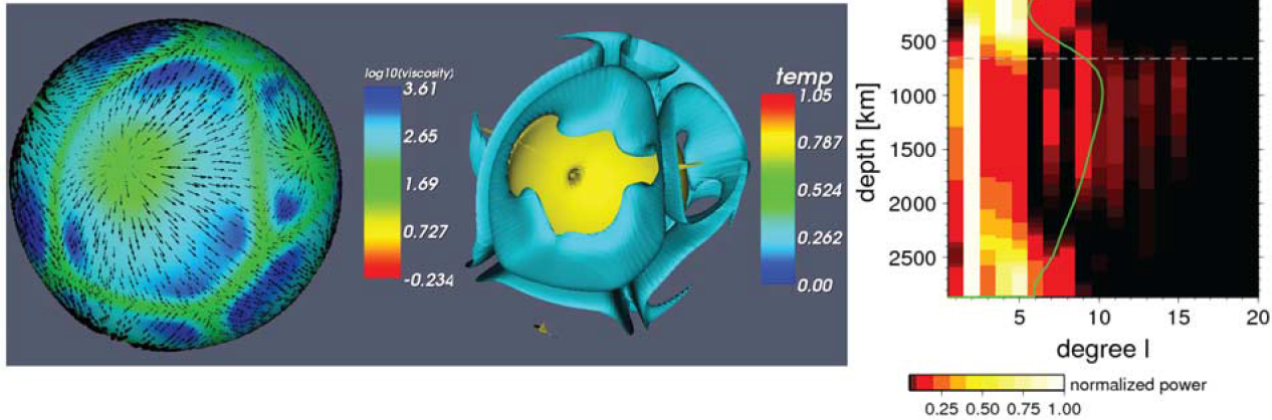
Mallard et al. (2016)

cf. Tackley (2000a,b), Richards et al. (2001),
van Heck and Tackley (2008), Foley and Becker (2009)

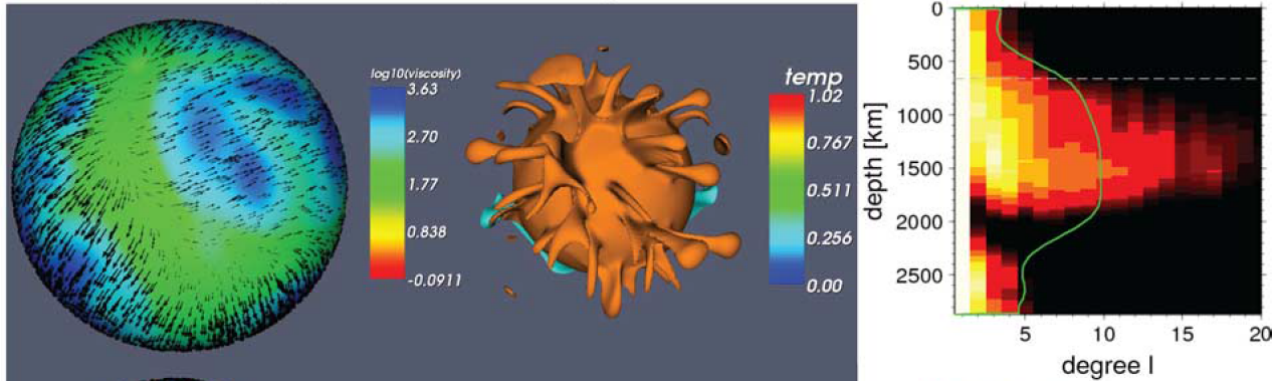
But: effect of asthenosphere,
internal vs. bottom heating, Ra #,
continents, and damage/memory

Plumes breaking plates

a. Pure bottom heating

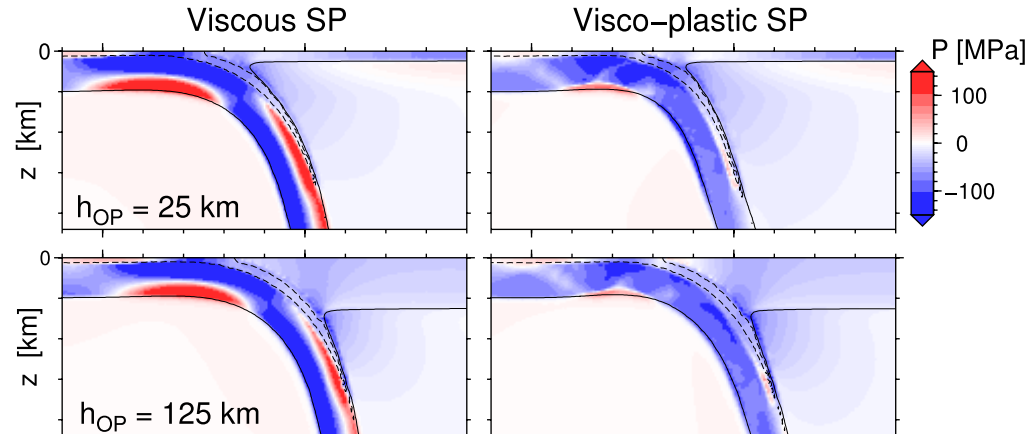
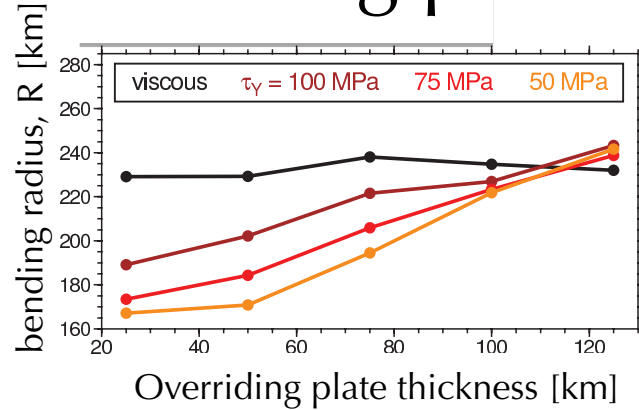


b. Bottom heating, 60 % internal heating

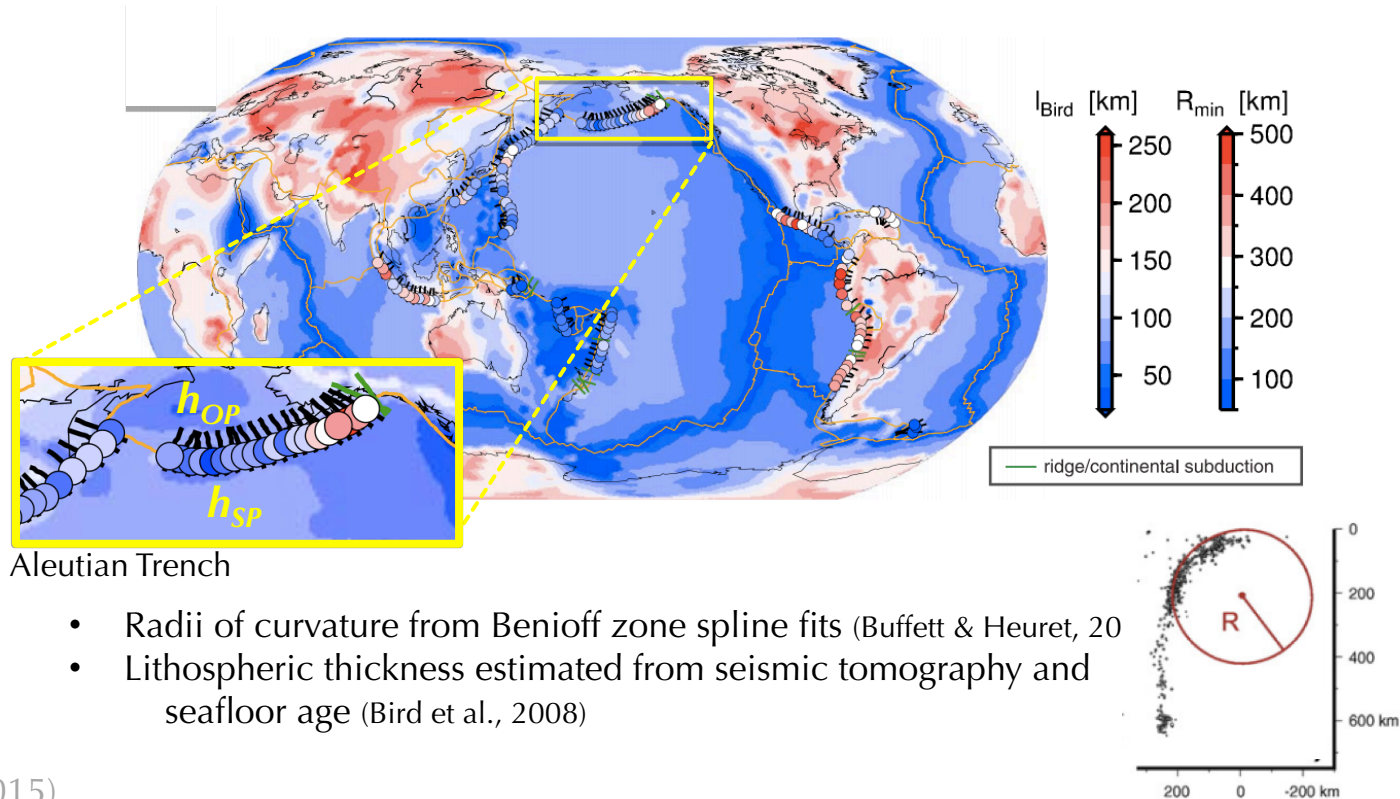


Effectively plastic slab: Bending radius affected by overriding plate thickness

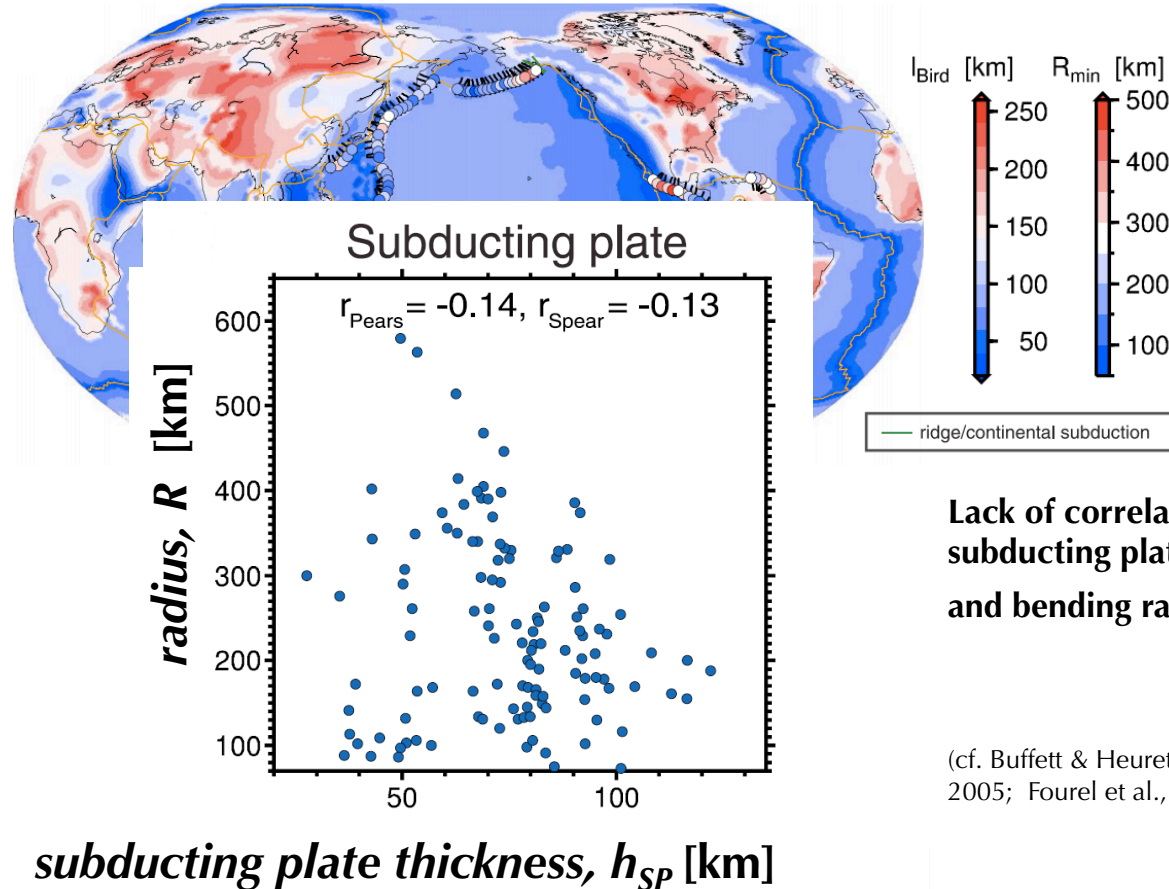
Plastic slab more strongly affected by lifting force ($F_{\Delta P}$) associated with overriding plate, weak bending stresses



Bending radius and plate thickness in nature



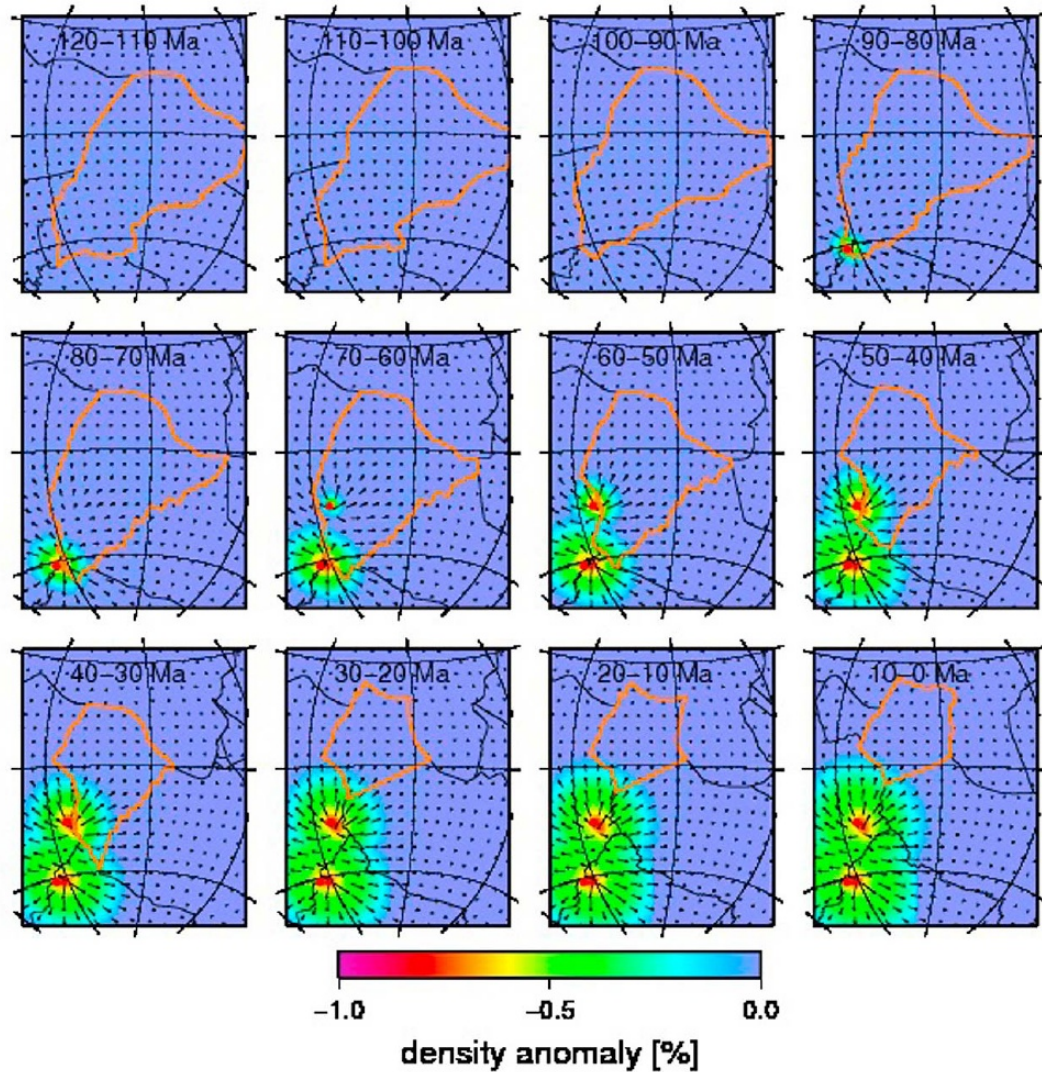
No subducting plate control on slab curvature



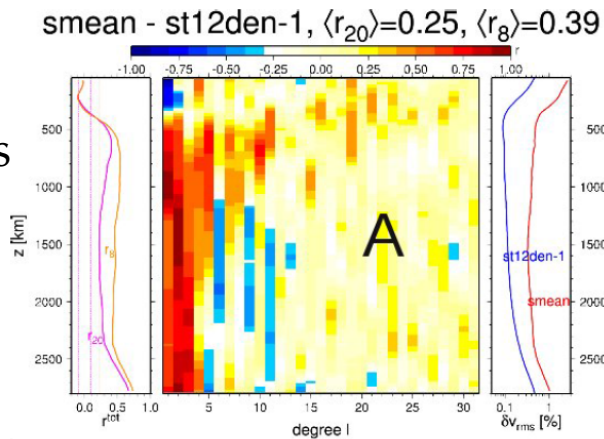
**Lack of correlation between
subducting plate thickness, h_{SP}
and bending radius, R**

(cf. Buffett & Heuret, 2011; Cruciani et al., 2005; Fourel et al., 2014).

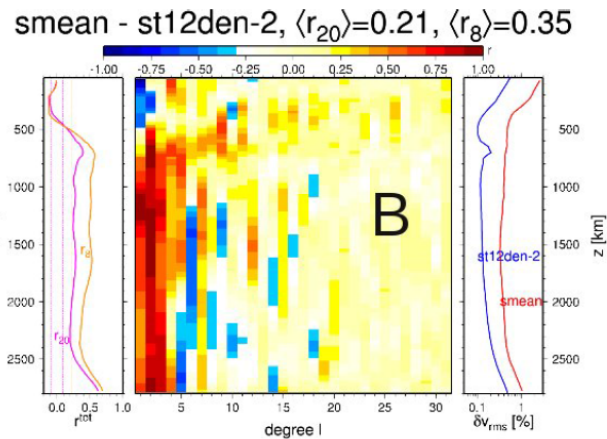
Plate driving forces: plumes



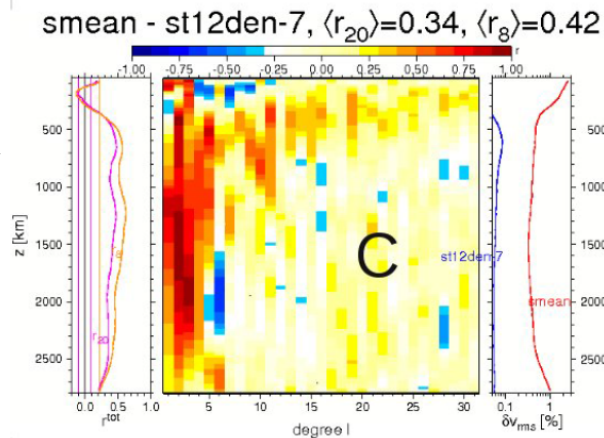
Slab plus
piles I



Slab plus
piles II



Slabs only



Slabs and
plumes

